

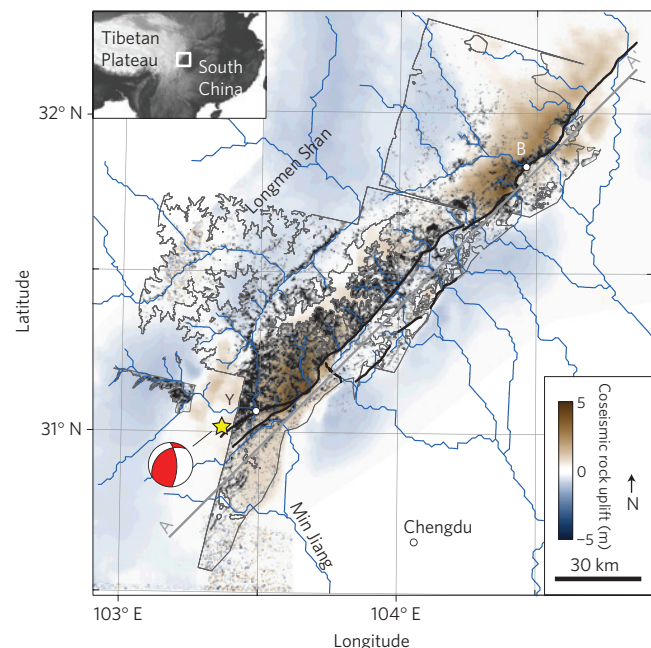
# Mass wasting triggered by the 2008 Wenchuan earthquake is greater than orogenic growth

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Shallow earthquakes are the primary driver of rock uplift in mountain ranges<sup>1</sup>. However, large shallow earthquakes also trigger widespread, coseismic landslides that cause significant but spatially heterogeneous erosion<sup>2–4</sup>. The interplay between rock uplift and the distribution and magnitudes of coseismic landslides thus raises a fundamental question as to whether large earthquakes and their associated landslides create or destroy mountainous topography. The 2008  $M_w$  7.9 Wenchuan earthquake in Sichuan, China triggered more than 56,000 landslides<sup>5</sup>, with a spatial distribution that was only partly related to the pattern of tectonic deformation<sup>6</sup>. Here we examine the potential changes in orogen volume using landslide area–volume scaling relationships<sup>4,7</sup> applied to high-resolution satellite imagery. We estimate that coseismic landsliding produced  $\sim 5\text{--}15\text{ km}^3$  of erodible material, greater than the net volume of  $2.6 \pm 1.2\text{ km}^3$  added to the orogen by coseismic rock uplift<sup>8</sup>. This discrepancy indicates that, even if only a fraction of the landslide debris is removed from the orogen over the likely  $\sim 2000\text{--}4000\text{ yr}$  earthquake return period<sup>6</sup>, the Wenchuan earthquake will lead to a net material deficit in the Longmen Shan. Our result challenges the widely held notion that large dip-slip or oblique-slip earthquakes build mountainous topography, and invites more careful consideration of the relationships between coseismic slip, mass wasting and relief generation.

It is axiomatic that earthquakes build topography through repeated vertical displacements<sup>1</sup>, yet large earthquakes are also a primary trigger of landslides<sup>2</sup>, which play a dominant role in the competition between tectonic and surface processes that drive mountain belt evolution<sup>9–12</sup>. Recent work<sup>2–4,7</sup> has shown that landslides are capable of generating sustained high rates of erosion (of the order  $1\text{--}10\text{ mm yr}^{-1}$ ), which poses a challenge to our understanding of how mountainous topography is generated: if the volume of erodible sediment produced by earthquake-triggered landsliding exceeds the coseismically-generated rock volume added to the orogen, then—assuming that this sediment is evacuated from the orogen by other erosional processes—the volume and mean elevation of the orogen must decrease. The relative roles of large earthquakes in generating coseismic rock uplift and facilitating landslide erosion<sup>13</sup> are thus critical for understanding the balance between crustal advection and denudation.

