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# Free gas accumulations in basal shear zones of mass-transport deposits (Pearl River Mouth Basin, South China Sea): An important geohazard on continental slope basins

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## Abstract

Free gas is an important trigger of instability on continental slopes, and resulting mass-wasting strata can potentially form competent seals to hydrocarbon accumulations. This work uses two high-quality 3D seismic volumes to investigate fluid accumulations at the base of mass-transport deposits in the Pearl River Mouth Basin, South China Sea. In parallel, IODP/ODP borehole data are used to document the petrophysical character of mass-transport deposits formed in similar continental-slope environments to the South China Sea. The interpreted data show gas accumulations as comprising enhanced seismic reflections that are discordant, or vertically stacked, below mass-transport deposits with chaotic seismic facies. Gas was accumulated in basal shear zones of mass-transport deposits in response to differences in capillary pressure and porosity. Free gas in Zone A covers an area of at least 18 km<sup>2</sup>. In Zone B, the free gas is sub-circular in plan view and covers an area of 30.58 km<sup>2</sup> for a volume of sediment approaching 1.5 km<sup>3</sup>. This work is important as it shows that vertical migration of gas is not significant in mass-transport deposits from the Pearl River Mouth Basin, but up-dip migration along their basal shear zones is suggested in multiple locations. As a result, free gas can pinch-out laterally to extend 1-2 km beyond these same basal shear zones. As a corollary, we show that free gas accumulations below mass-transport deposits comprise an important geohazard and should be taken into account when drilling continental-slope successions in both the South China Sea and continental margins recording important mass wasting. Strata charged with free gas form weak layers, hinting at a novel trigger of retrogressive slope failure on continental slopes worldwide.

**Keywords:** South China Sea, fluid migration, free gas, mass-transport deposits, seal units, geohazards.

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## 1. Introduction

Mass-transport deposits (MTDs) can shape both passive and active continental margins, transporting large volumes of sediment to continental-slope basins (e.g. Haflidason et al., 2004; Alves et al., 2009; Alves, 2010; Soares et al., 2012). They are of recognised importance as local geohazards when drilling (e.g. Varnes, 1978; Locat and Lee, 2002; Krastel et al., 2006; Masson et al., 2006; Lo Iacono et al., 2012), and originate when downslope-directed forces exceed resisting forces on a continental slope (e.g. Hampton et al., 1996; Masson et al., 2006; Dugan, 2012). Multiple factors such as high sedimentation rates (e.g., Dalla Valle et al., 2013; Noda et al., 2013), methane hydrate dissociation (e.g., Laberg and Vorren, 2000; Maslin et al., 2005), release of free gas (e.g., Bünz et al., 2005; Best et al., 2003; Dugan and Flemings, 2000), tectonic activity (e.g., Hance, 2003; Chadwick et al., 2012; Laberg et al., 2014), and slope oversteepening (e.g., Strozyk et al., 2010; Ikari et al., 2011) comprise important triggers for the generation of MTDs. In parallel, submarine creep zones act as a precursor of large-scale slope failure (Li et al., 2016).

Failed strata are transported downslope along a basal shear zone of variable thickness and lithology when submarine slope instabilities occur (Alves and Lourenço, 2010). As a result, three distinct domains are identified within MTDs based on their deformation styles (Martinsen, 1994; Lastras et al., 2002; Bull et al., 2009). They are: a) a headwall domain characterised by extensional deformation, including headwall scarps, extensional ridges and blocks; b) a translational domain comprising the main translated body of the MTD (including sidewall scarps, strike-slip deformation, ramps and flats, grooves and striations, blocks, slump folds, longitudinal shears and secondary flow fabrics), and c) a toe domain dominated by compressional deformation, including pressure ridges, small-scale folds and thrusts (Prior et al., 1984; Bull et al., 2009). Outcrop data from SE Crete has shown that glide planes of submarine landslides comprise a deformation zone in which strata is deformed, sheared and fractures to enhance porosity and fluid migration (e.g. Alves and Lourenço, 2010). Importantly, large volumes of fluids can be focused in vertical and horizontal surfaces within, or immediately below displaced blocks and MTDs, with these same blocks marking the sudden release of

overburden pressure near the seafloor (Sultan et al., 2004; Flemings et al., 2008; Alves, 2015). Notwithstanding this work, the roles of MTDs on fluid migration after their emplacement have seldom been addressed in the published literature.

In the Pearl River Mouth Basin, small-scale MTDs are observed on a canyon-incised continental slope (e.g. Dickinson et al., 2012; Gong et al., 2013; He et al., 2014; Li et al., 2016). Here, large-scale MTDs are grouped together to form the Baiyun Submarine Slide Complex (BSSC), and cover an area of more than 10,000 km<sup>2</sup>. They are imaged on multibeam bathymetry and 2D/3D seismic data on the continental slope off the Pearl River Mouth Basin (Sun et al., 2008; Li et al., 2014; Wang et al., 2014; Sun et al., 2017). The BSSC is composed of, at least, three vertically stacked MTDs, which were respectively emplaced at 0.19 Ma, 0.50 Ma and 1.59 Ma before present (Sun et al., 2017). Small-scale MTDs were probably caused by local tectonic oversteepening working in association with differences in the rheological behaviour of continental-slope strata (He et al., 2014; Li et al., 2016).

Recently, the origin of the BSSC was proposed to be more complex than previously assumed, as Wang et al. (2014) proposed that the Dongsha Tectonic Event as a key triggering factor of slope instability in the South China Sea. In parallel, Li et al. (2014) suggested that excess pore pressure aided by volcanic activity, tectonic oversteepening and gas hydrate dissociation, were the major triggering factors for the BSSC. This latter interpretation is possibly correct for parts of the BSSC, but there is no direct evidence to support them in the study area. After a detailed interpretation of new high-resolution 3D seismic data, the distribution of coarse-grained strata (turbidites) and associated seabed topographic highs, were found to be important in controlling the location of MTDs headwall scarps in parts of the South China Sea (Sun et al., 2017).

As the analysis of MTDs in the northern South China Sea is still in its infancy, and the importance of fluid migration in the triggering of MTDs has not been fully addressed, this paper focus on free gas accumulated below individual (Late Cenozoic) MTDs. A significant volume of free gas trapped by MTDs is clearly imaged in two high-quality 3D seismic surveys, providing an unique opportunity to explore their dual relationship. Hence, this work will address the following research questions:

- 1) What is the relationship between MTDs and underlying free gas in the Pearl River Mouth Basin?
- 2) What are the main paths for fluid migrating from below the studied MTDs?
- 3) What processes enhance the seal competence of MTDs in the Pearl River Mouth Basin?
- 4) What is the importance of the interpreted MTDs, and associated free gas, to future geohazard assessments in the Pearl River Mouth Basin and other continental margins?

## 2. Geological setting

The South China Sea is the largest and deepest marginal sea in the western Pacific Ocean and kept its present position since the Middle Miocene (Taylor and Hayes, 1983). Five major Cenozoic rift basins were formed along its northern margin, the Qiongdongnan, Pearl River Mouth, Yinggehai, Taixinan Basin, and Beibuwan Basins (Fig. 1). The structural evolution of these sedimentary basins is divided in three stages: a) an Eocene-Early Oligocene rift stage, b) a Neogene-Quaternary post-rift stage characterised by thermal subsidence, and c) a Late Oligocene transitional stage (Ru and Pigott, 1986; Gong and Li, 1997; Zhao et al., 2016) (Fig. 2). Except for regional tectonic movements, which occurred during the Late Miocene and caused local uplift and faulting (Lüdmann and Wong, 1999), tectonic activity was relatively moderate in the Pearl River Mouth Basin during the post-rift stage (Zhou et al., 1995) (Fig. 2).

The two 3D seismic surveys interpreted in this study are located on the continental slope off the Pearl River Mouth Basin (Fig. 1a). The Pearl River Mouth Basin lies in the central part of the northern South China Sea and occupies an area of  $17.5 \times 10^4 \text{ km}^2$  (Sun et al., 2012) (Fig. 1). It is the most favourable region for hydrocarbon exploration in the northern margin of the South China Sea (Gong and Li, 1997; Dong et al., 2009). In response to its structural evolution, depositional environments in the Pearl River Mouth Basin changed from lacustrine in its early evolution stages (Eocene-Oligocene rifting), to shallow marine (embayment), and finally open marine during Neogene-Quaternary post-rift (Zhu et al., 2009; Sun et al., 2012) (Fig. 2). Eocene-Lower Oligocene lacustrine strata (mudstone, black shale and coal) have a thickness of more

than 2000 m and comprise the main hydrocarbon source interval in the Pearl River Mouth Basin (Zhang and Huang, 1991; Huang et al., 2003). Deltaic depositional systems dominated the margins of the Pearl River Mouth Basin during the Early Miocene and deposited fine-grained strata with a total thickness of 2000–3500 m. These latter strata act as a regional seal for hydrocarbon prospects in the region (Zhu et al., 2009).

## 3. Data and methods

Two 3D seismic volumes acquired in two distinct regions (Zones A and B) are used in this study (Fig. 1). In Zone A, the interpreted seismic volume covers an area of  $2,000 \text{ km}^2$  and was acquired by the China Oilfield Services Limited in 2006. For the acquisition of the seismic volume were used a 3,000 m long streamer containing 240 channels (trace interval 12.5 m), and an array of tuned air guns with a total volume of  $8 \times 20 \text{ inch}^3$ , working at a shot interval of 25 m. The seismic data in Zone A is zero phased and has a frequency of 45 Hz. The 3D seismic volume has a bin dimension of  $12.5 \times 12.5 \text{ m}$  and a 2 ms vertical sampling interval. The vertical seismic resolution is 9 m for the depth of MTDs interpreted in this work.

The 3D seismic data in Zone B covers an area of  $300 \text{ km}^2$ . It was acquired by the China National Offshore Oil Corporation in 2008. The seismic acquisition system consisted of a 3,000 m-long, 120-channel streamer with a hydrophone group interval of 25 m and a sampling interval of 4 ms. The bin spacing of the time-migrated volume is  $12.5 \times 12.5 \text{ m}$ . The 3D seismic data is zero phased and has a dominant frequency of 35 Hz. The vertical seismic resolution of this 3D seismic survey is 11 m for the depth of MTDs interpreted in this paper.

