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Understanding process controls on groundwater recharge variability across Africa through Recharge Landscapes

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Abstract

Groundwater is critical in supporting current and future reliable water supply throughout Africa. Although continental maps of groundwater storage and recharge have been developed, we currently lack a clear understanding on how the controls on groundwater recharge vary across the entire continent. Reviewing the existing literature, we synthesize information on reported groundwater recharge controls in Africa. We find that 15 out of 22 of these controls can be characterised using global datasets. We develop 11 descriptors of climatic, topographic, vegetation, soil and geologic properties using global datasets, to characterise groundwater recharge controls in Africa. These descriptors cluster Africa into 15 Recharge Landscape Units for which we expect recharge controls to be similar. Over 80% of the continents land area is organized by just nine of these units. We also find that aggregating the Units by similarity into four broader Recharge Landscapes (Desert, Dryland, Wet tropical and Wet tropical forest) provides a suitable level of landscape organisation to explain differences in ground-based long-term mean annual recharge and recharge ratio (annual recharge / annual precipitation) estimates. Furthermore, wetter Recharge Landscapes are more efficient in converting rainfall to recharge than drier Recharge Landscapes as well as

27 having higher annual recharge rates. In Dryland Recharge Landscapes, we found that annual
28 recharge rates largely varied according to mean annual precipitation, whereas recharge ratio
29 estimates increase with increasing monthly variability in P-PET. However, we were unable to
30 explain why ground-based estimates of recharge signatures vary across other Recharge
31 Landscapes, in which there are fewer ground-based recharge estimates, using global datasets
32 alone. Even in dryland regions, there is still considerable unexplained variability in the
33 estimates of annual recharge and recharge ratio, stressing the limitations of global datasets for
34 investigating ground-based information.

35 **Keywords:** Groundwater recharge, Africa, recharge controls, ground-based estimates,
36 landscapes, comparative hydrology

37 **1 Introduction**

38 With an estimated storage of 0.66 million km³, groundwater is the largest store of freshwater
39 in Africa, exceeding annual volumes of streamflow by a factor of 100, and its development is
40 fundamental for securing current and future water supply (MacDonald et al., 2012).
41 Understandably it is often regarded as important for both urban (Foster et al. 2020; Oiro et al.
42 2020) and rural (Calow et al. 1997; Lapworth et al. 2013; MacDonald and Calow
43 2009) communities, though quantifying its role in water supply remains challenging for many
44 parts of the continent (Chávez García Silva et al. 2020). In Northern Africa, groundwater
45 irrigation practices (Siebert et al. 2010) have potentially led to the depletion of groundwater
46 resources (Aeschbach-Hertig and Gleeson 2012; Rodell et al. 2018). In contrast, groundwater
47 use in sub-Saharan Africa is primarily for domestic supply and sanitation services (Braune
48 and Xu 2010) and is potentially being under-utilized for crop production (Giordano 2006;
49 Siebert et al. 2010). In dryland river basins, securing water supply from surface water is
50 challenging due to the high inter-annual variability of streamflow and persistent dry periods

51 (Conway et al. 2009; Siam and Eltahir 2017; Sidibe et al. 2019). Therefore increasing
52 groundwater abstraction for more conjunctive use of surface water and groundwater could
53 reduce vulnerability to climate driven surface water shortages, particularly in rural
54 communities (Calow et al., 1997; Lapworth et al., 2013; MacDonald & Calow, 2009) and
55 generally improve water accessibility (Robins et al., 2006). In the future, rapid population
56 growth (Gerland et al. 2014; Parnell and Walawege 2011) and climate change could further
57 enhance the value of African groundwater resources (Taylor et al. 2013).

58

59 Yet, our understanding of the spatial variability of groundwater recharge processes across
60 Africa remains limited, constraining our ability to plan for the sustainable use of this resource
61 (MacDonald et al., 2021), though recharge rates should not be regarded as a safe-yield for
62 groundwater use (Aeschbach-Hertig and Gleeson 2012). Recent studies have tried to
63 overcome this problem in multiple ways: [1] Scaling up knowledge from a limited number of
64 detailed local studies. Cuthbert *et al.* (2019b) used multi-decadal groundwater level
65 timeseries in conjunction with local knowledge to develop site specific conceptual models
66 which allowed the authors to highlight a relationship between climate and recharge
67 frequency, sensitivity to precipitation and dominant recharge mechanisms. However, this
68 approach relies heavily upon rare long-term data as well as local knowledge and therefore it
69 is challenging to transfer findings to larger scales or different regions. [2] Recently,
70 MacDonald et al. (2021) used 134 ground-based annual recharge estimates compiled from
71 the literature along with global datasets to develop a continental statistical model. This model
72 enabled them to estimate long-term groundwater recharge rates across Africa using mean
73 annual precipitation without qualitative inclusion of different recharge processes. [3] Most
74 studies have based their continental scale estimates on process-based models. Global scale
75 hydrological models and land surface models can estimate groundwater recharge rates across

76 large spatial domains (Reinecke et al. 2021). However, recharge outputs from these models
77 have not yet been thoroughly evaluated (Bierkens, 2015; Telteu *et al.*, 2021; Wagener *et al.*,
78 2021). Furthermore, global models thus far only include a limited number of process
79 representations and neglect regionally dominant controls which could be important for
80 Africa, such as karst (Hartmann et al., 2015; Hartmann et al., 2014) or dryland-specific
81 hydrological processes (Quichimbo et al. 2021). This is likely because most (if not all)
82 available continental to global scale models available to estimate recharge have their origin
83 from regions outside of Africa.

84 In this study we want to take a step back to review what dominant controls should be present
85 in a model across Africa and investigate how well we can quantify these process controls
86 given the available data. In doing so, We specifically aim to answer three questions: (i) What
87 are the dominant controls on groundwater recharge already identified across Africa in
88 previous studies? (ii) Using global datasets only, what descriptors of controlling processes
89 can we define, and which regions of Africa should have similar recharge controls when
90 clustered using these descriptors? (iii) How do these regions for which we expect similar
91 controls compare to ground-based recharge observations? Due to the limited amount of
92 ground-based data on groundwater recharge in Africa, we adopt an approach which builds
93 strongly on our a priori understanding of recharge controls in Africa identified from the
94 literature. In doing so we build on previous efforts by Scanlon et al. (2006) who synthesized
95 qualitative local knowledge of recharge processes for the world's dry regions. In keeping
96 with the database compiled by (MacDonald et al., 2021), we only review the controls on
97 recharge which is distributed throughout the landscape. MacDonald et al. (2021) define
98 distributed recharge as both diffuse and focussed recharge but exclude focussed recharge
99 from large discrete features such as rivers or lakes. Where focussed recharge is widely
100 distributed through ephemeral rivers, depressions or rock fractures which are common over a

101 large area and contribute to regional recharge, they include this in their definition of
102 distributed recharge which we use. We follow the ideas of Winter's concept of hydrological
103 landscapes (Winter 2001) and define Recharge Landscape Units to represent areas for which
104 we expect similar recharge controls. We then compare these areas against an openly
105 available, comprehensive and thoroughly quality assured dataset of ground-based recharge
106 estimates in Africa, recently published by MacDonald et al. (2021). Although we use their
107 database in our analysis, this work has some key differences to the previous work by
108 MacDonald et al. (2021). Firstly, we attempt to explicitly link our analysis to both the
109 qualitative (understanding of recharge controls) and quantitative (i.e., ground-based recharge
110 estimates) findings in the literature, whereas MacDonald et al. (2021) only investigate the
111 quantitative data. Furthermore, our classification approach allows us to explore whether
112 relationships between environmental controls and recharge signatures vary between different
113 environmental settings. In contrast, the statistical approach taken by MacDonald et al. (2021)
114 only allowed them to investigate relationships which applied to the whole continent. Finally,
115 we investigate both long-term mean annual recharge rates of groundwater recharge and
116 recharge ratios. This allows us to understand how different recharge signatures vary and
117 interact in space, furthering our understanding of groundwater recharge beyond just looking
118 at annual rates.

119 **2. Review of process controls on groundwater recharge across Africa**

120 Most of the existing knowledge base on groundwater recharge processes, controls and rates in
121 Africa comes from a relatively small number of case studies investigating recharge at the
122 field, catchment, or sometimes regional scale. These studies use a wide range of methods to
123 understand recharge processes throughout the continent, with approaches often varying
124 according to environmental setting, data availability and the objective of the individual
125 studies (MacDonald et al. 2021). Details of the strengths and weaknesses of the different

126 methods can be found in Scanlon et al. (2002) and Healy (2010). We organize the review of
127 controls into four domains: climate and weather, topography, landcover/use, and soils and
128 geology. The aim of this review is firstly to identify dominant controls on groundwater
129 recharge, and secondly to understand whether these controls have clear positive or negative
130 relationships with groundwater recharge, or if their relationship with recharge is ambiguous.
131 We are considering processes that govern the potential recharge of an aquifer, which can be
132 more than the actual recharge due to interflow processes or if the potential recharge rate is so
133 large that it exceeds the rate at which water can flow laterally through the aquifer (Theis
134 1940). In the latter case, the aquifer can become over-full such that available recharge is
135 rejected. We show a summary of this review in Figure 1. An extended version of the review
136 can be found in the supplemental information.

137 *Climate and weather*

138 Annual scale components of the water-energy balance are a first order control on the spatial
139 variability of groundwater recharge (Kim and Jackson, 2012; Mohan *et al.*, 2018; Cuthbert *et*
140 *al.*, 2019b; MacDonald *et al.*, 2021), as they control the quantity of water available to be
141 partitioned into groundwater recharge, as well as the energy available to partially control
142 atmospheric losses (Budyko, 1974). Hence studies in Africa show variability of annual
143 recharge rates along a climate gradient, largely defined by precipitation due to the generally
144 high levels of energy available (MacDonald *et al.* 2021). In an upland catchment of
145 Cameroon where rainfall exceeds 3000 mm/year, estimated recharge rates exceed 900
146 mm/year (Kamtchueng *et al.* 2015), in comparison to recharge rates between 160 mm/year
147 and 330 mm/year in the Ethiopian Highlands where annual rainfall is approximately 1300
148 mm/year (Azagegn *et al.* 2015; Banks *et al.* 2021; Demlie 2015). Groundwater resources
149 throughout the deserts, which receive very little annual rainfall (Nicholson 2000), are
150 recharged at rates below 5 mm/year (Foster *et al.*, 1982; Dabous and Osmond, 2001; Zouari

151 et al., 2011), or may not even be actively recharged (Befus et al. 2017). In these regions deep
152 ‘fossil’ groundwaters recharged prior to the Holocene dominate aquifer stores (Sturchio et al.,
153 2004; Guendouz et al., 2006; Abotalib et al., 2016; Jasechko et al., 2017).

