

This is an Open Access document downloaded from ORCA, Cardiff University's institutional repository: <https://orca.cardiff.ac.uk/id/eprint/102148/>

This is the author's version of a work that was submitted to / accepted for publication.

Citation for final published version:

Cotton, Laura J., Janssen, Arie W., Pearson, Paul N. and Van Driel, Rens 2017. Pteropoda (Mollusca, Gastropoda, Thecosomata) from the Eocene/Oligocene boundary interval of three cored boreholes in southern coastal Tanzania and their response to the global cooling event. *Palaeontologia Electronica* 20 (3) , pp. 1-21.

Publishers page: <http://palaeo-electronica.org/content/pdfs/733.pdf>

Please note:

Changes made as a result of publishing processes such as copy-editing, formatting and page numbers may not be reflected in this version. For the definitive version of this publication, please refer to the published source. You are advised to consult the publisher's version if you wish to cite this paper.

This version is being made available in accordance with publisher policies. See <http://orca.cf.ac.uk/policies.html> for usage policies. Copyright and moral rights for publications made available in ORCA are retained by the copyright holders.



This is an incomplete preprint of a paper in press with Palaeontologia Electronica and should not be used or cited. It will be replaced by the published open access version in due course

Pteropoda (Mollusca, Gastropoda, Thecosomata) from the Eocene/Oligocene boundary interval of three cored boreholes in southern coastal Tanzania and their response to the global cooling event

Key words: planktic Gastropoda, *Altaspiratella*, *Heliconoides*, *Limacina*, *Bovicornu*, new species, geographical distribution, vertical ranges, Eocene-Oligocene transition (EOT)

24

25

26

27

28

29

30 Laura J. Cotton

31 Naturalis Biodiversity Center, P.O. Box 9517, 2300 RA Leiden, The Netherlands; present

32 address: Florida Museum of Natural History, 1659 Museum Road, University of Florida,

33 Gainesville, FL 32611, U.S.A and Department of Geological Sciences, University of

34 Florida, 241 Williamson Hall, Gainesville, FL32611, U.S.A .; lcotton@flmnh.ufl.edu

35

36 Arie W. Janssen (corresponding author)

37 Naturalis Biodiversity Center, P.O. Box 9517, 2300 RA Leiden, The Netherlands;

38 ariewjanssen@gmail.com

39

40 Paul N. Pearson

41 Department of Earth and Ocean Sciences, Cardiff University, Main Building, Park Place,

42 Cardiff, CF10 3AT, United Kingdom; PearsonP@cardiff.ac.uk

43

44 Rens van Driel

45 Buys Ballotsingel 85A, 3112 JD, Schiedam, The Netherlands; rensvdriel@gmail.com

46

47
48 \

49
50
51
52
53 ABSTRACT

54 The Eocene - Oligocene Transition was a period of major climatic and oceanographic
55 change, resulting in widespread biotic overturning. However, the record of many marine
56 organisms remains patchy. Planktic Mollusca (Pteropoda) from three cored boreholes
57 spanning the Eocene/Oligocene boundary (EOB) in southern coastal Tanzania are
58 represented by eight species, three of which are introduced as new: *Heliconoides nikkieae*
59 sp. nov., *Limacina tanzaniaensis* sp. nov. and *Limacina timi* sp. nov. Three of the other
60 species can only be identified in open nomenclature. The two most commonly occurring
61 species, *H. nikkieae* and *L. timi*, straddle the EOB without noticeable loss in abundance.
62 Two species, *Limacina robusta* (Eames, 1952) and *L. tanzaniaensis* disappear before the
63 EOB. The species *Bovicornu* aff. *eocenense* Meyer, 1886 disappears shortly after the
64 EOB. Two species were only found in a single sample each, in the Eocene part of the
65 succession. Response to changing environmental conditions seems to be demonstrated by
66 two or three of the pteropod species only that become extinct before or shortly after the
67 EOB.

70
71
72
73
74
75
76
77
78
79
80
81
82
83
84
85
86
87
88
89
90
91
92

Plain Language Summary

A major cooling event took place between 33.5 and 34 Ma, known as the Eocene-Oligocene transition (EOT). The response of many marine micro-biota, such as foraminifera, has been well documented, however, records of other marine organisms, such as pteropods, are less well-known. Pteropods are planktic gastropods which spend their life in the water column, and are sensitive to changes in the ocean environment. Here we describe the assemblage and ranges of well-preserved pteropods through the EOT from three borehole records from Tanzania, an exceptionally understudied region. In total eight species were found in this material, three of which are new. The two most common species pass through the climatic event with no noticeable change in abundance. Three other species disappear during the transition, at a level similar to the planktic extinctions. The remaining species are too few to comment on. This demonstrates a likely response to the changing conditions, similar to some foraminiferal groups. Despite the small sample sizes, this study represents an important new record from an understudied region.