There are no available borehole data in the study area and, therefore, a p-wave velocity of 1540 m/s was used in our thickness calculations based on data from ODP Site 1146, located 46 km to the east of Zone A (Figs. 1 and 3). The shallow strata drilled at ODP Site 1146 is dominated by fine-grained nannofossil clay (Fig. 3).

Seismic interpretation consisted in the mapping of the top and base of MTDs, computing and extracting variance volumes and slices, calculating RMS (root-mean square) amplitude maps, and computing isochron data on Schlumberger's Geoframe® 2012. After mapping the top

and base of different MTDs, their areas and volumes were also calculated on Geoframe® 2012.

In order to understand the petrophysical character of the MTDs in this study we compared ODP borehole data from the South China Sea with IODP and ODP data from the Gulf of Mexico and Amazon Fan, located in similar depositional environments to the study area, and presenting similar lithologies. Borehole data (including wireline logs) were obtained from the IODP/ODP database (<http://brg.ldeo.columbia.edu/logdb/files.php> and [http://www-odp.tamu.edu/publications/pubs\\_pr.htm](http://www-odp.tamu.edu/publications/pubs_pr.htm)) and provided by D.J.W. Piper (pers. comm.). The original log data was downloaded on Carbon 3.0. Sediment porosity ( $\phi$ ) in Site U1323 was calculated using the following equation:

$$\phi = (\rho_b - \rho_g) / (\rho_w - \rho_g) \quad (1)$$

where  $\rho_g$  and  $\rho_w$  are pore fluid density and solid-grain density, respectively (Dugan, 2012).

#### 4. Evidence of gas trapping at the base of MTDs in the Pearl River Mouth Basin

Enhanced negative seismic anomalies (ENSAs) are observed along the base of chaotic MTDs in Zone A (Figs. 4a-e). They occur around the MTDs' headwall scarps, but often continue into undisturbed strata to the east (i.e. upslope) of these same headwall regions (Figs. 4a-b, 4e). The ENSAs are discrete circular features, or present a linear distribution; they abruptly contact with non-enhanced seismic reflections (Fig. 4). To the west, they contact with high-amplitude strata with a thickness reaching 230 m, and onlapping seismic reflections being also observed (Fig. 4b). To the north and east, the interpreted ENSAs are bounded by low-amplitude features (Fig. 4e). However, their western limit is difficult to identify because of the predominance of high-amplitude strata in this same area. Thus, the black continuous line in Fig. 4b is the conservative limit of ENSAs to the west, providing a conservative area of  $\sim 18 \text{ km}^2$ , and a volume approaching  $0.44 \text{ km}^3$  (Fig. 4e and Table 1).

In Zone B, acoustic anomalies observed on seismic and RMS-amplitude data are vertically stacked (Figs. 5 and 6). These stacked enhanced seismic anomalies (SESAs) are trapped underneath chaotic MTDs (Figs. 5a and 5b) and

are sub-circular in plan view (Fig. 7). They occur along, or instead cross-cut, slope strata and seismic reflections that show a lateral inversion in polarity in these features (Figs. 6c-d). Several chaotic and dimmed seismic reflections sourced from deeper strata are observed directly below the SESAs (Figs. 5a-b, 6a). Their thickness is variable and shows four maxima that correlate with the underlying chaotic and dimmed seismic reflections (Figs. 5 to 7). The SESAs cover an area of  $30.58 \text{ km}^2$  (Fig. 5c; Table 1), and abruptly terminate to the west where they are bounded by a large normal fault (Figs. 5a-b).

Negative seismic anomalies observed at the base of MTDs are usually interpreted as basal shear zones, and are generated by differences in acoustic impedance between MTDs and underlying strata (Dugan, 2012). Therefore, the ENSAs observed in Zone A occur in basal shear zones as documented in Alves and Lourenço (2010) and Dugan (2012). Yet, they extend laterally beyond the MTDs (Figs. 4a-d), a character suggesting they are not basal shear zones through their whole extent. A sharp boundary at the gas hydrate stability zone (i.e., a bottom-simulating reflection or BSR) is another possible explanation for the origin of these negative acoustic anomalies. However, the ENSAs are not parallel to the seabed (Figs. 4a-d), indicating that they are not BSRs. Their abrupt contact with the surrounding strata is similar to the seismic character of free gas documented in multiple studies (e.g. Cartwright et al., 2007; Judd and Hovland, 2007; Sun et al., 2012) and, thus, the ENSAs observed in Zone A are interpreted as free gas. In contrast, the parallel high-amplitude seismic reflections in the western part of Fig. 4b and Fig. 4e are sub-parallel, and probably caused by strata with different physical properties to the imaged MTDs. They are interpreted as coarse-grained turbidites (Sun et al., 2017) i.e., likely regions with a high sand content as those documented by Gamboa et al. (2010) in SE Brazil.

The internal seismic character of SESAs in Zone B is distinct from that in ENSAs (free gas) and parallel high-amplitude seismic reflections (turbidites) in Zone A (Figs. 5 and 6). The SESAs are vertically stacked and show an inversion in seismic polarity, at places cross-cutting the slope strata (Figs. 6c-d). This character is typical of free gas accumulations (e.g. Judd and Hovland, 2007; Brown, 2011; Sun et al., 2012), and the SESAs in Zone B are also interpreted as comprising free gas in the

interior of porous sediment. The chaotic and dimmed seismic reflections directly below the free gas zone are interpreted as pathways (gas chimneys) for the migration of free gas (Figs. 5 to 7).

## 5. Seismic characteristics of gas-laden MTDs

### 5.1 Mass-transport deposits in Zone A

Shallow strata in Zone A are characterised by the interbedding of chaotic to parallel seismic reflections (Fig. 8). The chaotic seismic reflections in Zone A are similar to the seismic character of MTDs documented on many passive and active continental margins (e.g. Evans et al., 1996; Bøe et al., 2000; Imbo et al., 2003; Frey Martinez et al., 2005; Gee et al., 2006; Moscardelli et al., 2006; Bull et al., 2009; Omosanya and Alves, 2013; Gong et al., 2014; Li et al., 2014; Wang et al., 2014). Therefore, chaotic seismic reflections are interpreted as MTDs, here informally named MTDa. This MTDa extends beyond the coverage of 3D seismic data, with only its eastern part being well imaged (Fig. 4).

Apart from their chaotic seismic reflections, short sub-parallel seismic reflections are also observed within MTDa (Figs. 8b-c). They occur as structures accumulated close to headwall scarps or, in other regions of MTDa, as pointy features observed on seismic and coherence data (Fig. 9). Similarly to structures reported in Bull et al. (2009) and Alves. (2015), the pointy features are interpreted as linear blocks accumulated in the extensional domain of MTDa. The features are interpreted as rafted blocks that can occur at any zone within MTDs (Alves, 2015); 'ridge-like' structures are often observed and usually rise over, or pierce, the overlying strata (Figs. 8b-c). At places, onlapping seismic reflections occur at the flanks of these same ridges, indicating that the deposition of these onlapping strata occurred during, or immediately after, the emplacement of MTDa (Figs. 8b and 8d).

A steep headwall scarp is observed on seismic data, and has an elevation ranging from 60 to 95 m (Fig. 8). It is arcuate in plan view and extends from NW to SE (Fig. 9). At this same scarp, MTDa contacts with undeformed slope deposits that are observed to the east (Figs. 8c-d). In addition, a thin unit with near-constant thickness (~65.0 m) drapes and overlies MTDa.

MTDa has variable thickness, with an average value of

55.3 m (Table 1). The maximum thickness of MTDs is ~100 m and occurs against the headwall scarp (Fig. 8a). In detail, MTDa becomes thinner downslope from its headwall scarp, towards to west (Fig. 8a). MTDa extends beyond the 3D seismic survey to the west and south and cover an area of 69.5 km<sup>2</sup> (Table 1). Thus, 3.84 km<sup>3</sup> of remobilized strata are accumulated in Zone A. As both Zone A and the area around ODP Site 1146 are located in deep waters and far from the modern shelf break of the northern South China Sea, MTDa is likely composed of fine-grained deposits similar to those drilled at ODP Site 1146 (Fig. 3).

### 5.2 Mass-transport deposits in Zone B

Shallow strata in Zone B are wedge-shaped, with strata onlapping both the western and northern parts of the study area (Figs. 5a-b, 6a-b). Six seismic anomalies (informally named MTDb1 - MTDb6) are identified in Zone B. They have similar seismic characters to strata in MTDa and are interpreted to comprise MTDs (Figs. 5a-b, 6a-b). Locally, parallel seismic reflections are observed within these deposits, while sharp pointy features observed on coherent data and are interpreted as rafted blocks (Fig. 10). The uppermost two MTDs (MTDb1 and MTDb2) extend regionally to continue beyond Zone B (Figs. 10a, 11). The intermediate MTDs (MTDb3 and MTDb4) have smaller sizes and limited distributions compared to the other four MTDs. The lowermost MTDs (MTDb5 and MTDb6) are bounded to the west and north by, and onlap, continental-slope strata (Figs. 5a-b, 6a-b).

MTDb1 covers an area of 63.4 km<sup>2</sup> and has an average thickness of 27.2 m (Fig. 11; Table 1). Two thickness maxima are observed in the northeastern and central-western parts of MTDb1 (Fig. 11). Grooves caused by the scouring of rafted (remobilised) blocks, and often parallel to the flow directions of rafted blocks (Posamentier and Kolla, 2003; Moscardelli et al., 2006; Sun et al., 2011), are clearly identified at the base of MTDb1 (Fig. 10a). They strike roughly to the NW, indicating that strata were transported to the SE. Towards to the southern part of Zone B (downslope), MTDb1 becomes thinner and pinches out. MTDb2 has a similar distribution to MTDb1 and NW-SE grooves are also observed (Fig. 10b). However, MTDb3 and MTDb4 are limited to the central-northern part of the 3D seismic

survey (Fig. 12). In comparison, MTDb5 and MTDb6 have a similar distribution and cover areas of 51.4 km<sup>2</sup> and 46.9 km<sup>2</sup>, respectively (Fig. 12; Table 1). Rafted blocks and NW-trending grooves are also observed at the base of MTDb6 (Fig. 10c), indicating that all six (6) MTDs were sourced from the upper slope and moved towards the SE. MTDb5 and MTDb6 have a total thickness of ~55.8 m, and overlay the main body of free gas at depth (Fig. 11).