154 Groundwater recharge volumes are often biased towards the rainy season as elevated rainfall
155 is required to overcome high rates of evapotranspiration (Bromley *et al.*, 1997; Demlie *et al.*,
156 2007; Walraevens *et al.*, 2009; Mechal et al., 2015), and greater monthly and daily
157 precipitation intensity leads to a more efficient conversion of rainfall to recharge (Jasechko
158 and Taylor 2015; Owor et al. 2009; Taylor and Howard 1996). Groundwater level
159 observations in the Makutapora wellfield, Tanzania, suggest that recharge is dependent upon
160 months with the most extreme (> 95th percentile) rainfall (Taylor, Todd, et al. 2013) often
161 enhanced by the El Nino Southern Oscillation and the Indian Ocean Dipole. However, the
162 multiple climate oscillations known to affect climate patterns in Africa (Brown et al., 2010)
163 can have opposing effects in different parts of the continent (Nicholson and Kim 1997).
164 Nonetheless, wetting and drying cycles are being reflected in observed groundwater
165 hydrographs throughout Africa (Taylor *et al.*, 2013; Cuthbert *et al.*, 2019b; Kolusu *et al.*,
166 2019), showing both seasonally extreme recharge events as well as recharge events which are
167 more episodic in nature.

168 Episodic rainfall events are particularly important in arid landscapes where recharge often
169 depends upon a small number of days of intense rainfall (Vogel and Van Urk, 1975; Mazor *et*
170 *al.*, 1977; Van Tonder and Kirchner, 1990; Nkotagu, 1996; De Vries et al., 2000; Xu and
171 Beekman, 2003; Wanke et al., 2008). Döll and Fiedler (2008) stressed the importance of
172 heavy rainfall events in semi-arid and arid regions as they modelled groundwater recharge
173 globally, applying a rainfall threshold of 10 mm/day to drylands, below which they assumed
174 recharge would not occur. They identified this threshold via an independent analysis of 25

175 chloride profile estimates of annual recharge distributed throughout the world as well as
176 regional model estimates of recharge in Death Valley, California.

177 In summary, annual and seasonal precipitation as well as heavy rainfall events have a positive
178 relationship with groundwater recharge in Africa – largely driving inter- and intra-annual
179 recharge variability, while the amount of energy available from radiation has a negative
180 relationship with groundwater recharge. However, the influence of large-scale climate
181 oscillations on groundwater recharge in Africa is less clear as their effect on climate patterns
182 vary regionally.

183 *Topography*

184 Topographic slope controls the movement of water across the land surface and therefore
185 controls water infiltration into the subsurface and groundwater recharge, with gentler slopes
186 promoting more recharge than steeper slopes (Simmers 1990). The role of slope in
187 controlling groundwater recharge has been discussed throughout many different regions of
188 Africa, including Ethiopia (Gebreyohannes et al. 2013), Nigeria (Abdullateef et al. 2021;
189 Fashae et al. 2014), Botswana (Lentswe and Molwalefhe 2020) and Algeria (Boufekane et
190 al., 2020). Yet interestingly, McKenna and Sala (2018) found that recharge beneath flat
191 playas in the south-western United States is greater when they are surrounded by steeper
192 slopes which promote greater run-on onto the playa.

193 In dry regions, intense rainfall events are important drivers of focused recharge through flash
194 flooding (Sultan et al. 2000) and the formation of ephemeral water bodies and depression
195 storage (Lehner and Döll, 2004) , i.e. in areas where water accumulates on the land surface.

196 In Africa’s dry regions, alluvial aquifers underlying dry riverbeds are recharged episodically
197 or perhaps seasonally by river transmission losses following heavy rainfall (Tantawi, El-
198 Sayed and Awad, 1998; Sultan *et al.*, 2000; Gheith and Sultan, 2002; Benito *et al.*, 2010;

199 Walker et al., 2019; Seddon *et al.*, 2021). These storms can activate focused recharge
200 mechanisms despite negligible diffuse recharge in interfluvial regions due to high evaporation
201 (Favreau et al. 2009). In endoreic arid basins, surface water can also accumulate in salt pans
202 which typically occupy topographic depressions (Lehner and Döll 2004). (De Vries et al.,
203 2000) use chloride profiles to show that in the eastern fringes of the Kalahari Desert, recharge
204 is enhanced under these pans, with estimated annual rates of 50mm in comparison to 7mm for
205 the surrounding landscape.

206 Therefore, slope generally has a negative relationship with groundwater recharge since it will
207 provide an easier flow path for water to move downhill, whereas topographic depressions
208 have a positive relationship with (focused) groundwater recharge because they allow water to
209 accumulate.

210 *Landcover/use*

211 Landcover and use varies considerably across the African continent. Bare soils (33% of
212 Africa's land area) occupy most of northern Africa as well as parts of southern and eastern
213 Africa, whilst grasslands (15.4%), shrublands (13.4%) and agriculture (11.6%) are largely
214 distributed throughout the Sahel and Southern and Eastern Africa, and forests and woodland
215 (26%) spread across western, central and south-eastern regions (Mayaux *et al.*, 2004;
216 Tsendbazar et al., 2017; Xiong *et al.*, 2017). These vegetation patterns influence the spatial
217 variability of groundwater recharge (Kim and Jackson 2012) through their control over
218 transpiration, interception and soil evaporation fluxes (Gordon *et al.*, 2005; Schlesinger and
219 Jasechko, 2014; Good et al., 2015).

220 An estimated 7% of the continent's precipitation returns to the atmosphere via interception
221 evaporation, mostly occurring in the densely forested regions of Central Africa where this
222 flux can exceed 10% of the precipitation input (Miralles et al. 2010; Zhang et al. 2016; Zheng

223 et al. 2017). Globally, we could not find any studies directly discussing the relationship
224 between rainfall interception and groundwater recharge. However, it seems reasonable to
225 assume that by limiting the amount of precipitation reaching the land surface, interception
226 consequently reduces groundwater recharge.

227 An estimated 49% and 21% of precipitation over Africa returns to the atmosphere via
228 transpiration and bare soil evaporation, respectively (Zhang et al. 2016). The bulk of
229 continental transpiration is associated with the tropical forests (Gordon *et al.*, 2005; Good et
230 al., 2015), where tall vegetation with deep rooting systems increases the capacity of root-zone
231 moisture storage (Nijzink et al. 2016) and the access to deeper groundwater (Barbeta and
232 Peñuelas 2017). When investigating groundwater recharge at regional and catchment scales,
233 studies often find that recharge rates are lower in areas which are forested than in areas which
234 are unforested or have bare soils (Gebreyohannes et al. 2013; Houston 1982; Howard and
235 Karundu 1992; Stone and Edmunds 2012). Furthermore, the presence of woodland or forest
236 can restrict groundwater recharge to years of particularly high rainfall, even when recharge in
237 grass, crop or unvegetated parts of the catchment occurs annually (Houston 1982; Howard
238 and Karundu 1992). In the Kalahari Desert, dense bush and tree savannah is believed to
239 transpire much of the annual rainfall during the long dry season, leading to very little
240 recharge (De Vries et al., 2000; Sibanda et al., 2009). Similarly, chloride profiles in Senegal,
241 suggest that groundwater recharge rates decline as vegetation density increases (Edmunds
242 and Gaye 1994). Land clearing, often for agricultural expansion, can also enhance
243 groundwater recharge rates by reducing evapotranspiration (Taylor and Howard 1996; Været
244 et al. 2009).

245 Land clearing for agriculture does not only affect recharge through changes to
246 evapotranspiration, it can also alter the mechanisms through which recharge occurs, by
247 altering soil surface properties (Wirmvem et al. 2015) as well as runoff run-on processes

248 (Leduc et al., 2001; Leblanc *et al.*, 2008; Favreau *et al.*, 2009; Ibrahim *et al.*, 2014; Wirmvem
249 *et al.*, 2015). Agricultural land adjacent to many of Africa's largest lakes and rivers is
250 regularly equipped for irrigation (Siebert et al. 2015). Excess irrigation water can infiltrate
251 into the soil and percolate to the aquifer, increasing groundwater recharge rates (Bouimouass
252 et al. 2020; Scanlon et al. 2007). Nonetheless, as irrigation technologies become more
253 efficient, recharge via irrigation excesses is expected to decline (Scanlon et al. 2007).

254 Urban settings only account for less than 0.01% of the African landscape (Zhou et al. 2015).
255 Although, urbanisation is typically perceived as reducing groundwater recharge by reducing
256 the permeable surface area, recharge rates in urban areas can be as high as or even higher
257 than nearby rural areas (Lerner 2002; Sharp 2010). Urbanization can dampen existing
258 recharge mechanisms, but it can also introduce new mechanisms such as localised recharge
259 where there is little drainage infrastructure (Lerner 2002; Sharp 2010), as well as leakages
260 from on-site sanitation (Foster et al., 1999; Diouf, 2012; Lapworth *et al.*, 2017) and piped
261 distribution networks if such water supply is available.

262 In short, we find that the transpiration and canopy storage controls of different landcovers
263 show a negative relationship with groundwater recharge, whereas the additional supply of
264 water to agricultural land through irrigation has a positive relationship with recharge. Effects
265 of urbanisation on groundwater recharge on the other hand are more ambiguous.

266 *Soils and Geology*

267 Soils with larger sand fractions are more permeable and support higher recharge rates than
268 finer clay soils. In a global scale meta-analysis of recharge estimates, Kim and Jackson
269 (2012) show that on average sandy soils are 50% more efficient in converting water input
270 into groundwater recharge. Similar results are found at regional and catchment scales in
271 Senegal, Sudan and Zimbabwe, whereby higher recharge rates are estimated in areas where

272 the sand fraction is a more dominant component of the soil (Abdalla 2009; Butterworth et al.
273 1999; Edmunds and Gaye 1994). Lower recharge rates are found in clayey soils as the
274 vertical percolation of water through the soil profile is restricted (Attandoh et al. 2013;
275 Edmunds et al. 1992) and soil moisture is more exposed to evapotranspiration (Mensah et al,
276 2014; Yidana and Koffie, 2014; Kotchoni et al., 2018).

277 However, soil texture alone fails to recognise structural soil properties which enable
278 infiltration via preferential flow paths which bypass the soil matrix (Beven and Germann
279 1982). Macropores in the soil structure allow infiltration to bypass vegetation rooting zones
280 and impermeable soil layers (De Vries et al., 2000; Mazor, 1982; Van Tonder & Kirchner,
281 1990; Xu & Beekman, 2003) and facilitate recharge in conditions which would otherwise be
282 prohibitive. These preferential flow paths are an important mechanism for groundwater
283 recharge across a range of contrasting environmental settings. In the Botswanan Kalahari
284 Desert, semi-arid Tanzania and the tropical highlands of Ethiopia, the contribution of
285 preferential flows to groundwater recharge is approximately 24%, 60% and 36%, respectively
286 (Demlie et al. 2007; Nkotagu 1996; de Vries and Gieske 1990).