INTRODUCTION

The Eocene-Oligocene transition (EOT) was one of the most dramatic events of the Cenozoic, associated with major climatic disruption and widespread biotic turnover in both marine and terrestrial realms (see Coxall and Pearson, 2007 for review). High-resolution studies of the marine response have been carried out using planktic and benthic foraminifera, and nanno-fossils (e.g. Zachos et al., 1994; Molina et al., 1993; Diester-Haass et al., 2001; Lui et al., 2004; Wade and Pearson, 2008; Cotton and Pearson, 2011; Cotton et al., 2014). However, records of other marine groups are often patchy. Pteropods are planktic gastropods and are increasingly being used in studies of recent ocean records and climate change (Fabry, 1990; Wall-Palmer et al., 2012; Bednaršek and Ohman, 2015; references therein). Their aragonitic shells are susceptible to dissolution, making them excellent model organisms to study ocean acidification. However, palaeoecological data for older pteropod records are relatively uncommon, yet may add important information to the overall understanding of climatic impacts in the marine realm. Pteropods have a well-documented fossil record, but many regions remain understudied, and early well-preserved records are therefore exceptionally important (e.g., Janssen *et al.*, 2016).

The Eocene-Oligocene Transition (EOT) is an extended period of global oceanographic and climatic perturbation spanning between 33.5 - 34 Ma. The Eocene/Oligocene Boundary (EOB) is defined by the last occurrence of the planktic foraminiferal family Hantkeninidae at the Global Boundary Stratotype-Section and Point (GSSP) at Massignano, near Ancona, on the Italian Adriatic coast (Premoli Silva and Jenkins, 1993) which occurs at 33.7 Ma on the timescale of Cande and Kent (1995). However, an unreliable magnetostratigraphy and stable-isotope stratigraphy (Bodiselišch *et al.*, 2004)

116 and lack of carbonate (macro)fossils (pers. observ. AWJ, August 1992) in the Massignano
117 section prevented global correlation with climatic events. The additional fact that the
118 occurrence of Hantkeninidae is mostly restricted to tropical and sub-tropical
119 environments also made long-distance correlations (*e.g.*, to the North Sea Basin and in
120 particular to the Priabonian stratotype) problematic. Subsequent studies of Deep Sea
121 Drilling Project Site 522, showed that the extinction preceded the most positive oxygen
122 isotopes of the EOT, representing the Early Oligocene Glacial Maximum (Zachos *et al.*,
123 1994; Liu *et al.*, 2004). More recently this has been further refined.

124 The Kilwa District in southern Tanzania contains an apparently complete succession
125 through the EOT, which was recovered by three Tanzanian Drilling Project boreholes
126 (TDP 11, 12 and 17; Nicholas *et al.*, 2006; Pearson *et al.*, 2008). The dominant lithology
127 consists of hemipelagic clays interspersed with debris flow limestones (Nicholas *et al.*
128 2006). These clays contain exceptionally well preserved calcareous micro- and
129 nannofossils, including aragonitic preservation (Pearson *et al.*, 2008; Wade and Pearson,
130 2008; Bown *et al.*, 2008). This, coupled with the expanded nature of the sediments
131 allowed for high resolution stratigraphy chemo-and biostratigraphy to be carried out and
132 precise correlation to global records. This placed, for the first time, the extinction of the
133 Hantkeninidae between the two positive isotope shifts of the transition (Pearson *et al.*,
134 2008; Wade and Pearson, 2008); the first of which is attributed to temperature change and
135 the second largely to ice-volume (Lear *et al.*, 2008). In addition, the site contains many
136 other exceptionally well-preserved calcareous micro-fossils, including larger benthic
137 foraminifera (Cotton and Pearson, 2011), bryozoans (Martino *et al.*, 2017),
138 dasycladaceae, ostracodes and molluscs, including pteropods (Cotton *et al.*, in prep),

making Tanzania an exceptionally important site for the study of comparatively less ubiquitous fossils across the EOT. Studies of the molluscan record have shown that overall numbers of mollusc shells increase across the boundary (Cotton *et al.*, in prep), however, this study mainly examined the benthic forms. Here we present a detailed record of occurrences through the EOT interval and taxonomy of specifically the pteropod fauna from TDP 11, 12 and 17 (see Nicholas *et al.* 2006; Cotton and Pearson, 2011, for details on these boreholes). Though relatively few specimens were found due to the small sample size, the often excellent preservation and well-correlated nature of the record offers a unique insight into a region where no pteropod records were previously known.