## 6. Discussion

### 6.1 Seal capacity of mass-transport deposits

The relationship between free gas and MTDs has been previously assessed in the literature, with free gas being usually proposed as an important trigger of MTDs (e.g. Bünz et al., 2005; Wang et al., 2014). However, the two examples in this study show that free gas is directly trapped underneath MTDs, and that these deposits form competent seal intervals to promote lateral migration of gas into undeformed slope strata. There are no wells drilled through the MTDs investigated in this paper and, thus, the petrophysical properties of MTDs cannot be directly assessed. Yet, the MTDs in the study area are likely fine-grained as deduced from units drilled at ODP Site 1146, and are similar to strata documented in the Amazon Fan (Piper et al., 1997; Shipp et al., 2004), Nankai Accretionary Wedge (Strasser et al., 2011), Gulf of Mexico (Dugan, 2012), Ulleung Basin (Riedel et al., 2012) and Arabian Sea (Pandey et al., 2015). The MTDs in these locations were drilled by ODP/IODP and exploration wells. Moreover, all MTDs documented in these regions showed similar petrophysical characteristics (e.g. Piper et al., 1997; Dugan, 2012; Riedel et al., 2012; Alves et al., 2014; Pandey et al., 2015), suggesting that the petrophysical character of these same deposits is predictable and probably associated with a common lithology. Hence, the petrophysical data acquired by previous ODP/IODP expeditions can be used in this study.

The shear deformation caused by the mass movement changes the petrophysical character of slope strata (Shipp et al., 2004; Dugan, 2012). The bulk density and porosity of MTDs in the Gulf of Mexico is ~7% higher and ~14% lower than the underlying and overlying undeformed strata, respectively (Fig. 13a). The denser (and lower porosity) strata in MTDs will cause the compressional wave and

resistivity to be higher than confining, non-disturbed sediment (Fig. 13b), with shear deformation also contributing to the loss of water in MTDs (Piper et al., 1997; Shipp et al., 2004). For instance, the water content of MTDs in the Amazon Fan is ~30% lower than normal strata (Fig. 13b). Therefore, the MTDs are more consolidated than the overlying and underlying normally depositional strata and have a larger capacity to serve as seals than undeformed strata, because of the higher bulk density, lower porosity and water content of MTDs. This could justify why the free gas is trapped underneath the MTDs observed in the study area.

In Zone A, vertical migration pathways below free gas (e.g. tectonic faults, gas chimneys and mud diapirs) are not observed and free gas is laterally connected to continuous high-amplitude strata (Figs. 4b and 4e). These strata are interpreted as coarse-grained turbidites similar to those reported in the Espírito Santo Basin (Gamboa et al., 2010). Thus, the free gas is probably sourced from deeper strata below the MTDs and migrates upwards along these porous strata (Fig. 14a). When the free gas reaches the MTDs is sealed and cannot vertically migrate, but can laterally migrate updip along the base of MTDs (basal shear zone) or along the porous strata just below these latter. This phenomenon generates the ENSAs observed beyond the scarp of MTDs in Zone A (Figs. 4b, 14a).

Free gas is also sourced from deeper strata in Zone B. However, it migrates upwards along gas chimneys or faults as observed in Figs. 5-6 and 14b. In parallel, the distribution of free gas in Zone B is mainly controlled by their migration pathways and accumulated at the top of gas chimneys and faults as observed in Fig. 7. The gas chimneys are likely composed of fractures and small faults, and provide efficient pathways for fluid migration (Cathes et al., 2010). The transported fluids (gas) are obstructed by the more consolidated MTDs in Zone B and trapped within porous strata below the MTDs (Figs. 14b).

### 6.2 Mass-transport deposits and their importance as offshore geohazards in the South China Sea

Mass-transport deposits are considered to be significant geohazards when triggered on continental slopes (e.g. Moscardelli et al., 2006; Migeon et al., 2011). The large volumes and long transporting distances of strata involved in their *en-masse* transport have enough power to destroy



offshore infrastructures or, at times, trigger large destructive tsunamis with the potential of sweeping through low-lying coastal zones. For example, the tsunami triggered by the Storegga slide impacted the coasts of the Norwegian Sea, northern North Sea and southeastern North Sea (e.g. Long et al., 1989; Smith et al., 2004; Bondevik et al., 2005; Friergaard et al., 2015). However, the importance of MTDs as geohazards after their emplacement is seldom assessed. This study considers two main implications of MTDs and trapped free gas for future geohazard assessments.

First, local overpressure induced by the free gas is a significant hazard during drilling, particularly when considering the large volumes of free gas identified underneath the two MTDs in this study. Considering that the two MTDs are more consolidated than undeformed slope strata, overpressure will build underneath them. The free gas and consequent overpressure will greatly affect borehole stability during drilling in such an hydrocarbon-rich basin as the Pearl River Mouth Basin.

Second, the free gas-charged strata possibly serve as weak layers for new slope failures. In this study, trapped free gas under over-consolidated MTDs migrates laterally along the basal shear zone, and likely through porous strata located just underneath the MTDs. This free gas later migrates into the undeformed strata beyond the scarps of MTDs (e.g. Zone A; Figs. 4a-d). Here, the gas-bearing strata could possibly serve as weak layers and the strata overlying them will fail under certain conditions, such as during regional earthquakes. Therefore, new slope failures will occur and will share the same basal surface with the previous MTDs. This process probably contributed to the retrogressive failure of MTDs in the South China Sea, and other continental margins.

## 7. Conclusions

Mass-transport deposits and free gas were documented using two high-quality 3D seismic volumes. The main conclusions of this work can be summarised as follows:

- a) The MTDs are characterised by chaotic reflections, scarps and low coherence. In turn, free gas is revealed in the form of enhanced (bright) negative seismic reflections, or by stacked features of reversed polarity and cross-cutting strata;
- b) The free gas is directly trapped under the interpreted MTDs, which suggests they act as competent seals to hinder the vertical migration of free gas. However, the trapped free gas migrates updip beyond the scarps (and basal shear zones) of MTDs;
- c) The trapped free gas migrates vertically from deeper sources via gas chimneys, faults and porous strata.
- d) The seal competence of MTDs depends on petrophysical properties such as density, water content and porosity. These properties drive the MTDs to achieve a more consolidated state than adjacent (undeformed) slope strata;
- e) Overpressure caused by trapped free gas under the MTDs is a likely geohazard that must be taken into account when drilling through MTDs;
- f) The lateral migration of free gas is likely to generate new mass wasting on continental slopes. It potentially triggers multi-stage slope failure and can be responsible for the retrogressive evolution of MTDs on continental margins.

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## Figure Captions

Figure 1: (a) Geological setting and location of sedimentary basins in the northern South China Sea. Inset (top left): regional setting. The two study areas (highlighted by the red squares) are located within the Pearl River Mouth Basin (modified from Sun et al., 2012). ODP sites 1144–1148 are labeled; (b) Seabed morphology of Zone A. The MTD scarp is identified; (c) Seabed morphology of Zone B. Pockmarks (P) are observed and numbered.

Figure 2: Schematic stratigraphic column of the Pearl River Mouth Basin (modified from Pang et al. (2008) and Sun et al.

(2014)). The short red lines in ‘Tectonic Events’ represent the main stages of each tectonic episode.

Figure 3: Composite section of 2D and 3D seismic data crossing ODP Site 1146. MTDa is marked by the light purple polygon. The dashed green and cyan lines highlight the top and base of MTDa. The lines continue into correlative (and undeformed) continental slope strata. Lithology data is from Wang et al. (2000). Well log data is from Sun et al. (2017). DT = acoustic logging; RHOB = density logging. See location in Fig. 1a.

Figure 4: (a)-(d) 3D seismic profiles in Zone A showing the internal character of enhanced negative seismic anomalies (ENSAs) interpreted as free gas in the study area. The ENSAs are trapped underneath MTDa and continue laterally beyond its scarp. The parallel high-amplitude seismic reflections in (b) are interpreted as turbidites. The blue dashed lines in (b) highlight speculative fluid migration pathways (porous strata). The locations of (a)-(d) are labeled in (e); (e) RMS-amplitude map (50 ms (twt) window above and below the basal surface of MTDa) shows the distribution of ENSAs (black dashed line). The boundary (scarps) of MTDa superimposed on this map is shown by the dashed red line. Parallel seismic reflections show higher (RMS) amplitude than free gas.

Figure 5: (a) 3D seismic profile imaging free gas and MTDs in Zone B. (b) Corresponding interpretation of the seismic profile in (a). Six MTDs (MTDb1-MTDb6) showing chaotic seismic reflections and free gas form vertically stacked enhanced seismic anomalies (SESAs) underneath the MTDs. The labelled gas chimneys (blue arrows) and faults (dashed red lines) serve as major fluid pathways. See location of seismic section in (c); (c) RMS-amplitude map of Zone B (80 ms (twt) window below the basal surface of lowermost MTD) highlighting the sub-circular distribution of free gas.

Figure 6: (a) 3D seismic profile highlighting the relationship between free gas and MTDs in Zone B. (b) Corresponding interpretation of the seismic profile in (a). Gas chimneys (marked with blue arrows) are shown as vertical chaotic seismic reflections. Faults (marked with dashed red lines) are also identified. The stacked enhanced seismic anomalies (SESAs) with reversed seismic polarity (c) and cross-cutting seismic reflections (d) are interpreted as free gas.

Figure 7: Isochron thickness map of gas-bearing strata (stacked enhanced seismic anomalies) in Zone B. Several depocenters are identified and correlate with fluid migration pathways (gas chimneys and faults).