287 Rock fracturing (Nkotagu, 1996; Xu and Beekman, 2003; Adams et al., 2004; Kebede *et al.*,
288 2005; Kamtchueng *et al.*, 2015) and vertical conduits in karstic rock (Farid *et al.*, 2014;
289 Hartmann *et al.*, 2014, 2017; Chemseddine et al., 2015; Ayadi *et al.*, 2018; Leketa *et al.*,
290 2019) also provide preferential flow paths for groundwater recharge. In dry landscapes such
291 as the Kalahari Desert, rock fracturing at bedrock outcrops and isolated rock formations
292 called inselbergs (Burke 2003) can locally enhance groundwater rates (Mazor, 1982;
293 Butterworth *et al.*, 1999; Brunner *et al.*, 2004; Wanke et al., 2008). The distribution and
294 geometry of the superficial geology can also have a marked impact on recharge pathways and
295 rates in conjunction with the underlying bedrock and distribution of stream networks (Zarate

296 et al. 2021). Similar observations have been made regarding focused recharge opportunities
297 for water in karstic regions (Hartmann et al. 2017).

298 Soil perturbations such as crusting, cementation, compaction, weathering, and tillage can also
299 have a significant impact on recharge rates. Whilst studies mostly find that soil crusting
300 (Favreau et al. 2009; Jacks and Traoré 2014; Wakindiki and Ben-Hur 2002), cementation
301 (Nash et al., 1994; De Vries et al., 2000; Xu and Beekman, 2003; Francis *et al.*, 2007) and
302 compaction (Hamza and Anderson, 2005; du Toit et al., 2009) reduce the permeability of soil
303 layers and hence reduce groundwater recharge, the effects of deeply weathered soils known
304 as laterites (Bromley *et al.*, 1997; Ruedi *et al.*, 2005; Cuthbert and Tindimugaya, 2010;
305 Bonsor et al., 2014) and agricultural tilling practices (Abu-Hamdeh, 2004; Osunbitan et al.,
306 2005; Spaan et al., 2005; Strudley et al., 2008; Thierfelder and Wall, 2009; Abidela Hussein
307 *et al.*, 2019) on recharge are much less clear.

308 Therefore, in summation we find that, soil grain sizes, bedrock outcrops and properties which
309 promote preferential flow paths, such as soil macropores, rock fractures and karst geology,
310 have a positive relationship with groundwater recharge. Some soil perturbations such as
311 compaction, cementation and crusting have a negative relationship with groundwater
312 recharge, whereas others, including tilling and soil laterization, have a less clear relationship
313 with recharge.

314 *Interactions between controls*

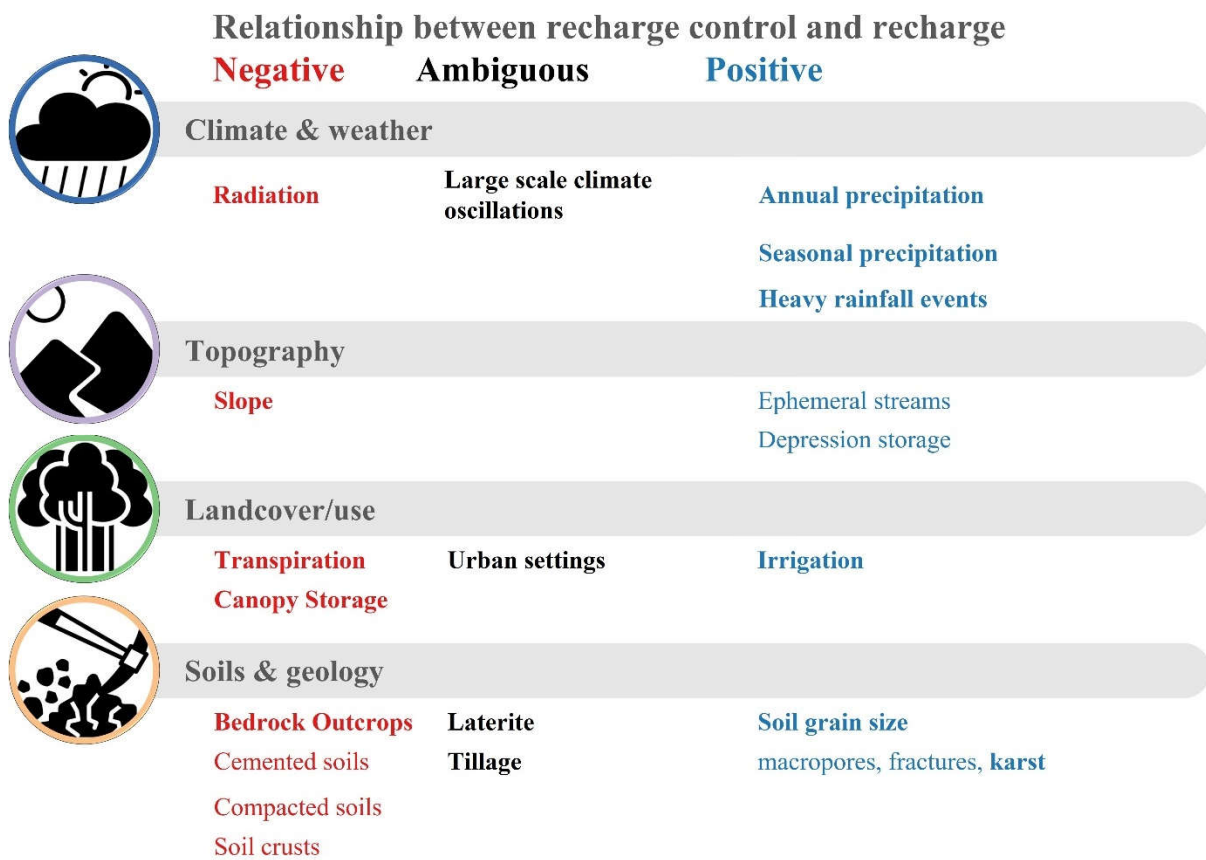
315 Up to now we have largely looked at landscape properties and their control over recharge
316 processes independently, in reality, groundwater recharge is a function of the interactions
317 between these controls. Hence at the continental scale, we would typically expect to find
318 some of the lowest recharge rates in areas with the most freely draining soils, as these regions
319 also have the lowest precipitation volumes. By identifying patterns in the landscape, i.e.

320 climate, topography, vegetation, soils and geology, we can begin to conceptualise recharge
321 processes of different environmental settings found in Africa. We can find these patterns as
322 landscapes are continuously co-evolving (Troch et al. 2013) via an array of physical and
323 biological processes which effect the uplift and deformation of bedrock and the erosion,
324 transportation and deposition of sediments (Dietrich and Perron 2006; Reinhardt et al. 2010).
325 This co-evolution, explains why we typically expect to find certain landscapes throughout the
326 continent, including rainforests, tropical woodlands and savannas and deserts.

327 We often regard climate as an external force driving the hydrological system, but it also
328 controls the spatial and temporal patterns of landcover (Zhou et al., 2014; Hawinkel et al.,
329 2016; Bouvet et al., 2018; Measho et al., 2019; Ndehedehe et al., 2019) and soils (Jenny
330 1941; Towett et al. 2015). Climate and vegetation patterns as well as soil properties are also
331 strongly affected by local topography. In mountainous areas we see vegetation becoming
332 shorter and less dense above the treeline, as temperatures decline and thinning soils make
333 ground conditions less stable (Harsch et al., 2009; Egli and Poulénard, 2016). Increased
334 precipitation and runoff due to orographic forcing as well as steeper slopes, promote more
335 active erosion and sediment transport fluxes at elevation and therefore prevents the
336 accumulation of soils (Acosta et al. 2015). In contrast, at lower elevations, vegetation can
337 assist the accumulation of soils by reducing surface water erosion and promoting infiltration
338 (Acosta et al. 2015; Descheemaeker et al. 2006; Descroix et al. 2009; Thompson et al. 2010).
339 In water limited regions, vegetation density often increases in topographic depressions such
340 as ephemeral streams, as accessibility to groundwater may be locally improved (Morin et al.,
341 2009; Steward et al., 2012; Ndehedehe et al., 2019; Grodek et al., 2020).

342

343



347 Figure 1. Summary of groundwater recharge controls for Africa identified in the literature. Controls are colour coded
 348 according to their relationship with recharge with red and blue representing negative and positive relationships, respectively.
 349 Bold font highlights controls which we can characterise using global datasets.

351 3. Materials and methods

352 To consolidate our understanding of groundwater recharge controls taken from reviewing the
 353 literature, we take a classification approach which we can then use as a tool to investigate
 354 why ground-based estimates of groundwater recharge vary spatially across the continent. We
 355 acknowledge their will be uncertainty in the classification due to the limitations of our own
 356 understanding and of global datasets. However, we aim to connect the qualitative and
 357 quantitative information obtained from local/regional findings to large scale regionalization
 358 approaches.

359 3.1 Global Datasets

360 We used nine global datasets to characterize the previously identified groundwater recharge
361 controls. Furthermore, controls were only integrated into our classification if the literature
362 indicated it had a clear positive or negative relationship with groundwater recharge and it
363 could be characterized using global datasets. The datasets used and the indices calculated are
364 summarized in Table 1.

365 Indices describing annual and seasonal climate attributes mostly characterise first-order
366 estimates of the water potentially available for groundwater recharge (P-PET) annually and
367 seasonally as well as its variability. This also builds on previous work by Wolock et al.
368 (2004) who used P-PET as the climatic index to delineate hydrological landscapes in the
369 United States. We characterised heavy rainfall across Africa using a threshold of 10 mm/day.
370 Several studies in Africa (Döll and Fiedler 2008; Owor et al. 2009; Taylor and Howard 1996)
371 have found annual recharge has a stronger correlation with the average volume of rainfall per
372 year on days with at least 10 mm of rain, than with mean annual precipitation and hence we
373 selected this as threshold for heavy rainfall in Africa. Though we acknowledge the rainfall
374 threshold for recharge occurrence likely varies across the continent. We characterized the
375 influence of landcover on groundwater recharge via transpiration and canopy storage
376 processes, by attributing vegetation specific transpiration coefficients to a landcover dataset
377 and by looking at the Leaf Area Index, respectively. This approach is also often taken when
378 parameterizing these processes in continental scale hydrological modelling (Telteu *et al.*,
379 2021). To avoid having multiple indices to describe soil textures we instead calculated the
380 ratio of soils which promote infiltration (i.e., sand) to those which restrict infiltration (i.e., silt
381 and clay) (Saxton *et al.*, 1986; Wösten et al., 2001). We used the depth to bedrock dataset of
382 (Pelletier et al. 2016) to highlight bedrock outcrop regions and the world map of carbonate
383 rock outcrops (Williams and Ford 2006) to highlight the extent of carbonate rock outcrops.

Table 1. Details of the recharge control indices we defined to characterise recharge controls across Africa and the global datasets we used to calculate them.