The  $M_w$  7.9 Wenchuan earthquake of 12 May 2008 in Sichuan Province, China, is ideal for examining the relationships between landsliding and orogen evolution because of its large magnitude, the steep regional topography, and the widespread occurrence of coseismic landsliding<sup>5,14</sup>. The earthquake occurred



**Figure 1 | Coseismic uplift and landslides triggered by the Wenchuan earthquake.** Black polygons show individual landslides. Heavy black lines show surface rupture traces<sup>18</sup> and the star indicates the epicentre. Grey outlines show the extent of imagery used in landslide mapping. Background is the coseismic rock uplift field based on SAR analysis, modified from de Michele and colleagues<sup>8</sup>. Heavy grey line shows the rupture-parallel section line onto which the results are projected. Beichuan (B); Yingxiu (Y).

in the Longmen Shan mountain range, which is underlain by a complex lithological assemblage comprising Proterozoic granitic massifs, a Palaeozoic passive margin sequence, a thick Triassic–Eocene(?) foreland basin succession, and minor exposures of poorly-consolidated Cenozoic sediment<sup>15</sup>. The faults in the Longmen Shan originated in the Late Triassic<sup>16</sup> and have remained active into the Quaternary as dextral-thrust oblique-slip faults<sup>17</sup>. The earthquake involved  $>10\text{ m}$  of oblique dextral-thrust surface slip on the Beichuan and Pengguan faults<sup>6,18</sup> (Fig. 1). Inversion of GPS and InSAR data<sup>6</sup> coupled with field observations<sup>18</sup> shows that the magnitude and proportion of dextral strike-slip and thrust dip-slip fault displacement varied significantly along the rupture trace, with two distinct zones of concentrated slip and moment release near Yingxiu and Beichuan (Fig. 1).

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To constrain landslide erosion, coseismic and immediate postseismic landslides were mapped within an area of 13,800 km<sup>2</sup> in the Longmen Shan using high-resolution satellite imagery collected within 30 days of the earthquake (see Methods). We resampled the raw landslide inventory data into landslide density  $P_{ls}$ :

$$P_{ls} = A_{ls}/A_t \quad (1)$$

where  $A_{ls}$  is the area of all landslides within a chosen window size  $A_t$  (ref. 19). Values of  $P_{ls}$  calculated with equation (1) vary from >60% (with  $A_t = 1 \text{ km}^2$ ) near the epicentre to 0% in the low-relief Sichuan Basin (Fig. 1).  $P_{ls}$  also varies significantly along strike, with high values along the Min Jiang valley near Yingxiu (Fig. 1) and secondary clusters to the northeast, particularly associated with major transverse river valleys. This partly, but not fully, reflects along-strike variations in surface rupture<sup>18</sup>. Strong variations in  $P_{ls}$  between different lithologies were noted by Dai *et al.*<sup>5</sup>, along with complex relationships between  $P_{ls}$  and the distance from the earthquake source. Given that landslide occurrence is not solely tied to coseismic deformation, there is potential for mismatch between patterns and volumes of tectonic rock uplift and landslide erosion.

Understanding the balance between tectonic and mass wasting processes in the Wenchuan earthquake requires a scaling relationship to convert individual landslide area  $A_i$  to total volume  $V_{ls}$ :

$$V_{ls} = \sum_1^n \alpha A_i^\gamma \quad (2)$$

where  $n$  is the number of landslides and the scaling parameters  $\alpha$  and  $\gamma$  are constants that vary with setting and hillslope process (for example bedrock or shallow landslides). We applied equation (2) using published scaling parameters<sup>4,7</sup> as well as those derived from field measurement of 41 landslides in the study area. The results (Table 1) are strikingly consistent and place first-order constraints on the likely volume of material involved. Application of a global best-fit relationship for all landslide types from Larsen *et al.*<sup>4</sup> with  $\gamma = 1.332 \pm 0.005$  yields  $V_{ls} = 5.73 + 0.41/\gamma - 0.38 \text{ km}^3$ . A global best-fit relationship for bedrock landslides from Larsen *et al.*<sup>4</sup> ( $\gamma = 1.35 \pm 0.01$ ) and a relationship derived from field measurements ( $\gamma = 1.388 \pm 0.087$ ) both yield similar values of  $V_{ls} \approx 9 \text{ km}^3$ , whereas a global relationship from Guzzetti *et al.*<sup>7</sup> yields  $V_{ls} = 15.2 + 2.0/\gamma - 1.8 \text{ km}^3$ . The predicted volumes in Table 1 are minima, because the images span most but not all of the surface rupture (see Methods), but are consistent with spatially-averaged denudation of 0.42–1.1 m over the 13,800 km<sup>2</sup> mapped area. Conversion of these estimates to landslide erosion rates requires knowledge of the recurrence intervals of large landslide-triggering earthquakes on the Beichuan fault, but these are poorly constrained by limited dating at a few widely-spaced trench sites<sup>20,21</sup> or inferred rates of strain

accumulation<sup>6</sup>. Assuming plausible recurrence intervals of 2,000–4,000 yr (refs 6,20) yields a long-term, spatially-averaged erosion rate due to landsliding alone of 0.1–0.6 mm yr<sup>-1</sup>, similar to the pre-earthquake total erosion rates of 0.2–0.6 mm yr<sup>-1</sup> in the eastern Longmen Shan estimated from cosmogenic nuclide analyses over similar millennial timescales<sup>22</sup>.