Figure 8: (a) Isochron thickness map for MTDa, Zone A. The thickness of MTDa is variable and is reduced from the scarp (northeast) downslope (southwest). (b)-(c) Internal seismic character (chaotic seismic reflections) of MTDa in Zone A. The dashed blue lines are speculative fluid migration pathways (porous strata) interpreted in the study area. ENSAs (enhanced negative-phased seismic anomalies, free gas), blocks, internal ridges and scarps are labelled in the figure.

Figure 9: Variance slices across MTDa, Zone A. (a) Variance slice of the flattened top of MTDa. (b) Variance slice 30 ms below the flattened top of MTDa. (c) Variance slice 30 ms above the flattened base of MTDa. (d) Variance slice of the flattened base of MTDa. Blocks, linear structures and scarps are identified in the figures.

Figure 10: Variance slice highlighting the internal character of MTDs in Zone B. (a) Variance slice of the flattened base of MTDb1. (b) Variance slice of the flattened base of MTDb2. (c) Variance slice of the flattened base of MTDb6. (d) Variance slice at a two-way time depth of 2850 ms, highlighting the distribution of free gas (dashed pink line) and MTDb5 and MTDb6 at this depth. Blocks and grooves are identified and indicates that MTDs were transported from NW to SE.

Figure 11: Isochron thickness map of MTDb1, Zone B. MTDb1 is shown to pinch out towards the southern part of the

study area.

Figure 12: Relative distribution of MTDb1 to MTDb6 and free gas in Zone B. The main volume of free gas is overlain by MTDb5 and MTDb6.

Figure 13: Well log data from (a) IODP Site 1323 and (b) ODP Site 936A highlighting the petrophysical character of MTDs at these locations after their emplacement. The water content and porosity of MTDs decrease, whereas the density, compressional velocity and resistivity increase after their emplacement. This indicates that the MTDs are more consolidated strata above and below, and form competent seal intervals.

Figure 14: Conceptual models for fluid migration and accumulation underneath MTDs in Zone A (a) and Zone (b). In Zone A, vertical fluid migration chiefly occurs by diffusion along porous strata. After the fluid reaches the base of MTDa is trapped by overconsolidated strata in this deposit, and migrates laterally along its base. In Zone B, vertical fluid migration chiefly occurs via gas chimneys and pipes. Fluid is trapped by MTDb5 and MTDb6 and accumulates within porous strata below the two MTDs, particularly in their basal shear zones.

**Table caption**

Table 1: Geometric characteristics of MTDs and free gas in Zones A and B. AT = average thickness; MTD = Mass-transport deposit. Please note that the average thickness and volume of free gas reflect the thickness and volume of gas-bearing strata in the two studied zones. \* means a conservative value.