Attribute	Description	Units	Period	Data source	Reference
Climate attributes					
P-PET	Mean annual precipitation minus mean annual PET.	mm/year	1979-2015	1. MSWEP v1.2 (Precipitation)	1. (Beck et al. 2017)
P-PET in season	Mean annual volume of precipitation in excess to PET in months considered in-season. A month is considered in-season when P exceeds PET.	mm/year	1979-2015	Spatial res.: 0.25° Temporal res.: Daily 2. CRU v4 (PET)	2. (Harris et al., 2020)
σ (P-PET)	The standard deviation of monthly P-PET	mm/month	1979-2015	Spatial res.: 0.5° Temporal res.: Monthly	
P10	The average volume of rainfall per year on days with at least 10 mm of rain.	mm/year	1979-2015		
Topography attributes					
Slope	Geodesic slope of the DEM using a 3 by 3 moving window.	Degrees	-(Lehner, Verdin, and Jarvis 2013)	HydroSHEDS Spatial res.: 15 arc seconds	(Lehner et al., 2013)
Landcover/use					
Kveg	Vegetation coefficient related to transpiration. Vegetation-specific annual values (L. J. Gordon et al. 2005) applied to a landcover classification. Mean value from 1992-2005.	-	1992-2015	ESA-CCI v2.0.7 Spatial res.: 300m Temporal res.: Yearly	(Defourny et al. 2017)
LAI	Mean leaf area index (based on 12 monthly means from 1981-2015)	-	1981-2015	GIMMS-LAI3g v2 Spatial res.: 0.25° Temporal res.: Monthly	(Mao and Yan. 2019)
Irrigation	Area equipped for irrigation multiplied by the fractional area actually irrigated.	km ²	2005	Global Map of Irrigation Areas Spatial res.: 5 arc minutes	(Siebert et al., 2013)
Soil attributes					
Sand / (Clay + Silt)	The ratio of sand (>0.05mm) to silt (0.002-0.05mm) and clay (<0.002mm) in the fine earth fraction of the top 2m of the	-	-	SoilGrids250m Spatial res.: 250m	(Hengl et al. 2017)

soil profile.
Proportions of each soil texture are by weight. Take the depth weighted harmonic mean across intervals of 0-5cm, 5-15cm, 15-30cm, 30cm-60cm, 60-100cm, 100-200cm.

Geology attributes					
Depth to bedrock	Average soil and sedimentary deposit thickness. Maximum of 50m.	m	-	Gridded Thickness of Soil, Regolith and Sedimentary Deposit Layers Spatial res.: 30 arc seconds	(Pelletier et al. 2016)
Karst	Extent of carbonate rock outcrop areas.	-	-	World Map of Carbonate Rock Outcrops V3.0	(Williams and Ford 2006)

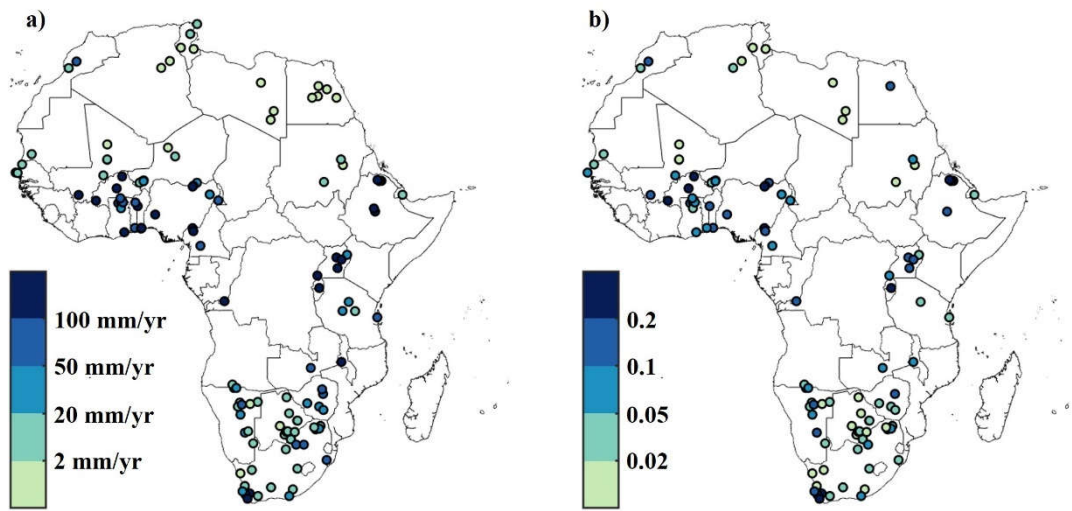
387

388 3.2 Ground-based annual recharge and recharge ratio estimates

389 We used the database compiled by MacDonald et al. (2021) of long-term mean annual
 390 recharge estimates compiled from case studies in the literature. We selected this database
 391 above other meta-datasets (Moeck et al. 2020; Mohan et al. 2018) because of its focus on
 392 Africa, the thorough quality assurance conducted throughout its compilation, and the
 393 additional meta-data provided, such as recharge estimate uncertainty ranges. Through quality
 394 assurance steps, MacDonald et al. (2021) removed 182 datapoints (from an initial 316), due
 395 to duplicative studies in the same location and findings which were solely dependent upon
 396 hydrological modelling. Additional screening removed data points where the site co-ordinates
 397 and date of the study period were not provided. Finally, we removed estimates dated prior to
 398 1979 or after 2015, as they would not correspond to the timing of the climate datasets we
 399 used. Ultimately, we were left with 129 ground-based estimates of annual groundwater
 400 recharge distributed across Africa. 111 of these sites/studies also reported corresponding
 401 mean annual precipitation rates, so we could estimate long-term mean recharge ratios at these

402 locations (Figure 2). Spatially, 31 of these estimates reflect recharge rates over spatial scales
403 less than 100 km², a further 41, 29, and 28 are for spatial scales of 100-2500 km², 2500-
404 62500 km² and greater than 62500 km², respectively.

405



406

407 Figure 2. The remaining annual recharge and recharge ratio estimates collected from case studies in the literature by
408 MacDonald et al. (2021), after initial screening of the data. The recharge ratio is defined as the fraction of precipitation being
409 converted to recharge (recharge / precipitation).

410 3.3 Clustering

411 To delineate regions with expected similar recharge control indices (i.e., Recharge Landscape
412 Units) we use a fuzzy c-means clustering algorithm (Bezdec 1981). This fuzzy clustering
413 algorithm allows for pixels to belong to multiple units simultaneously, albeit with varying
414 degrees of membership, thus enabling us to study the gradual transition between units (e.g.,
415 reflecting different landscapes). The degree of overlap in membership allowed us to
416 determine the uniqueness of each delineated Recharge Landscape Unit. The degree of
417 membership is dependent upon how close in value each pixel's recharge control indices are to
418 the centroid of each unit, which is regarded as being representative for a unit. Membership
419 scores vary from 0 to 1, with 0 representing no similarity and 1 suggesting the pixel's

420 recharge control indices are equal to the values of the unit's centroid. Further details on the
421 algorithm and on application details are provided in the supplemental material. Ultimately,
422 we attributed each pixel to the unit with which it has the highest degree of membership,
423 which we refer to as its primary unit.

424 **3.4 Random Forests**

425 We used classification-based Random Forests to expand our classification for recharge
426 controls in Africa to the rest of the world. Random Forests is a machine learning algorithm
427 which combines multiple trees to produce an ensemble of predictions (Breiman 2001;
428 Breiman et al. 1984), which link predictor variables (recharge control indices) to a response
429 (Recharge Landscape Units). Each individual tree develops rules for predicting responses
430 which are structured as a binary decision tree composing of nodes and branches. At each
431 node a conditional binary split is applied to one of the predictor variables. The split forms
432 two branches which link to nodes in the overlying stratum. This splitting continues until the
433 terminal node (the leaf) is met and the outcome is predicted. Each classification tree in the
434 ensemble model is trained on observations (Pixels of classification for recharge controls in
435 Africa) which were randomly selected with replacement from a sub-sample of 70% of the
436 total observations ('in-bag' observations). The random forest model consists of 25 trees each
437 with a maximum of 400 decision splits. Increasing the number of trees or decision splits did
438 not significantly improve model performance. Addor et al., (2018) previously used Random
439 Forests to predict observed streamflow signatures across the USA and Stein et al., (2021)
440 used random forests to explore how climate and catchment attributes influence flood
441 generating processes.