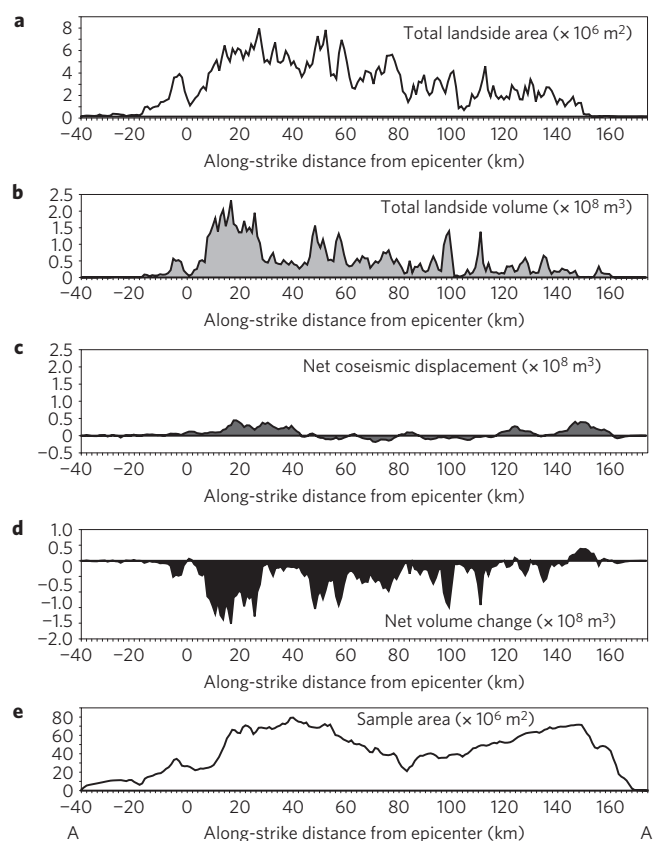
These landslide volume estimates can be compared with the volume of material added to the orogen in the earthquake through coseismic rock uplift. de Michele *et al.*<sup>8</sup> inverted ascending and descending mode Synthetic Aperture Radar (SAR) data (see Methods) to obtain the three-dimensional surface displacement vectors at ~350 m intervals across the region (Fig. 1). We sum the vertical component of these data (equation (3)) over the area of our landslide mapping to obtain a net positive volume gain  $V_t = 2.6 \pm 1.2 \text{ km}^3$ . This is more than one standard error less than all estimates of landslide volume (Table 1), and implies that the earthquake added much less volume to the Longmen Shan than was potentially released by landsliding (Fig. 2). There are, however, two important caveats to this direct comparison. First, the SAR data were obtained between November 2006 and August 2008, and thus record surface change due to coseismic and postseismic landslides as well as coseismic and postseismic deformation. Landsliding affects only about 4% of the 13,800 km<sup>2</sup> mapped area, however, minimizing the effect of landsliding on  $V_t$ . Also, disruption of the ground surface by landsliding causes local incoherence in the SAR analysis, and incoherent pixels are not used in the calculation of surface displacements<sup>8</sup>. The displacement magnitudes and directions determined from the inversion closely match field observations<sup>8,18</sup>, suggesting that at the orogen scale the displacement estimates are not strongly biased by landslide-induced surface change. Second, and more significantly, estimated landslide volume does not necessarily equate to eroded volume; conversion to an orogen-scale erosion rate requires that the landslide debris be efficiently flushed from the orogen<sup>13</sup>. Although there was some sediment storage along major Longmen Shan river valleys before the earthquake, the overall preponderance of bare-bedrock hillslopes and general lack of thick (>100 m) sediment stores<sup>22,23</sup> suggest that coseismic landslide debris is likely to be efficiently removed over the entire earthquake cycle, but the lack of pre- and post-earthquake sediment discharge data prevents us from quantifying the rate of removal<sup>13,24</sup>.