Figure 1

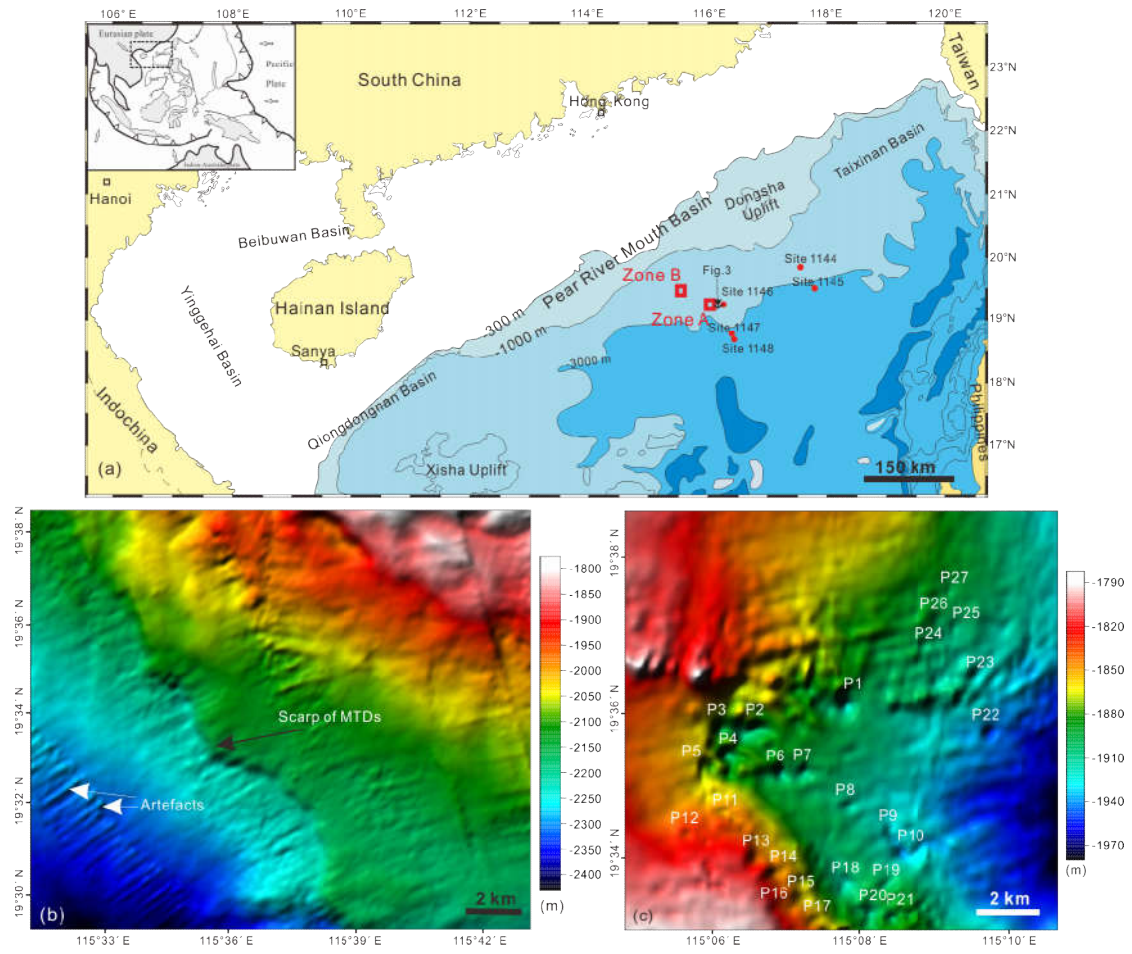


Figure 2

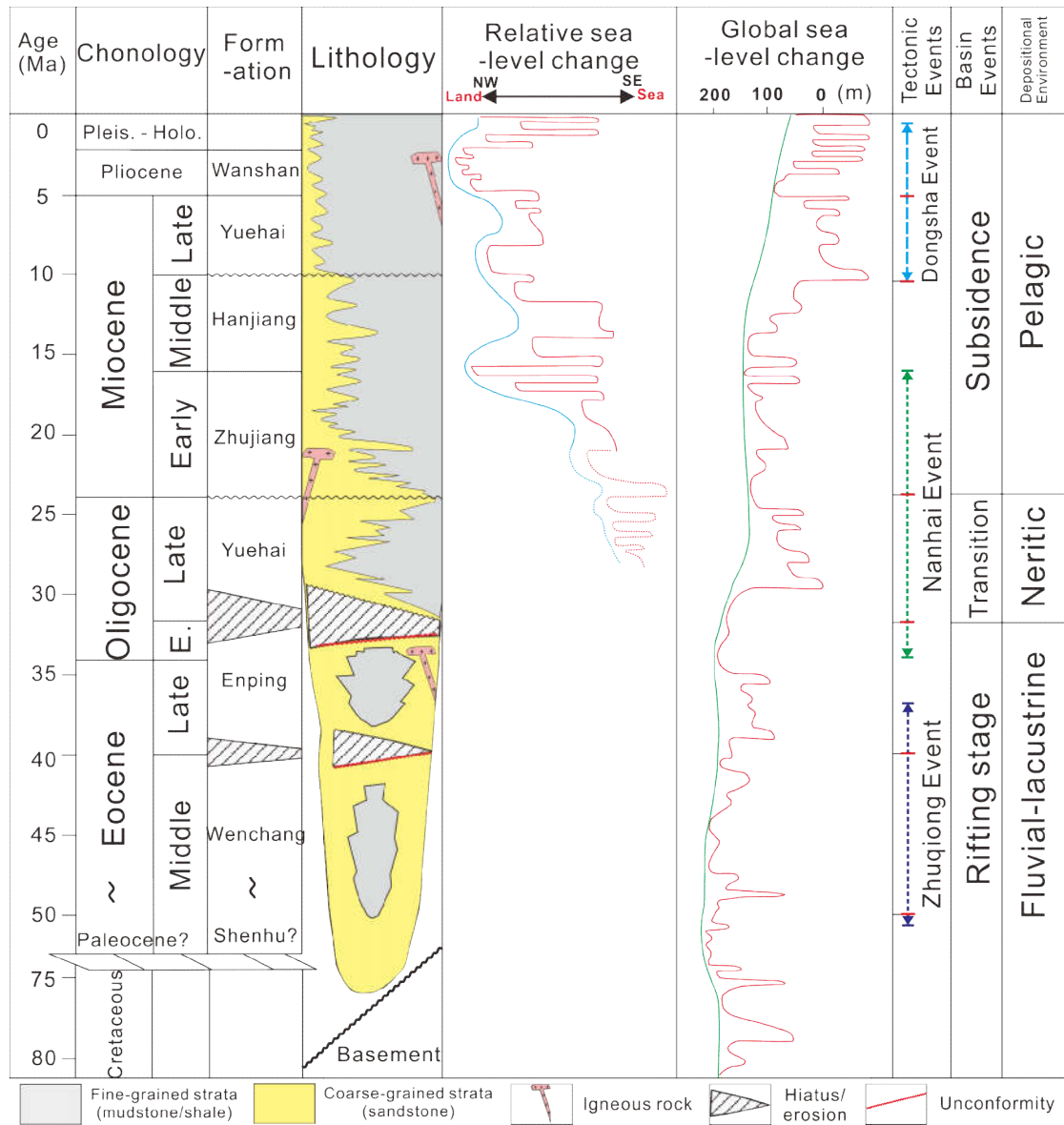




Figure 3

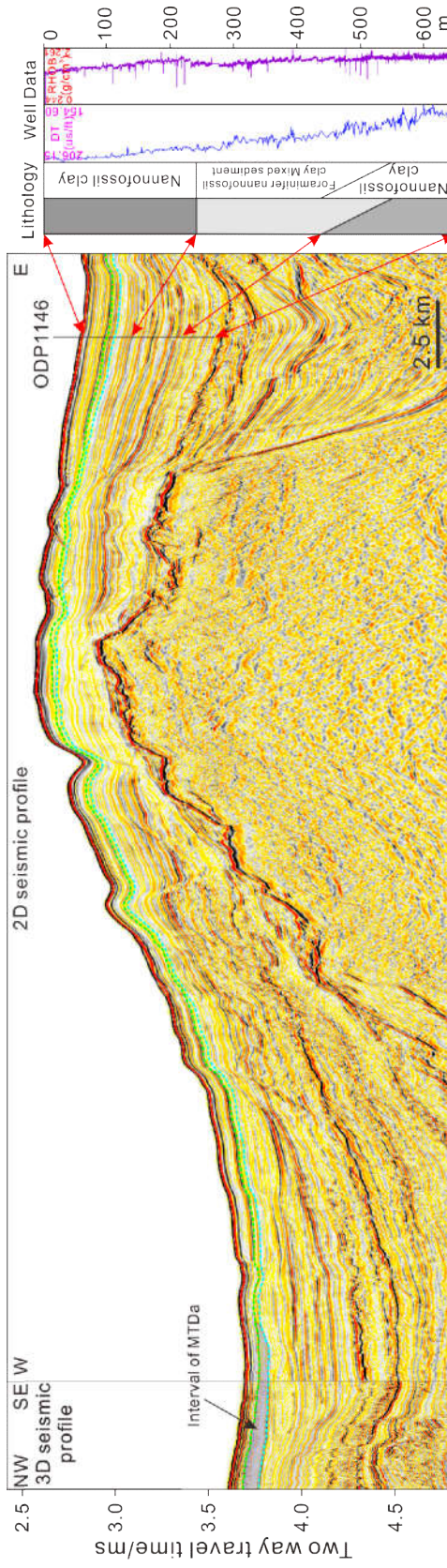


Figure 4

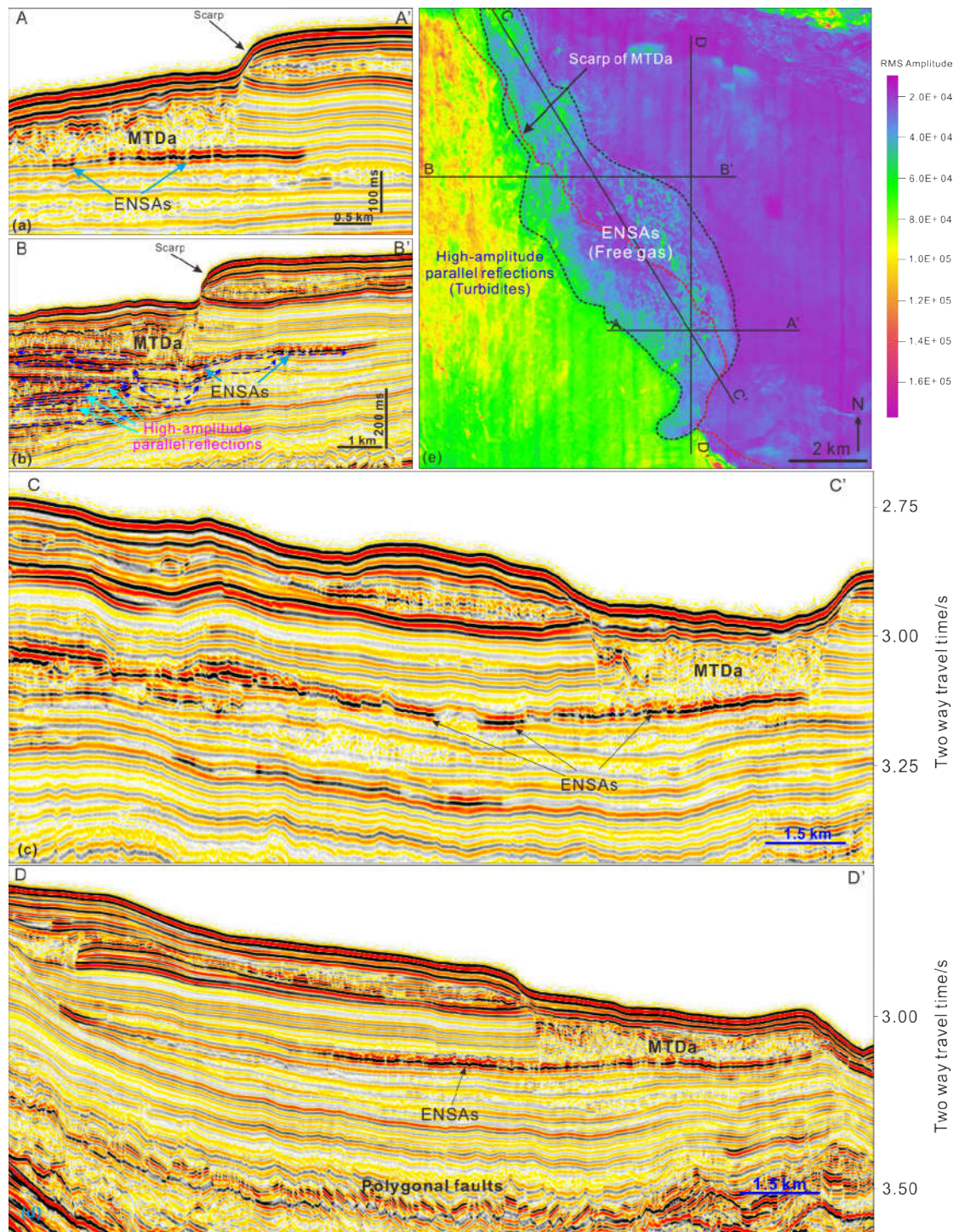