442

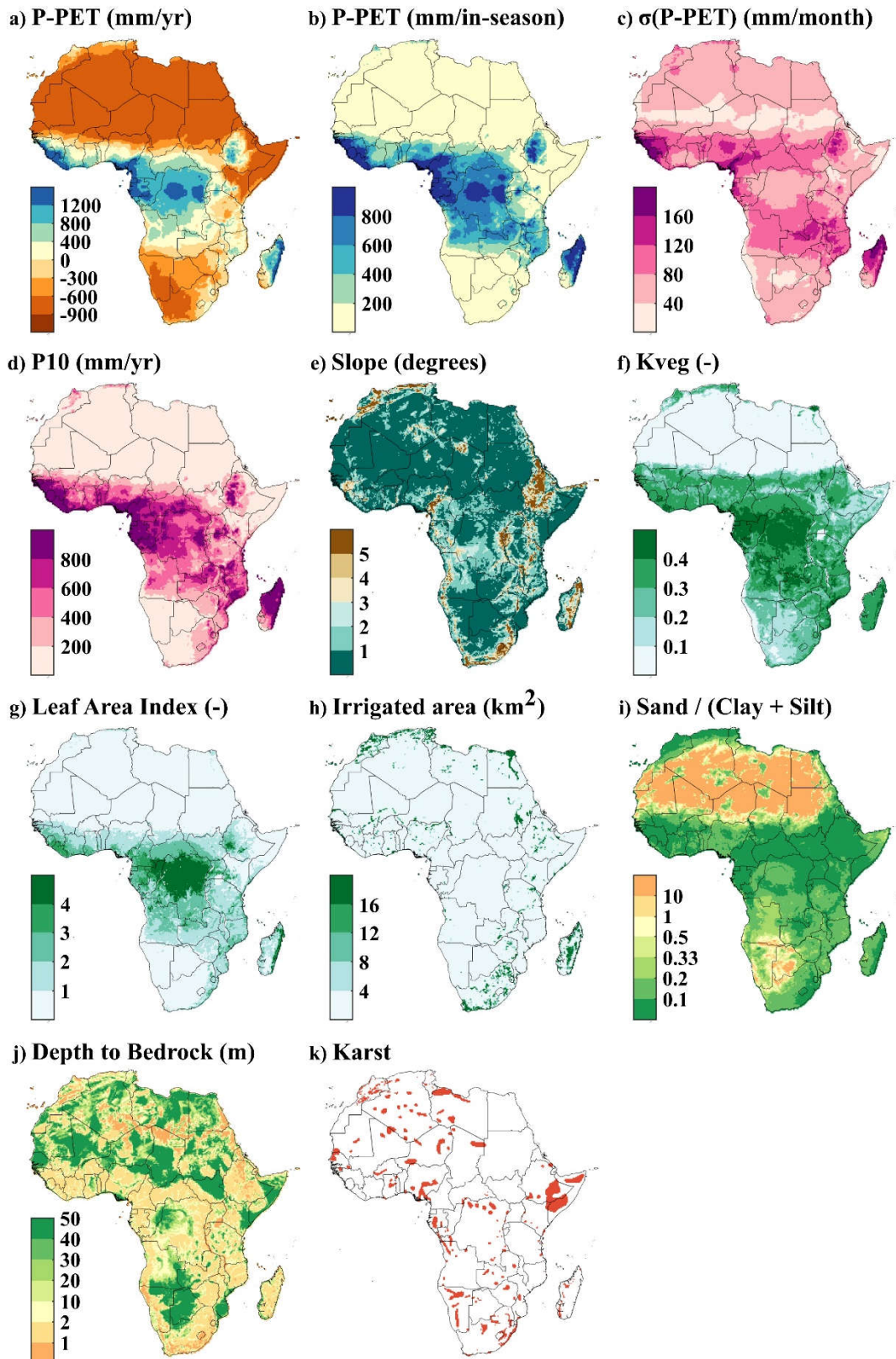
443

444 **4 Results**

445 **4.1 Recharge Landscape Units outline regions with similar recharge**

446 **controls in Africa**

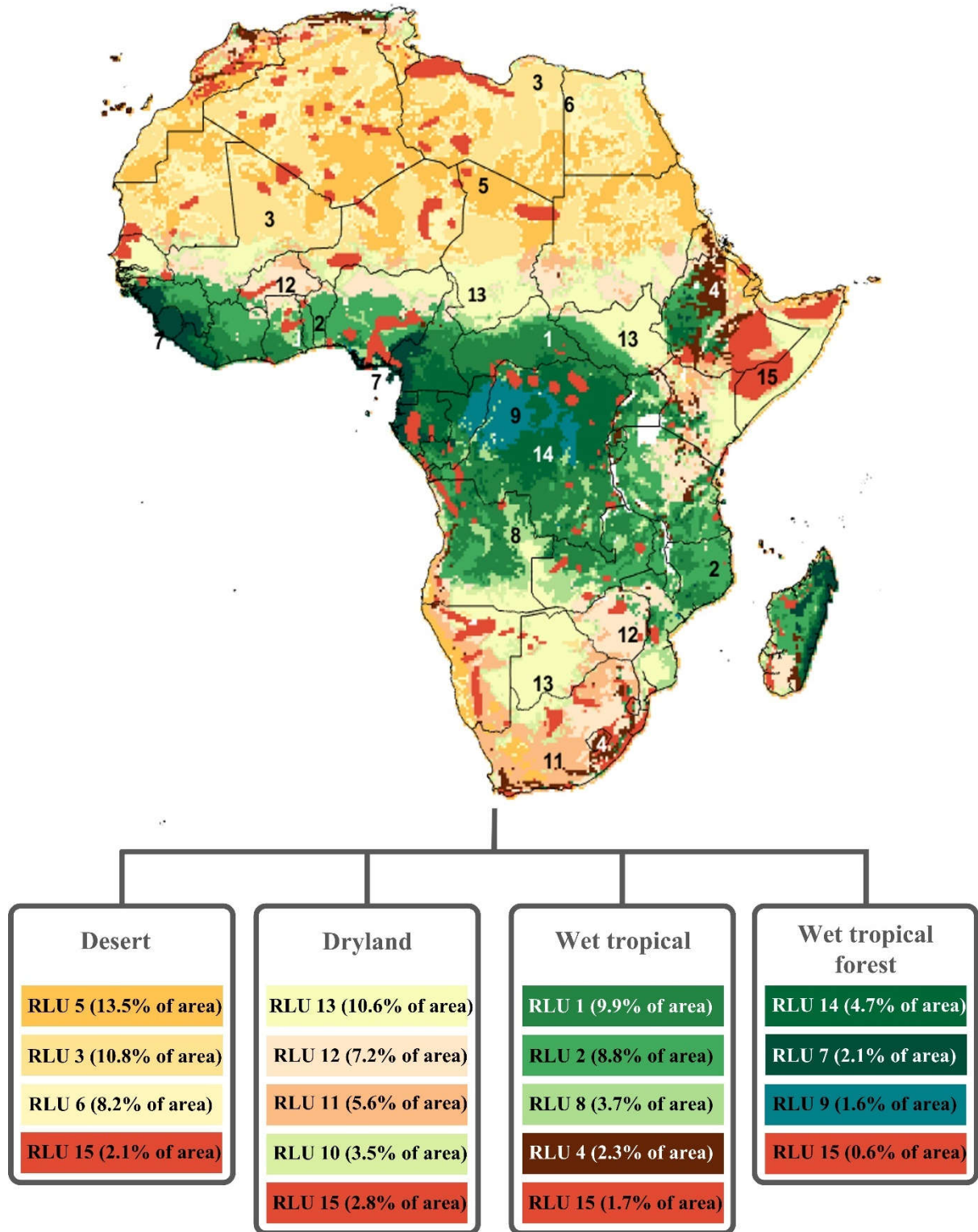
447 Based on our review in section 2, we defined and calculated 11 indices to characterise the
448 different controls on distributed groundwater recharge we identified in our review (Figure 1).
449 To avoid using redundant information for each control, we checked the correlations between
450 each of the indices initially considered and removed indices such that none of the indices for
451 a given control had Pearson correlation coefficients greater than or equal to 0.7 with one
452 another (see supplemental information) (Dormann et al. 2013).



453

454 Figure 3. 11 recharge control indices characterising controls identified in the literature using global datasets. a) P-PET; b) P-
 455 PET in-season; c) σ (P-PET); d) P10; e) Slope; f) Kveg; e) Leaf Area Index; h) Irrigated area; i) Sand / (Clay + Silt); j) Depth
 456 to bedrock; k) Karst. The definitions of each index the datasets used for their characterisation are stated in Table 1.

457 The cluster analysis combines the 11 indices into 15 Recharge Landscape Units with similar
458 recharge control indices of which 9 cover over 80% of the African land area (Figure 4). We
459 initially identified 14 units using fuzzy clustering, as additional units did not greatly reduce
460 the dissimilarity within individual units. The 15th unit which delineates potential karst regions
461 was manually superimposed. Even though we expect recharge to vary significantly between
462 the different settings in which karst is found, we delineate the group as a whole, because we
463 expect the recharge mechanism associated to karst environments to be a dominant control on
464 recharge processes. We can see the continent has been roughly organised into very dry
465 regions in the north and south of the continent and wetter regions spanning from West Africa
466 down through Central Africa towards Mozambique and Madagascar. Even though the spatial
467 organisation of the units suggest proximity is a reasonable indicator for similarity, we do find
468 regions with similar recharge control indices which are also far away from each other. For
469 example, hyper arid regions with shallow soils can be found along Namibia's coastline as
470 well as the coastlines of Egypt and Sudan and throughout the Sahara Desert (unit 5) and
471 extremely wet regions can be found on the coast of West Africa and eastern Madagascar (unit
472 7). Likewise dry highland regions with high slope can be found in South Africa, the East
473 African Rift, Ethiopian Highlands and in the Atlas Mountains (unit 4) and flat regions with
474 thick soil profiles can be found throughout the Sahel, South Sudan and the Kalahari basin
475 (unit 13). In contrast, we also find Recharge Landscape Units which appear to represent
476 unique and spatially concentrated areas, such as the Congo Basin Rainforest (units 9 and 14),
477 as well as regions where properties appear more diverse with multiple units appearing within
478 smaller areas, such as Madagascar and Ethiopia.



479

480

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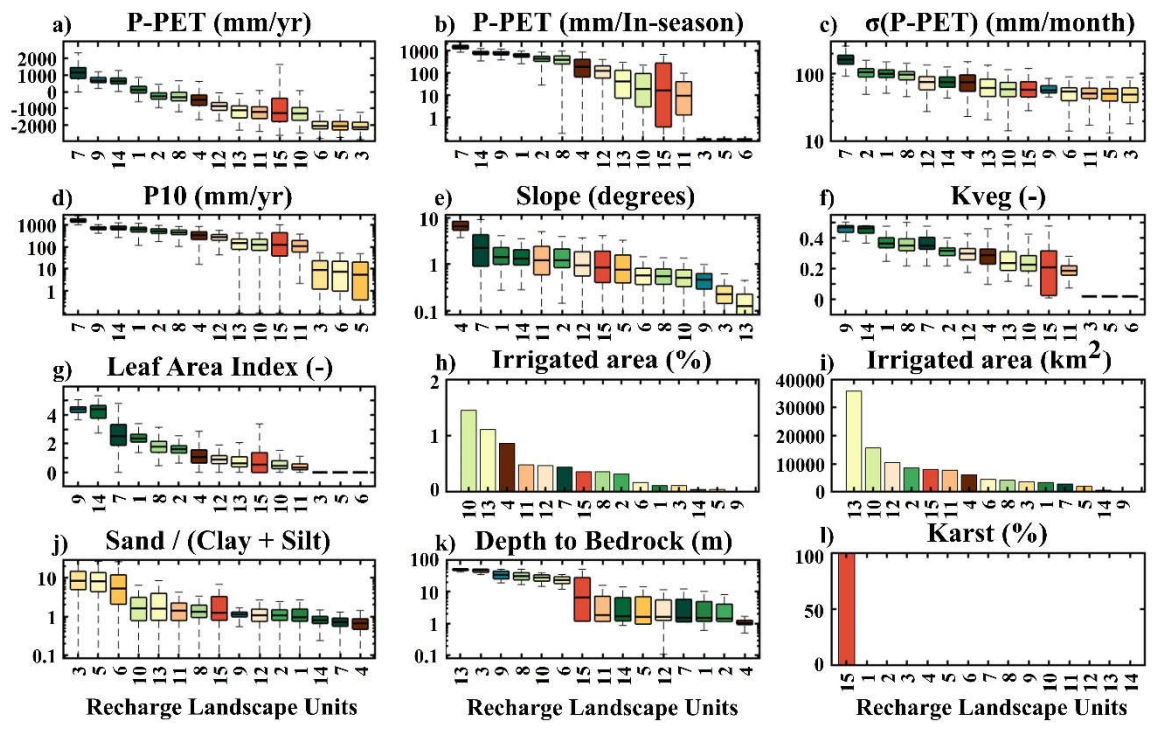
482

Figure 4. Map of the 15 Recharge Landscape Units of our classification for a priori understanding of recharge controls in Africa. We group the Recharge Landscape Units into broader groups of similar units which we call Recharge Landscapes.