Thus, if hillslope and fluvial processes can remove the Wenchuan landslide debris before the next large landslide-triggering earthquake, then the earthquake will likely have caused a significant net volume loss from the orogen. How does this imbalance affect the growth of topography in the Longmen Shan? We stress that our results are an instantaneous measure of the competition between erosional and tectonic processes and bear only indirectly on the long-term volumetric balance that defines an orogen<sup>11</sup>. It is possible that the range is in topographic decay, as suggested by Godard *et al.*<sup>25</sup>, with rates of erosion outpacing those of rock uplift,

**Table 1 | Landslide scaling relationships and volume estimates.**

Relationship*	$\alpha$	$\gamma$	Volume <sup>†</sup> (km <sup>3</sup> )	Mean erosion (m) <sup>‡</sup>	Erosion rate (mm yr <sup>-1</sup> ) <sup>§</sup>	Reference
L1 (all landslides)	0.146	1.332 ± 0.005	5.73 + 0.41/γ - 0.38	0.42	0.1–0.2	4
L2 (all bedrock landslides)	0.186	1.35 ± 0.01	9.36 + 1.39/γ - 1.21	0.68	0.2–0.3	4
L3 (mixed Himalayan landslides)	0.257	1.36 ± 0.01	14.9 + 2.2/γ - 1.9	1.1	0.3–0.5	4
G (all landslides)	0.074	1.450 ± 0.009	15.2 + 2.0/γ - 1.8	1.1	0.3–0.6	7
Field measurements	0.106	1.388 ± 0.087	9.08 + 22.2/γ - 6.35	0.66	0.2–0.3	This study

\*L1: global relationship for all landslides from Larsen *et al.*<sup>4</sup>; L2: global relationship for all bedrock landslides from Larsen *et al.*<sup>4</sup>; L3: relationship for mixed bedrock and soil landslides in the Himalaya from Larsen *et al.*<sup>4</sup>; G: global relationship for all landslides from Guzzetti *et al.*<sup>7</sup>. <sup>†</sup>Uncertainties are expressed by applying equation (2) with ±1 s.d. error on  $\gamma$ . <sup>‡</sup>Mean erosion represents the average lowering of the ground surface due to landsliding and is calculated by dividing the estimated volume by the total study area ( $A_{map}$ ). <sup>§</sup>Spatially-averaged landslide erosion rate is determined by dividing mean erosion range by the approximate earthquake recurrence interval of 2,000–4,000 yr (refs 6,20).



**Figure 2 | Along-strike variations in landslide occurrence and coseismic displacement.** All data are projected onto a rupture-parallel line A–A' (Fig. 1) at 1 km intervals. **a**, Total area of landslides within each 1-km wide strip. **b**, Landslide volume derived from the global bedrock landslide scaling relationship<sup>4</sup> applied to individual landslides within each 1-km wide strip; other relationships show similar patterns. **c**, Net coseismic volume change<sup>8</sup> in each 1-km wide strip. **d**, Net volume change determined by subtracting landslide volumes from coseismic volume change. **e**, Along-strike distribution of sample area covered by satellite imagery. Local minima in landslide area and volume are not correlated with small sample areas.

although this model remains to be tested through more focused thermochronological investigation. A second possibility is that some of the long-term rock uplift is accumulated through interseismic deformation<sup>26</sup> or afterslip<sup>27,28</sup>, although the latter mechanism in particular has tended to yield a small fraction of the coseismic displacement. Alternatively, an important fraction of long-term rock uplift may occur in more frequent smaller, or deeper, earthquakes that generate lower peak ground acceleration values<sup>29</sup> and trigger a much lower volume of landslides<sup>2,3</sup>. In that scenario, large or shallow earthquakes would serve primarily to reduce the tectonic topography constructed by smaller or deeper earthquakes and maintain hillslopes at threshold gradients. In support of this idea, Ouimet<sup>30</sup> noted that short-term ( $10^3$  yr) erosion rates in the Longmen Shan are  $0.2\text{--}0.3\text{ mm yr}^{-1}$ , lower than rates over Myr timescales ( $0.5\text{--}0.7\text{ mm yr}^{-1}$ ; ref. 25), and suggested that large earthquakes allow erosion rates to catch up with longer-term rock uplift rates. Climatic conditions will also probably play a role in determining the precise pattern and volume of landslides in response to a given earthquake; given the order-of-magnitude agreement between our estimated rates of landslide erosion and both long- and short-term regional erosion rates, however, temporal variations in climate are unlikely to exert significant changes on the volume balance. A further possibility is that the balance between rock uplift and landslide erosion in the Wenchuan earthquake was anomalous and cannot