Figure 5

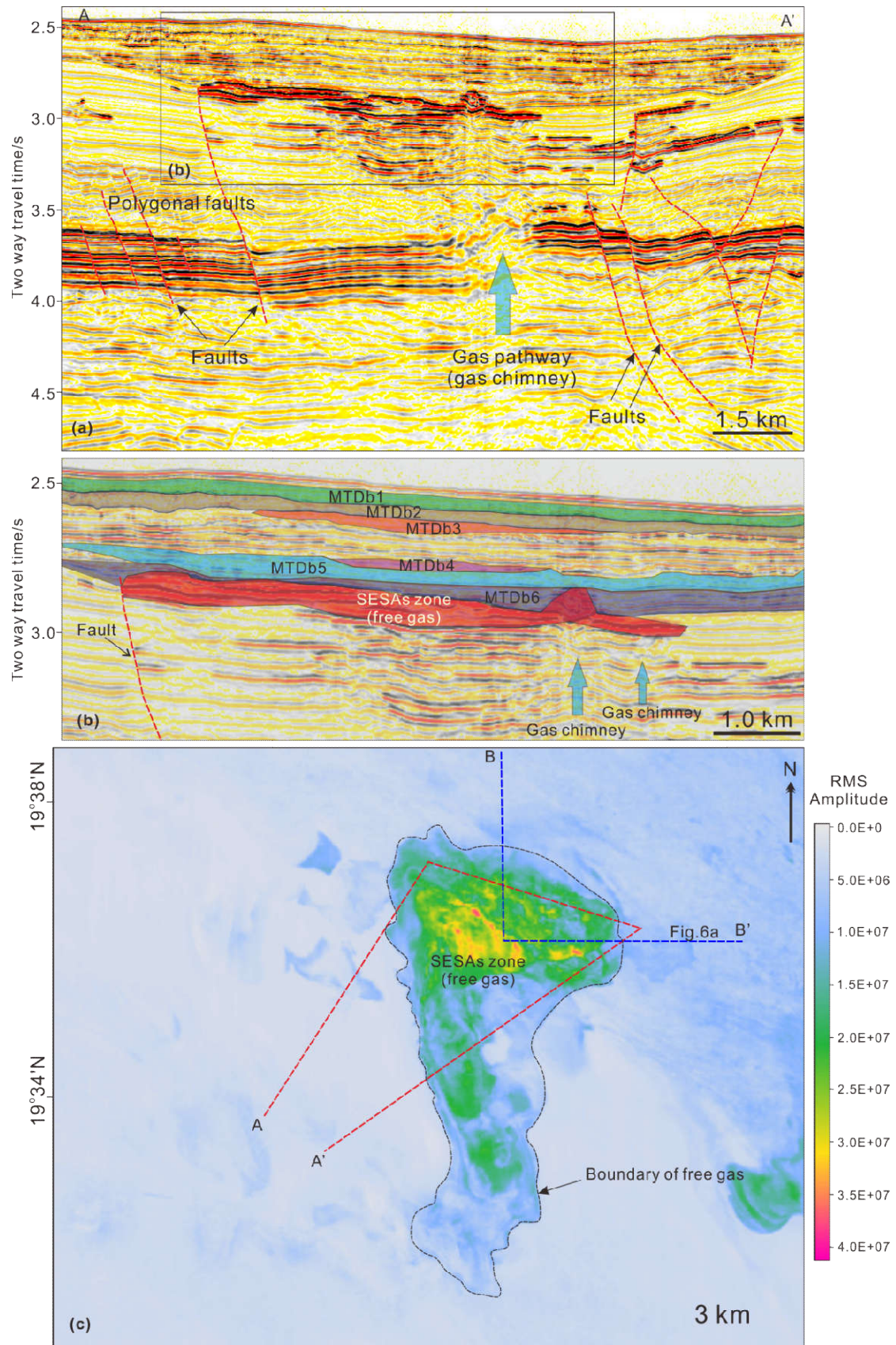




Figure 6

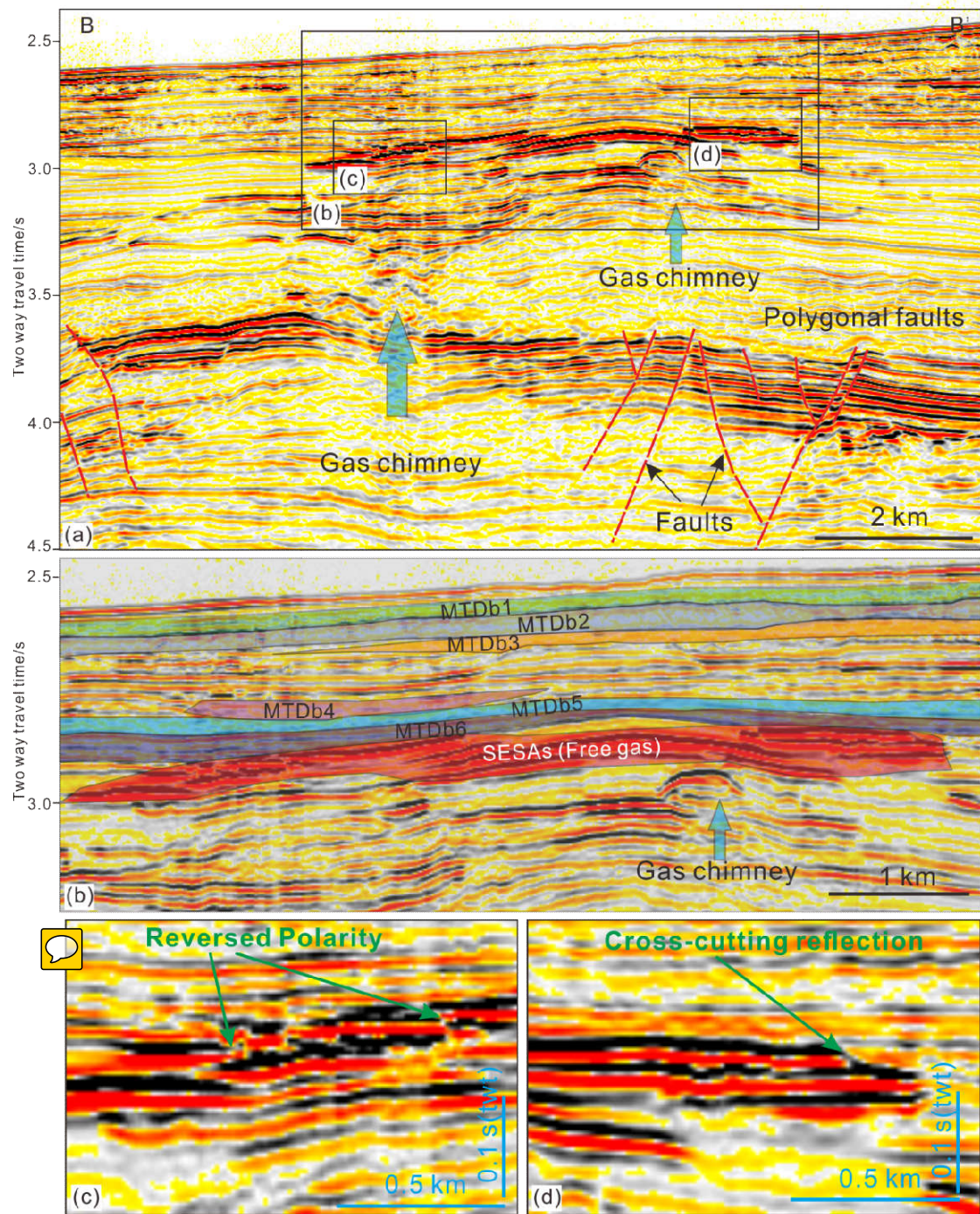


Figure 7

Two way isochron thickness of gas-charged strata in Zone B