483 We found that grouping Recharge Landscape Units into broader Recharge Landscapes
484 suitably organises the African landscape into regions with noticeably different distributions
485 of long-term average annual recharge and recharge ratio (Figure 6). Further disaggregation of
486 the landscape did not allow us to explain any further variability in the ground-based estimates
487 of recharge signatures when using global datasets. These broader Recharge Landscapes also
488 aggregate Recharge Landscape Units with similar recharge control indices, as shown by the
489 boxplots in Figure 5. For each index, boxplots are organized by the median values of each
490 unit, ordered from left to right in descending order. In Dryland and Wet tropical Recharge
491 Landscapes, we see that climate and weather, landcover and soil texture indices transition
492 smoothly across all units. Units within Wet tropical forest Recharge Landscape are typically
493 associated to high Kveg and Leaf Area index values and fine soil textures, whilst units of the
494 Desert Recharge Landscape have low Kveg and Leaf Area values as well as predominantly
495 sandy soils. Similarly, most units have similar topographic slopes except for unit 3, 4 and 13
496 which represent highland and flat plain regions. There is a clear divide in the depth of soils in
497 each of the units, with six of the units showing deeper soil profiles and 8 showing a tendency
498 towards shallow soils. We can see that unit 15 which represents karst regions occurs in a
499 wide range of different climate, topographical, landcover and soil settings. Irrigated areas do
500 not contribute to large areas of any of our Recharge Landscape Units.

501 Desert Recharge Landscapes could only be further differentiated by their depth to bedrock,
502 while other landscape types were dis-aggregated by climate seasonality, slope, landcover and
503 slope, as well as the depth to bedrock. Desert Recharge Landscape Units are differentiated
504 according to where depth to bedrock is less than 13.5 m (unit 5), where the bedrock depth is
505 between 13.5 m and 33.9 m (unit 6) and where the depth to bedrock is greater than 33.9 m
506 (unit 3). This reflects differences in topography throughout Desert Recharge Landscapes, as
507 mountainous Desert Recharge Landscapes with greater slopes also have smaller bedrock

508 depths. Dryland Recharge Landscapes are also largely dis-aggregated according to the depth
 509 to bedrock, with unit 13 representing where bedrock depth is greater than 37m, unit 10 where
 510 bedrock depth is between 16.3m and 37m and units 11 and 12 where the bedrock depth is less
 511 than 16m.

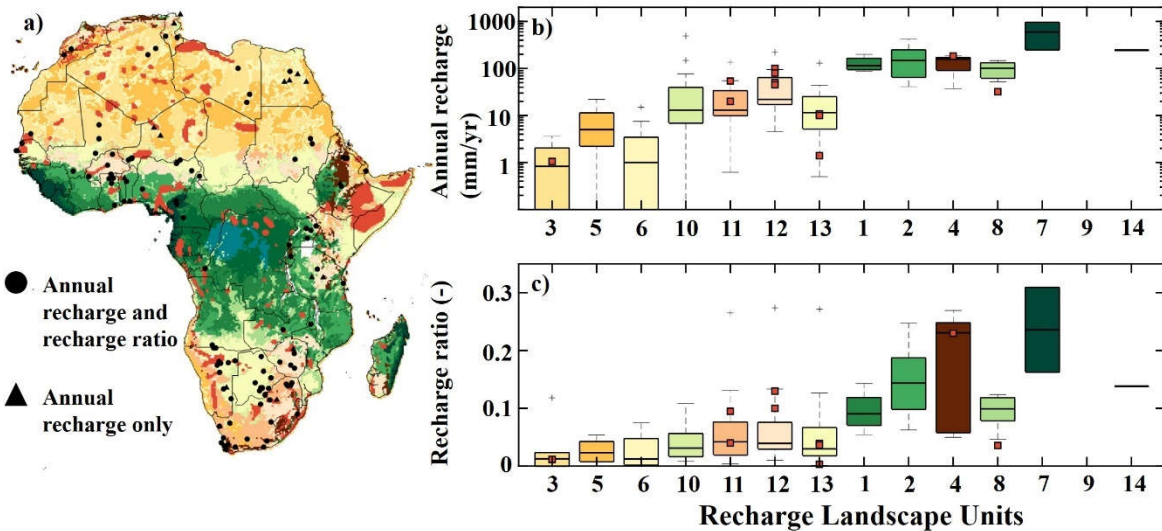


512
 513 Figure 5. Boxplots showing the index values in each of the Recharge Landscape Units we identified. Boxplots are organised
 514 from left to right in descending order of the median values in each unit. We show irrigated area as both the total area
 515 irrigated within the Recharge Landscape Units (h) and as a percentage of the areas for each Recharge Landscape Unit (i).

516 Ground-based estimates of annual recharge (recharge ratio) are bias towards drier settings
 517 with 20 (15), 66 (58), 28 (25) and 3 (3) data points in Desert, Dryland, Wet tropical and Wet
 518 tropical forest Recharge Landscapes, respectively. Recharge Landscapes which have high
 519 annual recharge rates also have higher recharge ratios suggesting that as well as being
 520 generally wetter, they are more efficient in converting that rainfall into recharge (Figure 6).
 521 The variability of ground-based annual recharge estimates within Landscapes is greatest in
 522 wetter settings, as shown by standard deviations of 5.6 mm/year, 66.0 mm/year, 84.0
 523 mm/year and 400.1 mm/year for Desert, Dryland, Wet tropical and Wet tropical forest
 524 Recharge Landscapes, respectively. The standard deviation between the mean annual

525 recharge estimates of each Recharge Landscape is 217.2 mm/year in contrast to a standard
526 deviation of 113.7 mm/year when looking across the whole dataset population. Similarly for
527 ground-based estimates of recharge ratio, the standard deviations within Desert, Dryland, Wet
528 tropical and Wet tropical forest Recharge Landscapes are, 0.033, 0.059, 0.070 and 0.092,
529 respectively. Again, standard deviation between the mean recharge ratio estimates of each
530 Recharge Landscape is greater than across the whole population, each being 0.079 and 0.072
531 respectively.

532 We also investigated the possible influence of the different groundwater recharge estimation
533 methods to see whether this explained any of the variability in annual recharge and recharge
534 ratio estimates within the individual spatial units (see supplemental information). However,
535 in agreement with (MacDonald et al. 2021) we did not find a relationship between the
536 estimation methods used and the recharge signatures. Additionally, one of the benefits of the
537 database compiled by MacDonald et al. (2021) is that they provide uncertainty ranges for
538 each of their ground-based estimates. Although [figure 6](#) does not show these uncertainty
539 ranges, we found that these uncertainty ranges were **lowest** for both annual recharge and
540 recharge ratio in Desert Recharge Landscapes and were largest in Wet tropical and Wet
541 tropical forest Recharge Landscapes (see supplemental). However, uncertainty ranges relative
542 to the ground-based estimates were largest in Desert Recharge Landscapes and lowest in Wet
543 tropical forest Recharge Landscapes. Below we discuss the larger Recharge Landscapes.



544

545 Figure 6. a) Map of ground-based estimate data points distributed across the Classification of recharge controls in Africa.
 546 Boxplots of the ground-based estimates of long-term mean annual recharge (b) and recharge ratio (c) found in each of the
 547 Recharge Landscape Units. No data points are located within Unit 9 and hence it is not shown. Only one data point is located
 548 within Unit 14. Unit 15 representing karst does not have its own boxplot. Instead, we have superimposed (red dots) these
 549 data points above the units which they would have otherwise been attributed to. For this plot we simply use the average
 550 ground-based estimates of annual recharge and recharge ratio at each data point and ignore the uncertainty ranges.

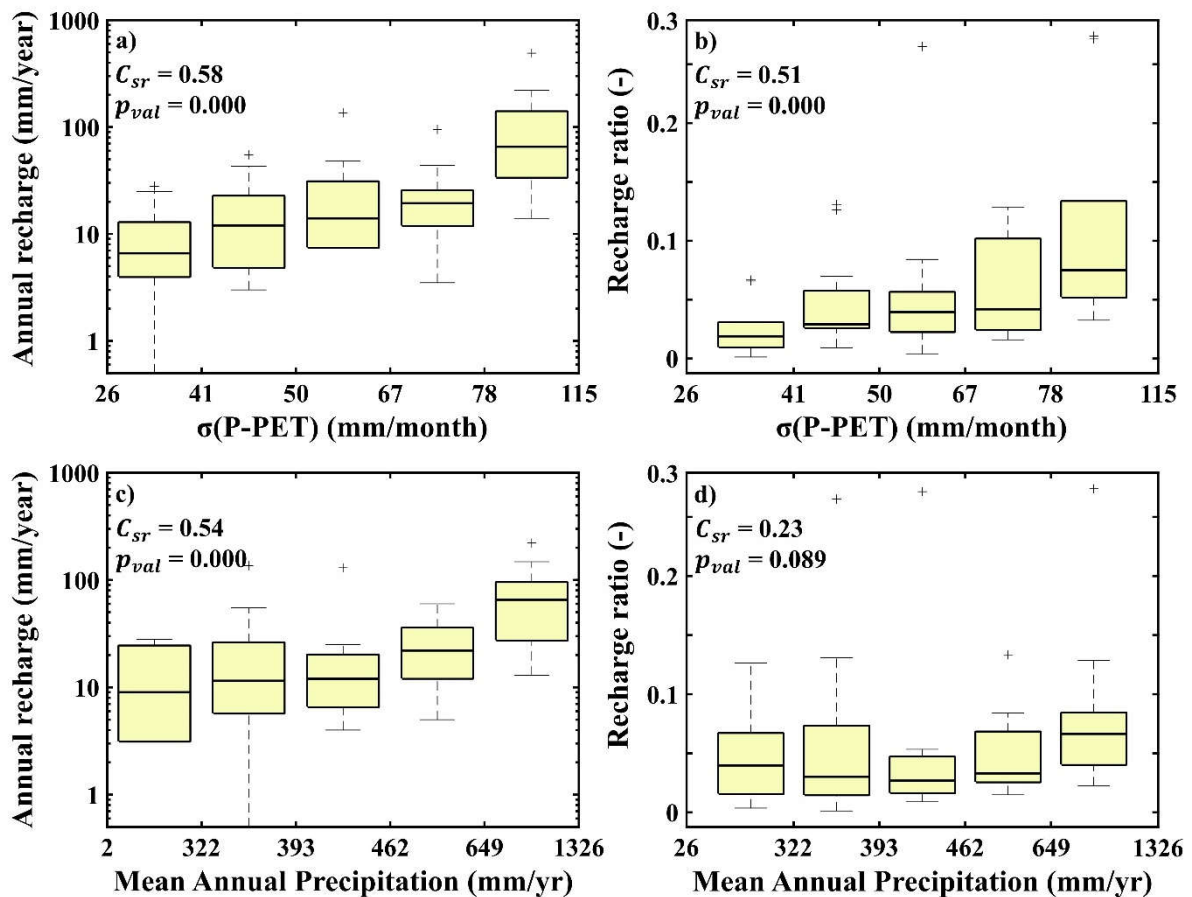
551 *Desert (RLU 3, 5, 6)*

552 Desert Recharge Landscapes are characterised by low moisture availability (P-PET), low
 553 vegetation cover (kveg) and very high sand content in its soils (Figure 5). These properties
 554 lead to the lowest annual recharge and recharge ratio estimates occurring in Africa, as 80% of
 555 annual recharge (recharge ratio) estimates in Desert Recharge Landscapes are below
 556 5mm/year (4%). Low recharge ratios suggest that even when rain does fall, only a small
 557 fraction is converted to recharge, despite the sandy soils, owing to the very high potential
 558 evapotranspiration demand. In these regions evapotranspiration can draw on moisture from
 559 substantial depths which prohibits the downward flux of moisture towards to water table
 560 (Lehmann et al. 2019; Scanlon et al. 2006). We also find ground-based recharge estimates in
 561 Desert Recharge Landscapes show very little variability. Although we find marginally greater
 562 annual recharge rates and recharge ratios in unit 5, we cannot explain why, and differences
 563 may not be significant as there are only 20 data points across this region.