be extrapolated over multiple earthquake cycles. It seems likely that earthquakes with a larger component of shortening will lead to a net addition of rock volume, whereas dominantly strike-slip events will cause a net loss due to widespread landsliding but limited rock uplift. Dextral and thrust slip in the Wenchuan earthquake were highly partitioned between different fault strands<sup>18</sup>, and the ratio of rock uplift to lateral slip on those strands may vary between earthquakes<sup>17</sup>. Large differences in that ratio in successive earthquakes would thus be expected to yield major temporal variations in the net volume balance, even if the pattern and total volume of landsliding remained the same. In any case, the apparent and provocative mismatch between tectonic and erosional volumes involved in the Wenchuan earthquake points to a need for much greater understanding of the role of large earthquakes in setting regional erosion rates and long-term patterns of orogen evolution.

## Methods

**Landslide mapping.** We developed a semi-automated detection algorithm using EO-1 and SPOT 5 imagery for objective mapping of individual landslides (see Supplementary Information). Landslide areas were extracted from EO-1 imagery using an intensity threshold and a  $20^\circ$  gradient mask to remove false positives in valley floors; independent work<sup>5</sup> shows that areas with a gradient of  $<20^\circ$  have very low landslide densities. Unsupervised classification with a  $20^\circ$  gradient mask was used to delineate landslide areas in SPOT 5 imagery. A series of feature-oriented filters were applied to remove false positives produced by roads and fields, and the map was visually inspected and corrected. This resulted in a landslide map with a total area of  $13,800\text{ km}^2$  (Fig. 1) that covers 150 km of the 225 km surface rupture<sup>6,18</sup>, so that the total landslide area and volume calculated here are minimum values. Comparisons with field evidence<sup>18</sup>, fault models<sup>6</sup>, and SAR analysis<sup>8</sup>, and with independent landslide maps compiled by hand from imagery and aerial photographs<sup>5</sup>, however, suggest that the mapped area covers the majority of coseismic slip and represents a significant sample of the main impact zone of the earthquake.

**Coseismic volume estimation.** By combining C- and L-band space-borne SAR amplitude data, de Michele *et al.*<sup>8</sup> derived the three-component coseismic surface displacement field due to the Wenchuan earthquake. Here we used the up, or vertical, component to calculate the net coseismic volume change in the Longmen Shan, ignoring elevation change in the low-relief Sichuan Basin (Fig. 1). Within the area of the Longmen Shan covered by the landslide mapping (Fig. 1), we calculated the net volume change as

$$V_t = A \sum_{x=1}^n U_x \quad (3)$$

where  $A$  is the cell area,  $U_x$  is the vertical displacement for each cell, and  $n$  is the number of cells, yielding  $V_t = 2.6 \times 10^9\text{ m}^3$ . The standard deviation of the difference between the displacements and ground truth data is not a good statistical indicator of the uncertainty in  $V_t$ , because random (uncorrelated) errors are likely to lead to a negligible net contribution to the total volume over the mapped area. Instead, we estimated the uncertainty in  $V_t$  by evaluating the magnitude of statistical variation in  $U_x$  within a non-deforming area far from the earthquake rupture. We chose a  $36\text{ km} \times 36\text{ km}$  area in the Sichuan basin, 45 km away from the fault rupture, containing a high level of noise (mean of 0 m and standard deviation of 1.5 m). We extracted 30 profiles, each 36 km long, within the selected area, and used the least squares method to fit each profile by linear regression. Because the  $y$ -intercept value influences the volume estimation beneath each  $36\text{ km} \times 1\text{ pixel}$  area, we examined the  $y$ -intercept parameter for each of the profiles and calculated the Root Mean Square Error (RMSE) between the 30  $y$ -intercept parameters and the ground truth data. This yields an RMSE of 0.10 m; when applied over the entire mapped area, this is equivalent to an estimated uncertainty of  $1.2 \times 10^9\text{ m}^3$  on  $V_t$ .

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### Author contributions

R.N.P. and S.W. conducted the landslide mapping and analysis. A.L.D., S.W., Y.L., R.H. and D.N.P. collected field data on the rupture and landslide characteristics. M.d.M. derived the tectonic mass flux. A.L.D. conceived the idea and wrote the paper with input from R.N.P., N.J.R., D.N.P. and M.d.M.

### Additional information

The authors declare no competing financial interests. Supplementary information accompanies this paper on [www.nature.com/naturegeoscience](http://www.nature.com/naturegeoscience). Reprints and permissions information is available online at <http://www.nature.com/reprints>. Correspondence and requests for materials should be addressed to A.L.D.