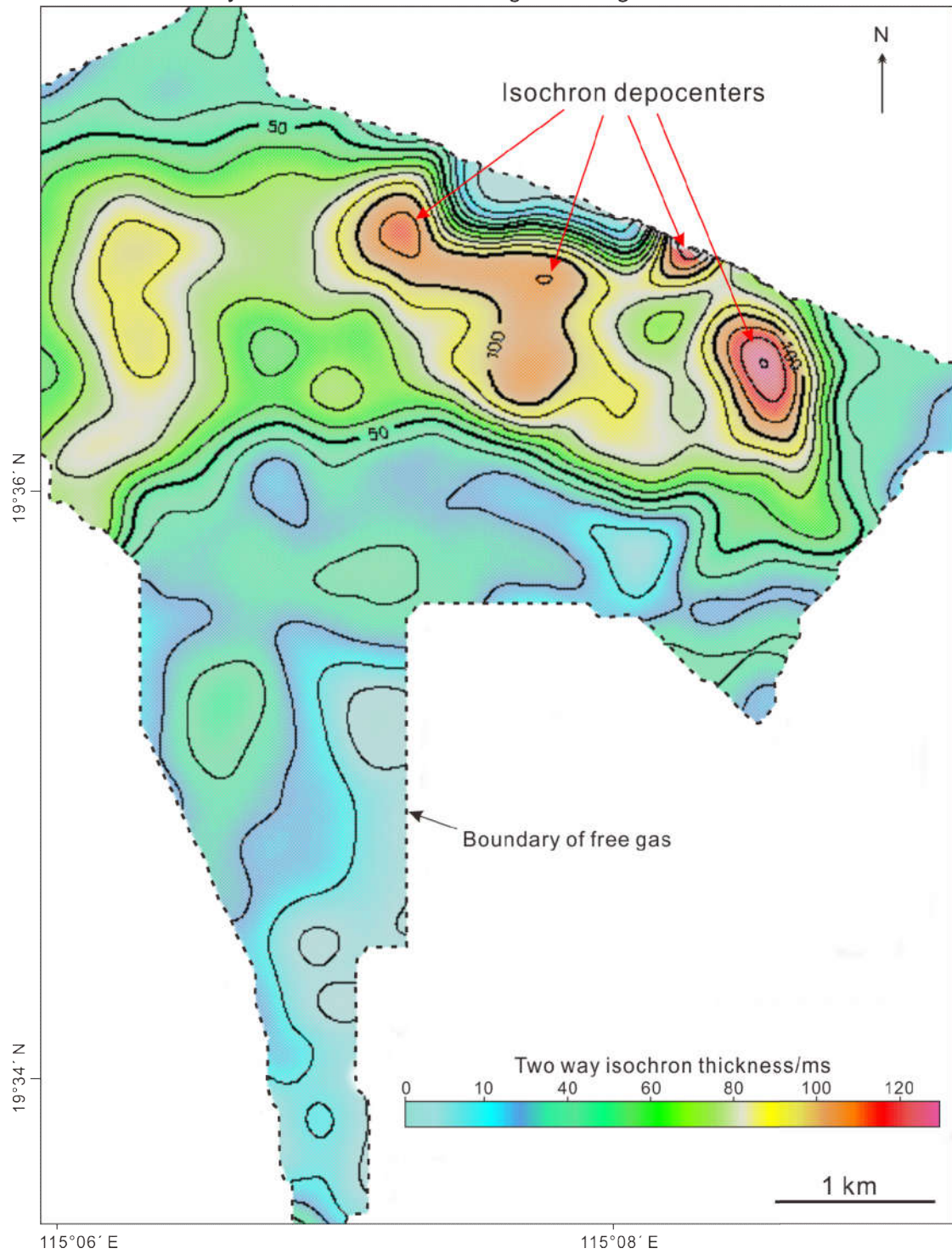




Figure 8

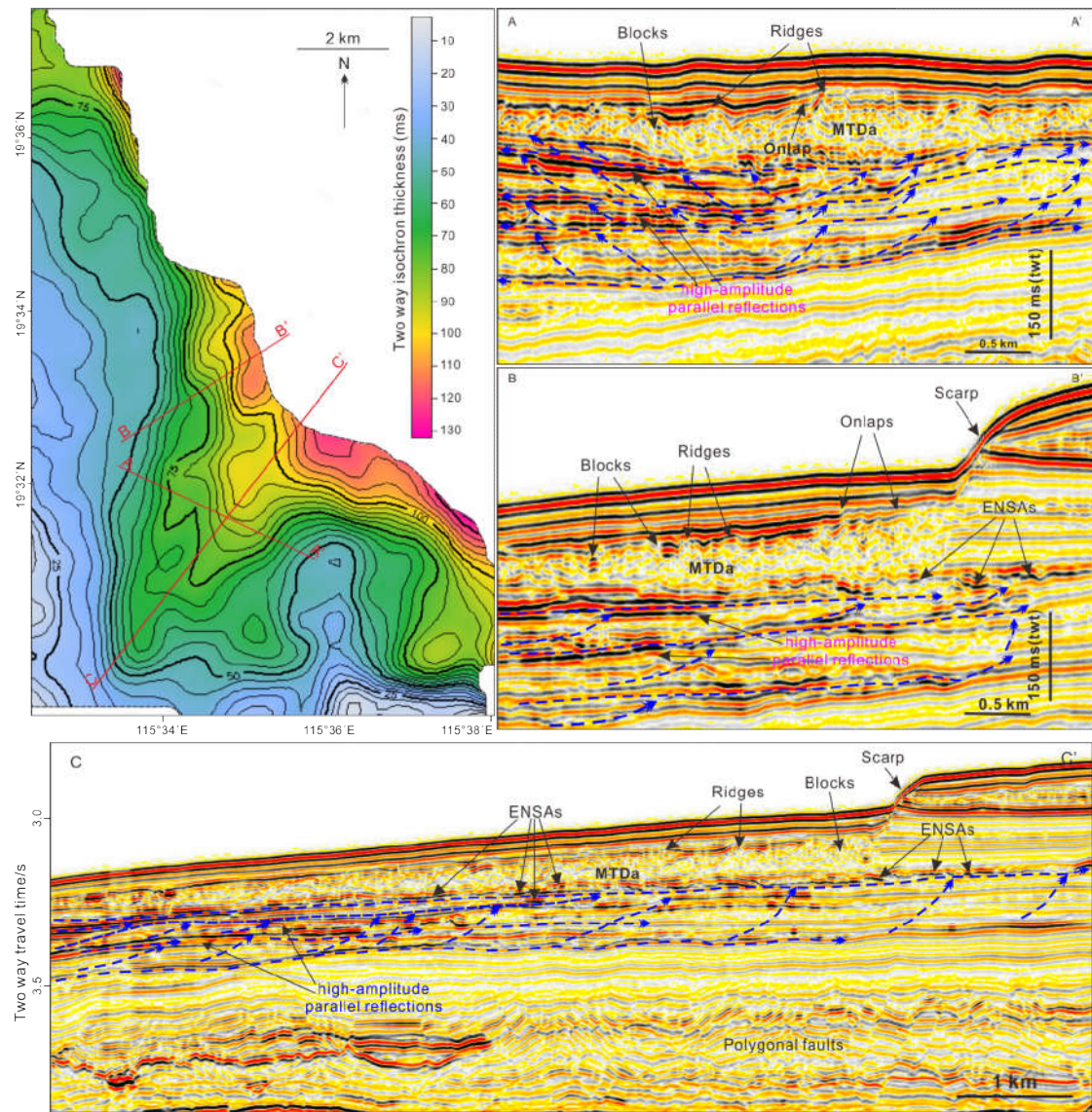


Figure 9

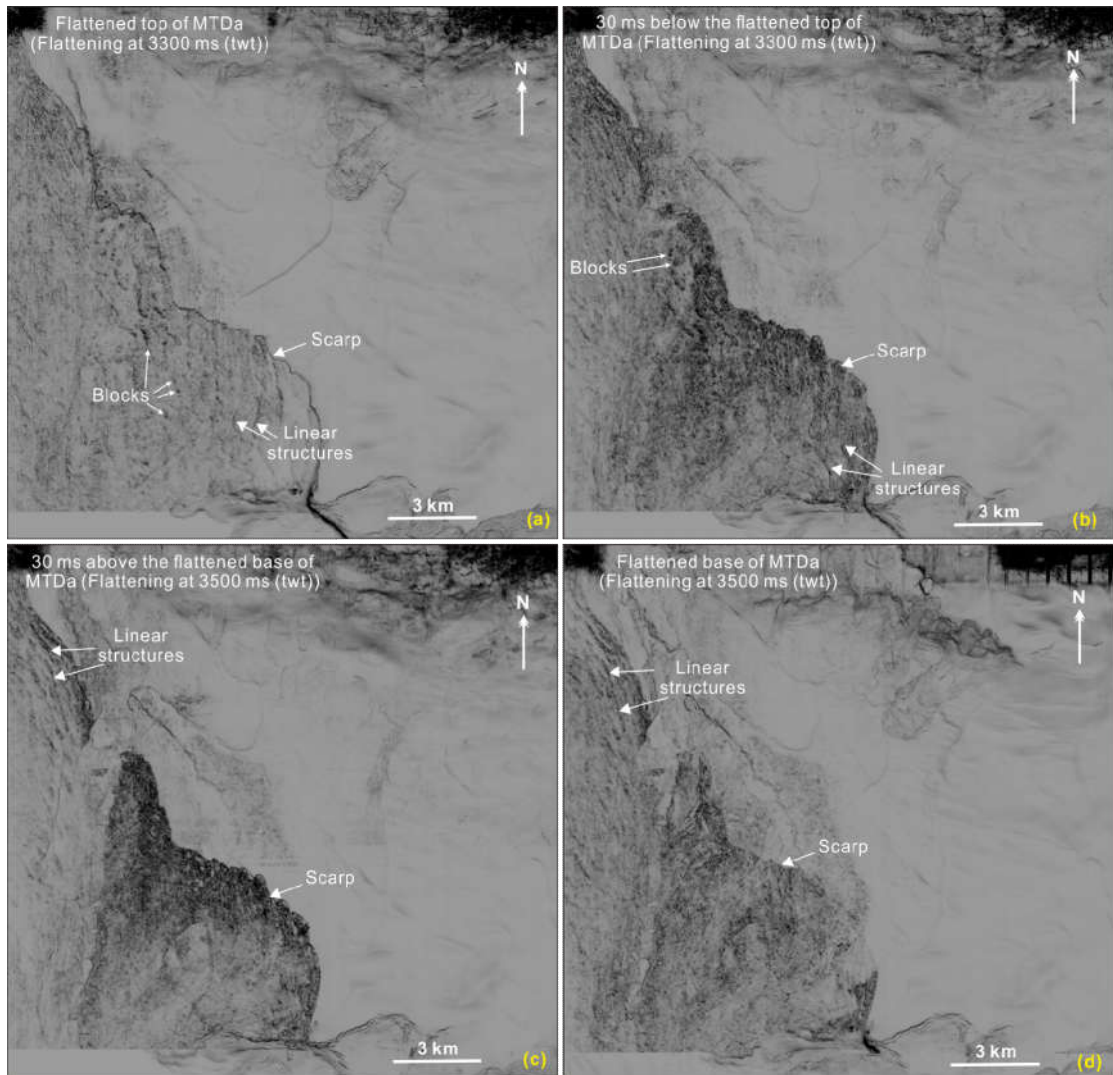




Figure 10

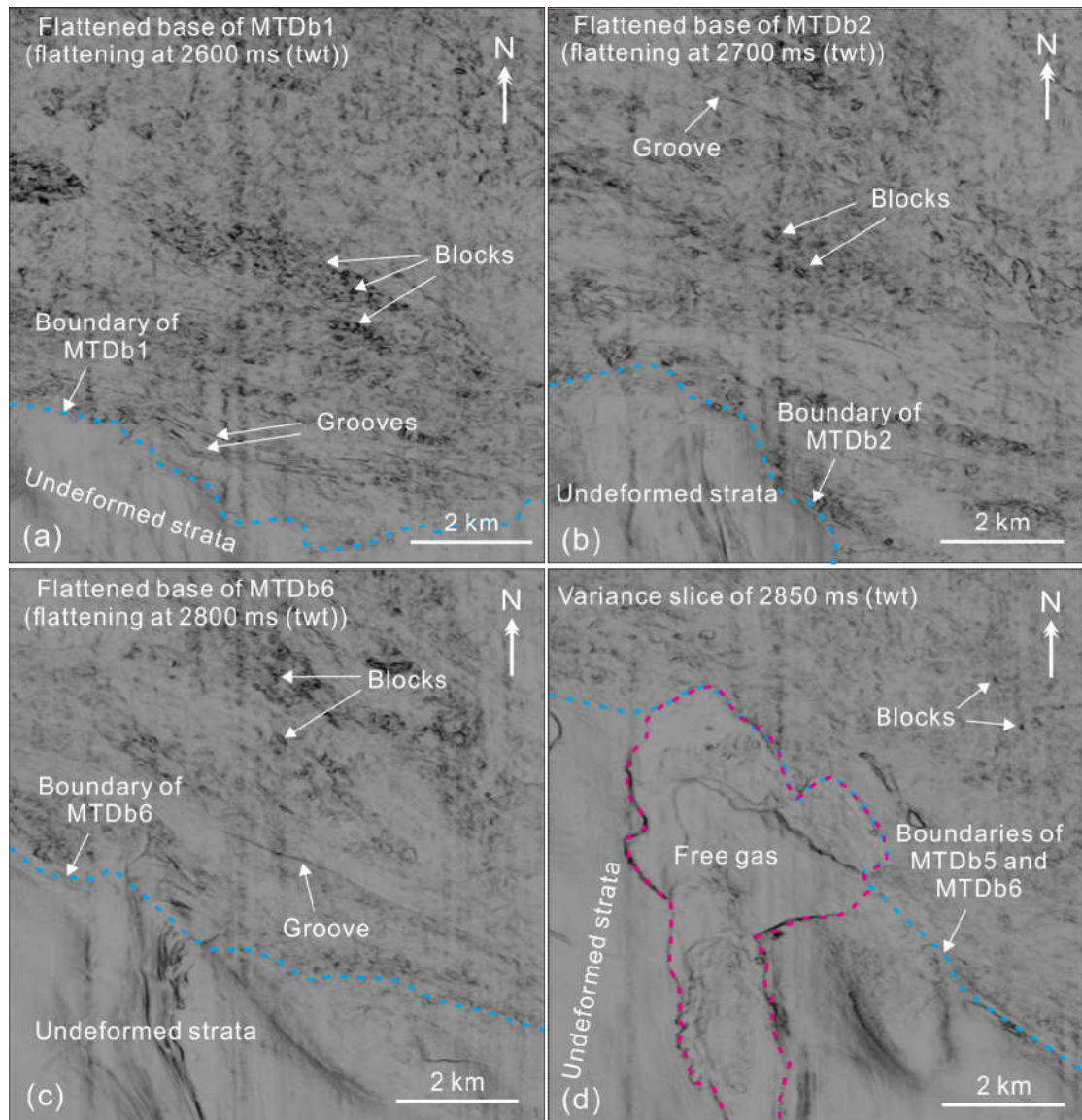