564

565 Dryland (10, 11, 12, 13)

566 About 51% of the 129 ground-based estimates are sited in Dryland Recharge Landscapes
567 where water is generally only available for recharge seasonally (units 10, 11, 12 and 13). 70%
568 of these sites have annual recharge rates between 3-30 mm/year and a further 18% of these
569 sites have rates between 30-100 mm/year. Typically, in these regions less than 10% of
570 rainfall is converted to recharge, with only 9 of the 58 sites recording higher recharge ratios.
571 In this Recharge Landscape, we find that long-term estimates of annual recharge vary
572 according to mean annual precipitation, whereas recharge ratios are greater at sites with
573 greater monthly variability in P-PET (Figure 7).



574

575 Figure 7. Boxplots showing how ground-based estimates of mean annual recharge (a, c) and recharge ratio (b, d) vary
576 according to monthly variability of P-PET and Mean Annual Precipitation in Dryland Recharge Landscapes. Recharge
577 signatures are binned according to percentiles (0-20; 20-40; 40-60; 60-80; 80-100) of the controlling variable. In the top left
578 corner of each sub-plot, we show the spearman rank correlation and the p-value for testing the hypothesis of no correlation.

579

580 *Wet tropical (1, 2, 4, 8)*

581 18 (26) out of 28 annual recharge estimates in the Wet tropical landscapes (units 1, 2, 7, 8)
582 exceed 100mm/year (50mm/year). These sites are also the more efficient in converting
583 rainfall to recharge with 56% (92%) of them having recharge ratios greater than 10% (5%).
584 The wetter conditions as well as seasonal periods of heavy monsoon rain allows deeper
585 drainage, despite increased partitioning of rainfall at the land surface by vegetation, steeper
586 terrain, and less permeable soils. Most of the variability between and within Wet tropical
587 landscape units is attributed to differences in annual and seasonal scale water excess (P-PET)
588 and heavy rainfall events (P10).

589 Differences in annual recharge and recharge ratio estimates of units 1 (median annual
590 recharge 115mm/year; median recharge ratio 9%) and 2 (median annual recharge
591 148mm/year; median recharge ratio 14%) could be attributed to greater LAI and Kveg
592 properties in unit 2. However, when comparing the properties of the individual sites we do
593 not find this relationship. Highland areas (unit 4) show a particularly large variability in the
594 fraction of precipitation being converted to recharge. This perhaps reflects the high degree of
595 variability we can expect in highland regions depending upon landscape positioning.

596 *Wet tropical forest (7, 9, 14)*

597 These areas are characterised by the highest vegetation cover (LAI) and moisture availability
598 (P-PET). We only have three ground-based estimates of annual recharge and recharge ratio
599 within this Recharge Landscape: two in unit 7 and one in 14. The highest annual recharge
600 estimate in our database is located in unit 7, with 31% of rainfall being converted to recharge
601 to allow a rate of 941 mm/yr. Referring to existing literature, we find that in addition to high
602 annual precipitation rates (3050 mm/yr) extensive bedrock fracturing near the land surface
603 enables rapid infiltration and recharge (Kamtchueng et al. 2015).

604 *Karst – present across the other Landscapes (15)*

605 We do not find a clear pattern whereby the presence of karst at a site indicates higher annual
606 recharge rates or recharge ratios than other sites within a similar setting (Figure 6). When
607 investigating the individual studies, some studies reported karst despite not being identified
608 as such by the global dataset, and vice versa Within settings defined as karst by global
609 datasets, annual recharge rates and recharge ratios increase with increasing annual scale P-
610 PET (see supplemental information).

611 **5 Discussion**

612 **5.1 Which regions of Africa show similar recharge controls when clustered** 613 **using descriptors derived from global datasets?**

614 We find 15 Recharge Landscape Units within which we expect recharge processes to be
615 similar, according to our clustering result. Only 9 Recharge Landscape Units are needed to
616 characterize over 80% of the continent's land area. Although we initially wanted to
617 investigate the potential membership of pixels to multiple Recharge Landscape Units, more
618 than 80% of pixels have a membership greater than 0.8 to their primary unit (please see
619 supplemental information). This is likely due to the high dimensionality of the dataset we are
620 using in the cluster analysis. In light of this, and because we further aggregated our 14 (out of
621 15) Recharge Landscape Units into broader Recharge Landscapes (largely according to
622 climate), we simply used the primary unit membership. The Recharge Landscapes we
623 identify are Desert, Dryland, Wet tropical and Wet tropical forest, which account for 32.5%,
624 26.9%, 24.6% and 8.4% of Africa's land area respectively (total of 92.4%). An additional
625 7.25% of the continent's land area is defined by its geology (i.e. karst) and can be found
626 distributed across each of the four previously mentioned Recharge Landscapes (as we would
627 expect according to previous studies, e.g. Hartmann *et al.*, 2017). At the resolution of our

628 classification, climate indices have strong positive correlations with landcover indices
629 (pearson correlation coefficient > 0.7). It is not surprising that our Recharge Landscapes
630 strongly resemble previous climate classifications (Peel et al., 2007; Knoben et al., 2018),
631 because climate is a dominant control on the long-term evolution of land surface and near
632 surface landscape characteristics including topography (Chen et al. 2019), soils and
633 vegetation (Pelletier et al. 2013). It is important to recognise that the classification of places
634 may vary temporally (Aleman et al. 2020; Tierney et al. 2017), however as existing
635 continental scale datasets for ground-based recharge estimates only provide long-term mean
636 annual rates, we were not currently able to investigate the temporal variability of groundwater
637 recharge and how this relates to changing landscape classification (Sawicz et al.
638 2014). Furthermore, we regard the classification as a tool for analysis rather than something
639 unchanging in time.

640 Our Recharge Landscapes broadly resemble the ecozones in classifications by Olson *et al.*
641 (2001) and Jasechko *et al.* (2014), which identify five and three different regions across
642 Africa respectively. They are also similar to the five regions delineated by MacDonald *et al.*
643 (2021) when using aridity classes to investigate the spatial variability of recharge across
644 Africa. Unlike Olson *et al.* (2001) and Jasechko *et al.* (2014) we do not aggregate deserts and
645 xeric shrublands, which we instead include in our Dryland Recharge Landscapes. Hence our
646 Desert Recharge Landscapes more closely align with the hyper-arid regions delineated by
647 MacDonald *et al.* (2021), whilst our Dryland Recharge Landscapes also align with their arid
648 and semi-arid regions. By separating dry systems according to the occurrence of vegetation,
649 we differentiate between regions where transpiration has a greater effect on recharge
650 processes (Scott *et al.*, 2006; Cavanaugh et al., 2011; Gebreyohannes *et al.*, 2013).
651 Consequently, we organise the Kalahari Desert as a Dryland, as it is affected by transpiration
652 (Foster et al. 1982). Our Dryland Recharge Landscapes can be found throughout the desert,

653 shrubland and tropical biomes of classifications by Olson *et al.* (2001) and Jasechko *et al.*
654 (2014). Thus, previous ecozone classifications may have delineated these regions too broadly.
655 We also see that by identifying Dryland Recharge Landscapes with low slope and high
656 bedrock depths (RLU 13), we identified a landscape unit where large seasonal wetlands are
657 likely to occur (Olson *et al.* 2001). These wetlands include the Okavango delta, the Kafue and
658 Barotse floodplains in Southern Africa; the Sudd Swamps in Eastern Africa; and the inland
659 Niger delta, Hadejia-Nguru wetlands and wetlands of Southern Chad in the Sahel. Such
660 wetlands can be significant sources of annually occurring focused groundwater recharge,
661 given soil conditions do not restrict infiltration (Edmunds *et al.*, 1999; Wolski *et al.*, 2006).
662 Unlike the classifications of Olson *et al.* (2001), Jasechko *et al.* (2014) and MacDonald *et al.*
663 (2021), we further disaggregate Desert Recharge Landscapes according to depth to bedrock.
664 In Desert Recharge Landscapes, shallow bedrock depths largely align with mountainous
665 regions, which are often regarded as important recharge zones for current episodic recharge
666 events (Gheith and Sultan 2002; Sultan *et al.* 2007) and more regular recharge events in
667 previous paleoclimate periods (Sturchio *et al.* 2004). Our Wet tropical forest Recharge
668 Landscapes largely align with the tropical and subtropical moist forests shown in Olson *et al.*
669 (2001). Though further disaggregation into units identifies unique regions such as the Swamp
670 forests of the Congo Basin and regions with extreme monsoonal rainfall in the Gulf of
671 Guinea. In contrast, neither Jasechko *et al.* (2014) nor MacDonald *et al.* (2021) identify the
672 forested regions of their tropical and humid classes, respectively.

673 **5.2 How do regions with similar controls compare to ground-based** 674 **recharge estimates?**

675 In Africa, Recharge Landscapes with greater long-term mean annual recharge rates are also
676 more efficient in converting precipitation to recharge, as shown by the higher long-term mean

677 recharge ratio estimates. We do not know whether this relationship is found across other
678 continents or regions as previous studies investigating the controls on ground-based recharge
679 estimates across large spatial scales assess the spatial variability of annual recharge rates only
680 (Moon *et al.*, 2004; Mohan *et al.*, 2018; Moeck *et al.*, 2020; MacDonald *et al.*, 2021).

681 Investigating how recharge signatures interact in space allowed us to advance our
682 conceptualisations of recharge processes across Africa. Though comparative hydrology is
683 only just starting to be recognised by observational investigations within the groundwater
684 community (Haaf *et al.* 2020; Heudorfer *et al.* 2019), it is well established within the surface
685 water community (Addor *et al.* 2018; Sawicz *et al.* 2011, 2014) and has already been used in
686 global scale groundwater investigations using global scale modelling products (Cuthbert *et*
687 *al.*, 2019a).