Figure 11

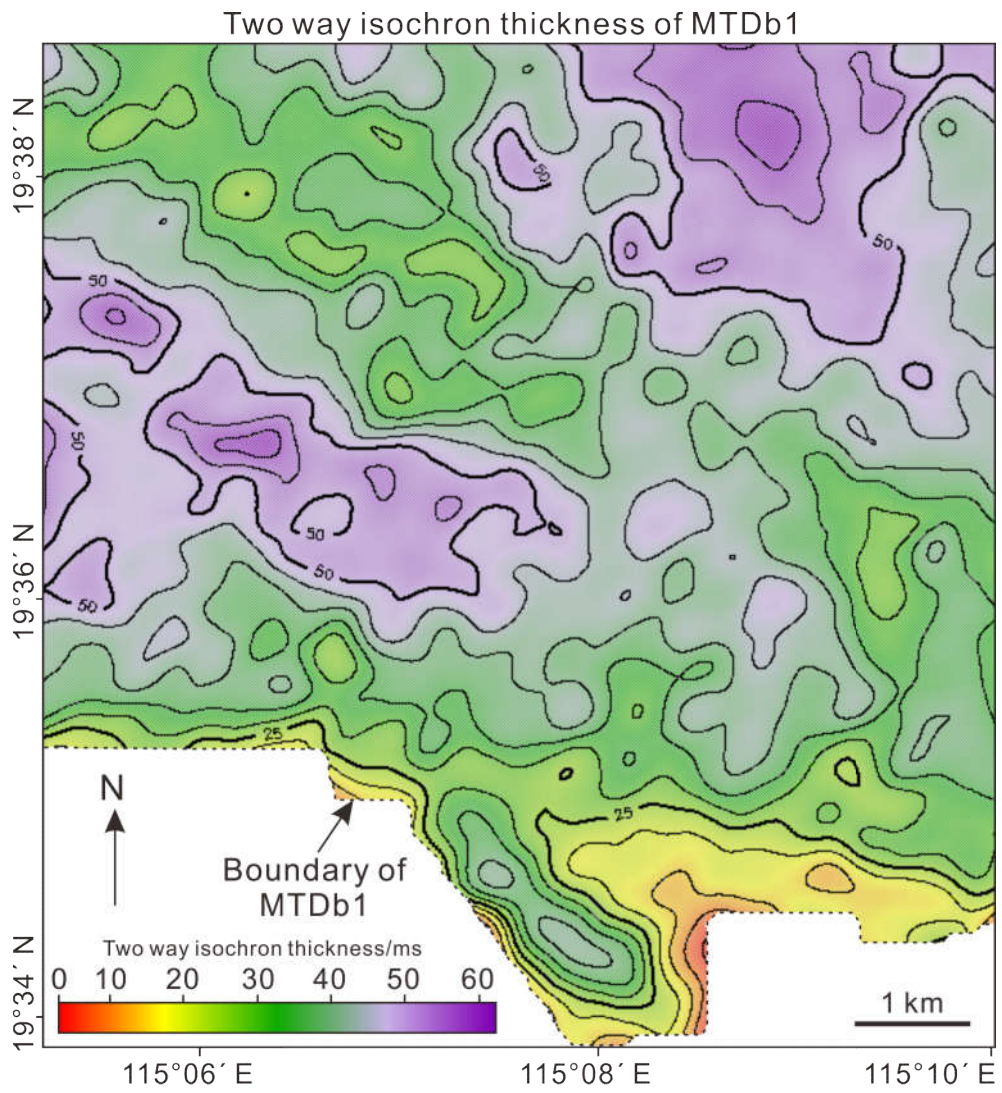


Figure 12

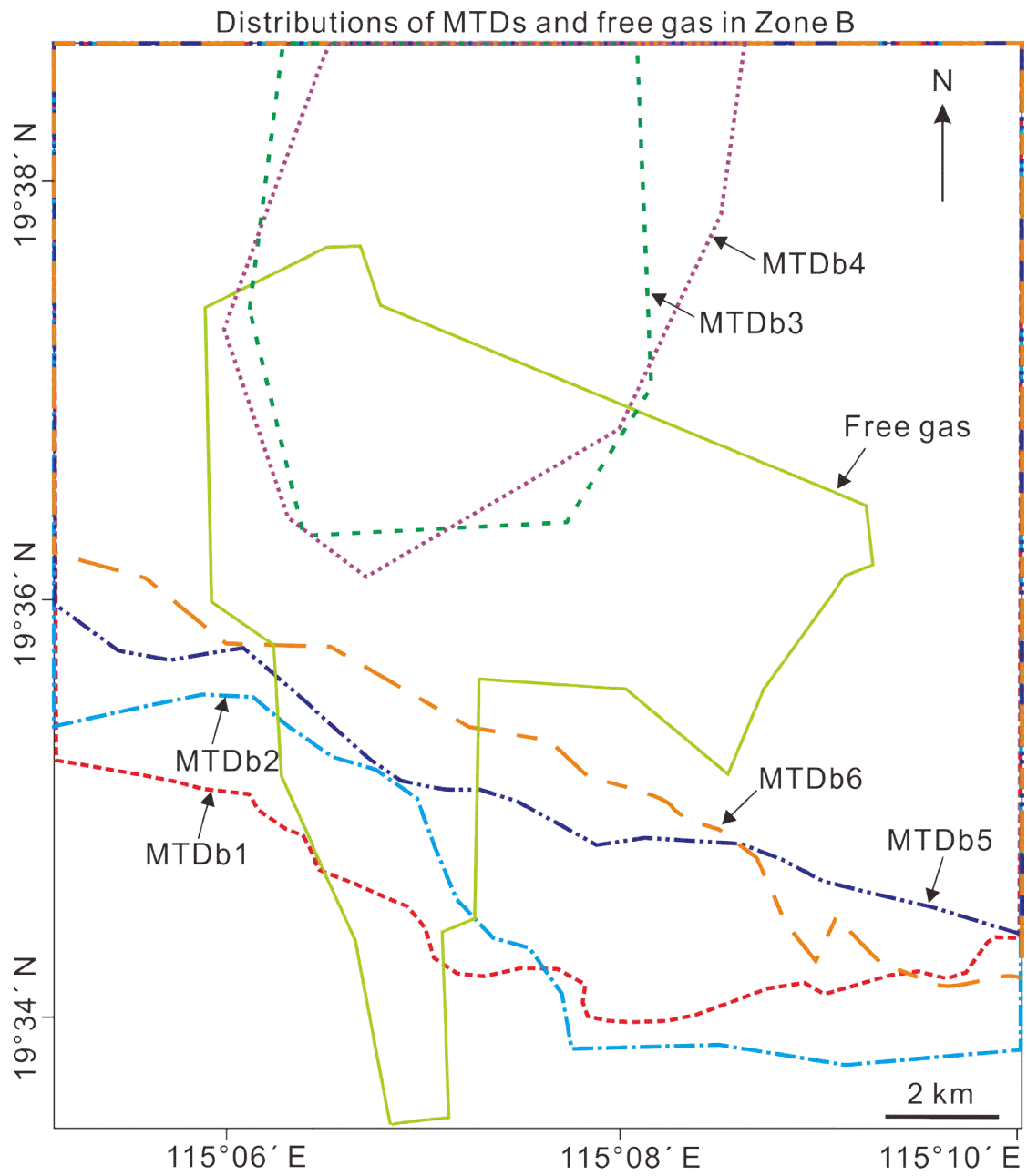


Figure 13

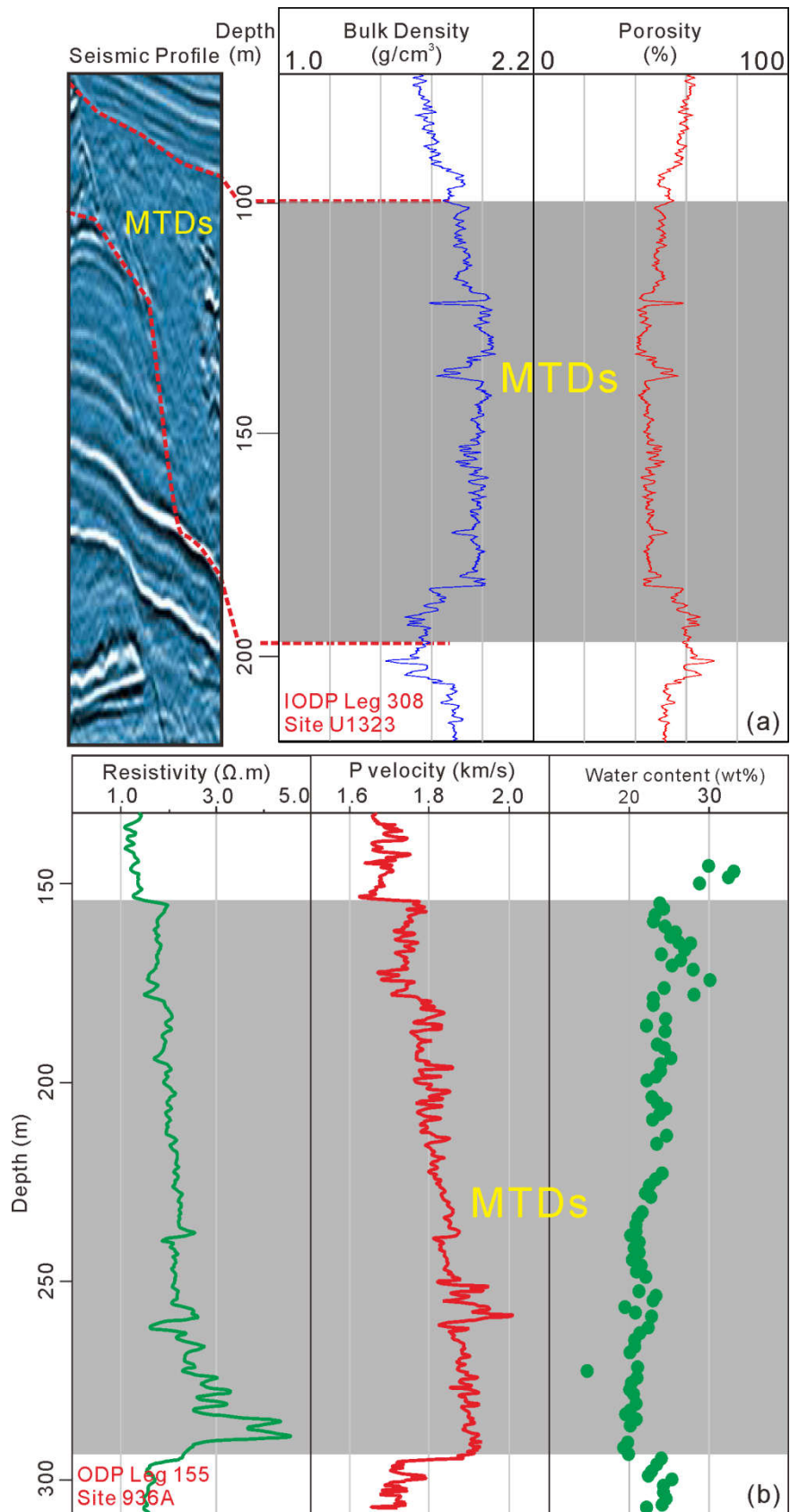




Figure 14