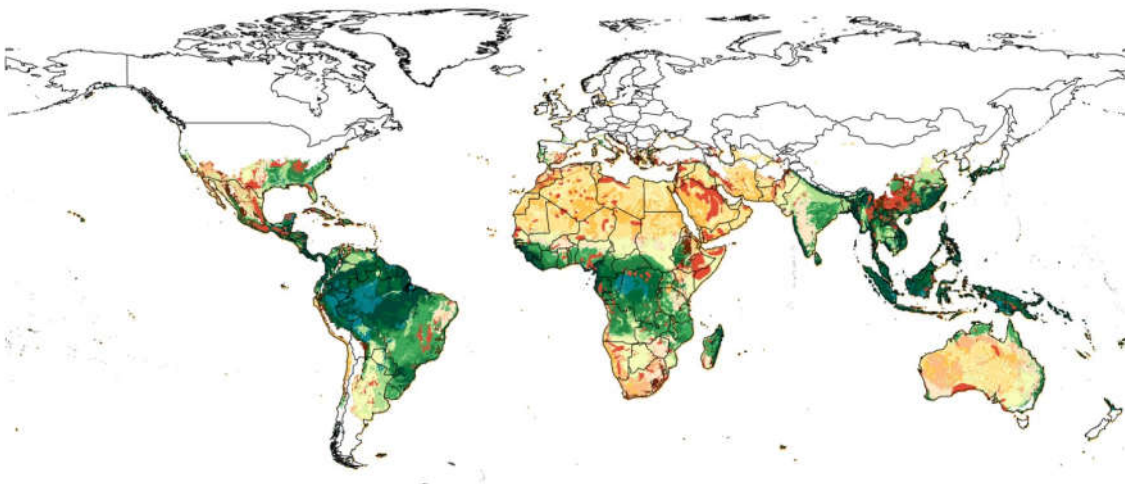
688 Even though we can explain the variability of ground-based estimates of annual recharge and
689 recharge ratio between different Recharge Landscapes, we have very limited ability to
690 explain why they vary within Recharge Landscapes using global datasets. Wet tropical and
691 Wet tropical forest Recharge Landscapes receive higher rates of annual recharge and are also
692 more efficient in converting precipitation to recharge than Dryland and Desert Recharge
693 Landscapes, as shown by the higher recharge ratio estimates in these places. This is not
694 surprising, as heavy seasonal, monthly and daily rainfall is already known to be important for
695 recharge processes in both tropical and dry regions of Africa (Döll and Fiedler 2008;
696 Jasechko and Taylor 2015; Owor *et al.* 2009; Taylor, Todd, *et al.* 2013). Furthermore, in
697 agreement with Taylor *et al.* (2013), we find that mean annual recharge ratios in Dryland
698 Recharge Landscapes, increase with monthly variability in P-PET. However, interactions
699 with other large-scale physical or biological indices offer little further explanation for why
700 ground-based estimates of annual recharge and recharge ratio vary within individual
701 Recharge Landscapes. For the most part, our inability to explain the spatial variability of

702 ground-based recharge estimates within Recharge Landscapes stresses the limitations of
703 global datasets for describing the complex interactions between landscape properties and how
704 they control more local recharge processes. It could also be attributed to the scale differences
705 in the resolution of our classification and the representative areas of each of our ground-based
706 estimates. Previous studies trying explain the spatial variability of recharge processes at
707 continental and global scales also mostly establish relationships with broad climate and eco-
708 hydrological patterns (Jasechko *et al.*, 2014; Cuthbert *et al.*, 2019b; MacDonald *et al.*, 2021).
709 Furthermore, MacDonald *et al.* (2021) were also unable to explain the more regional/local
710 spatial variability in ground-based estimates of recharge using global datasets. More
711 specifically, they found that there are spatial correlations in long-term average recharge rates
712 across Africa up to distances of 900 km, which cannot yet be explained by environmental
713 properties. Ultimately, this suggests a gap between what we can learn from local insight and
714 from large scale regionalization, regarding the interaction of environmental properties and
715 their control over recharge processes. This potentially has wider implications as global-scale
716 models which are frequently used to estimate groundwater recharge at these scales (Döll and
717 Fiedler 2008; Wada *et al.* 2010), typically rely upon assimilating global datasets for climate
718 forcing, characterising the land surface and model parameterisations (Telteu *et al.* 2021).
719 Nonetheless, understanding which recharge controls are currently identifiable using these
720 methods could help guide the evaluation and future development of continental scale models
721 (Gleeson *et al.* 2021). For example, recharge ratio estimates across the Sahara from PCR-
722 GLOBWB are typically greater than 0.2 (Jasechko *et al.* 2014), whilst our analysis shows
723 they are mostly below 0.04.

724 **5.3 Looking ahead**

725 Given the limited explanatory power of global datasets as shown in our and other previous
726 studies, it is likely that continental and global scale modelling of groundwater recharge can

727 benefit from the implementation of landscape-based conceptualisations of recharge processes
728 and controls (Gao et al. 2018). Hartmann et al. (2015) showed (for carbonate rock regions
729 across Europe and Northern Africa) that even relatively simple process conceptualizations
730 capture main differences in recharge dynamics between different large landscape groups.
731 Such conceptual models characterize largely our prior understanding of groundwater recharge
732 in different landscapes. This is likely to be particularly important in data sparse regions where
733 we cannot reasonably rely upon model parameterisation schemes that rely heavily on the
734 reliability of soils and other data (Wagener et al. 2021). Adding information through the
735 definition of simple system conceptualizations, would enable us to further combine expected
736 hydrologic behaviour of the landscape with widely available datasets (e.g. Cuthbert *et al.*,
737 2019b). By focussing on regionally dominant recharge controls, we can develop more
738 parsimonious mathematical models that are also more appropriate for the data scarcity found
739 in many places (Sarrazin *et al.*, 2018), or specific hydrologic processes of most relevance
740 (Quichimbo et al. 2021).



741
742 Figure 8. Application of the recharge landscape classification framework to domains outside of the study region. We used a
743 random forest to transfer our Recharge Landscape Units across the rest of the world, with the previously discussed recharge
744 control indices acting as predictor variables. The random forest model is an ensemble of 25 classification trees each with a
745 maximum of 400 decision splits. The model was trained on data points in Africa which were randomly selected with
746 replacement from a sub-sample of 70% of the Africa data points ('in-bag'). Model testing on 'out of bag' data points found a
747 misclassification rate of just 4%. Areas shown in white are significantly dissimilar to the study region. The criterion for this
748 separation was having mean temperatures below 13.5°C or above 35.5°C and snow fractions above 0.1. We estimated snow
749 fractions by using a simple temperature threshold. Precipitation on days with an average temperature below 1°C is regarded
750 as entirely snowfall whereas it is entirely rainfall on days with an average temperature above 1°C (Berghuijs et al., 2014).

751 We use a global gridded dataset of daily temperature provided by the Climate Prediction Center, NOAA
752 (NOAA/OAR/ESRL PSD). Further details are provided in the supplemental information.

753

754 The value of comparative hydrology in this context could lie in identifying regions of
755 similarity beyond the direct study domain. As discussed here, specific studies with ground-
756 based estimates of groundwater recharge are rare – certainly across Africa. Figure 8 shows
757 how the classification approach introduced here would classify other regions of the world if
758 applied globally. All areas shown in white are significantly dissimilar to our study domain
759 and hence unsuitable for comparison. However, areas in colour map onto some areas in our
760 domain and thus offer the potential for transferability of knowledge gained from outside our
761 direct study domain. For example, studies in karst regions (shown in red) might complement
762 the rather sparse ground-based measurements available inside Africa, thus offering an
763 opportunity to expand on existing datasets like that compiled by MacDonald et al. (2021).

764 **6 Conclusions**

765 We set out to study the variability of groundwater recharge across Africa through the use of a
766 classification of groundwater recharge controls as landscape elements, utilising global
767 datasets to characterize our *a priori* understanding following an extensive literature review.
768 Our final classification consists of 15 recharge landscape units which are similar across the
769 11 indices we used to describe recharge controls across the continent. We aggregated these
770 Recharge Landscape Units into four larger Recharge Landscapes, including Desert, Dryland,
771 Wet tropical, and Wet tropical forest, which broadly agrees with classifications by Olson *et*
772 *al.* (2001) and Jasechko *et al.* (2014). Karstic environments are treated separately, scattered
773 across each of the Recharge Landscapes we have found.

774 A classification approach has allowed us to consolidate most of the findings from previous
775 studies into a spatial representation of expected recharge controls across the African

776 continent. Much of our previous understanding of recharge processes in Africa was point or
777 plot based, originating from the case studies which have assessed recharge processes and
778 controls throughout the region. We hypothesize that the small number of Recharge
779 Landscapes needed to characterize the broader recharge controls of the African landscape, is
780 explained by the dominance of climatic controls, likely connected with the co-evolution of
781 vegetation, soils, and topography. These Recharge Landscapes were useful in organising
782 ground-based estimates of annual recharge and recharge ratio. Yet, in exception of Dryland
783 Recharge Landscapes, we were not able to explain the variability of estimated recharge
784 signatures within each of the Recharge Landscapes using global datasets alone.

785 There is still considerable variability in ground-based estimates of groundwater recharge
786 which cannot yet be explained using global datasets. This result highlights the limits of using
787 global datasets to decipher the complex interactions of landscape properties in controlling
788 recharge processes. Nonetheless, future data-based modelling of groundwater recharge at
789 continental scales could be advanced by using methods which explore the relationships
790 between controls and recharge within regions of similarity, instead of across the entire
791 continent (MacDonald et al. 2021). Further advancement is also likely to come from the
792 development of system conceptualizations which allow us to add more information than that
793 embedded in global datasets (Wagener et al. 2021). This would lead to a convergence of top-
794 down strategies (such as classification) with other more bottom-up approaches like the one
795 taken by Cuthbert *et al.* (2019b). Further expanding the study domain using similarity
796 principles might offer a strategy for expanding existing strategies. Furthermore, considering
797 the co-evolution of multiple landscape properties could help further separate the
798 hydrologically relevant behaviour of different places (Troch et al. 2013), which in turn could
799 help the predictive ability of global datasets used in model parameterisations. Currently such

800 expected hydrologic behaviour (derived from literature reviews), is only considered through
801 the definition of appropriate predictor variables.

802 Finally, as meta-analysis databases become more common in continental and global scale
803 hydrological studies (Moeck et al. 2020; Wang et al. 2020), we would like to stress the
804 importance of thorough quality assurance in their initial development. Our findings from
805 these studies depend upon strong underlying datasets and it is unlikely future studies will
806 assess the quality of these datasets when investigating or expanding upon them. For the same
807 reasons, the initial development of these databases should also ensure that additional meta-
808 information is comprehensive.

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