Commented [LC1]: Not sure what these refs relate to
AWJ: tried to make that a bit more clear

MATERIAL AND METHODS

The Tanzania Drilling Project (TDP) was initiated in 2002 (Pearson *et al.*, 2004; Nicholas *et al.*, 2006, 2007) after preliminary field observations beginning in 1998 (Pearson *et al.*, 2001) on numerous outcrops between the towns of Kilwa and Lindi, southern Tanzania, had demonstrated the presence of well-preserved microfossil assemblages of Cretaceous to Paleogene age. The TDP focused on generating litho- and biostratigraphic, geochemical, micropalaeontological and palaeoclimatic records, and resulted in recovery of over 40 cored boreholes, covering Cretaceous (Aptian) to Paleogene (Oligocene) sediments (Pearson *et al.*, 2004, 2006; Nicholas *et al.*, 2006; Jimenez-Berrucoso *et al.*, 2010, 2012, 2015). These sediments are formally defined as the Kilwa Group and were initially subdivided into four formations: The Nangurukuru, Kivinje, Masoko and Pande formations (Nicholas *et al.*, 2006), to which a fifth was recently added, the Lindi formation (Jimenez-Berrucoso *et al.*, 2015). The Pande

162 Formation contains an apparently conformable succession through the EOT and was
163 recovered in three boreholes drilled approximately along strike of each other (Figure 1):
164

165 TDP 11: South of Stakishari (Tanzania, Kilwa region), coordinates UTM 37L 560250
166 8983211; pteropod specimens are available from four samples TDP11.26.2, 64-74 cm, to
167 TDP11.33.2, 74-84 cm.
168

169 TDP 12: South of Stakishari (Tanzania, Kilwa region), coordinates UTM 37L 560222
170 8981309; pteropods are available from 16 samples TDP12.11.4, 20-26 cm, to TDP12.29.2,
171 20-30 cm.
172

173 TDP 17: Stakishari (Tanzania, Kilwa region), coordinates UTM 37L 560539 8984483;
174 pteropod specimens available from 36 samples 17.15.1, 50-63 cm, to 17.41.3, 0-15 cm.
175

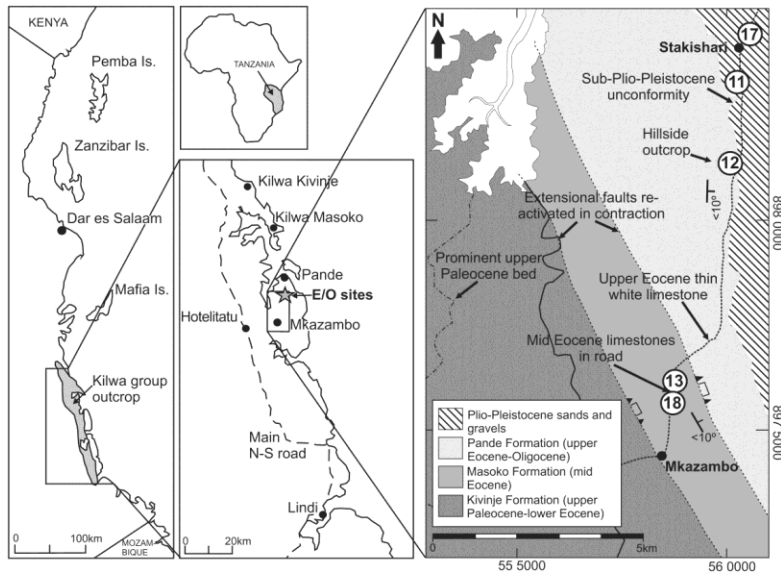


Figure 1. Location and geological map of the Tanzanian Drilling Project

Eocene/Oligocene boundary sites (TDP 11, 12 and 17), additional Tanzanian Drilling Project sites in the area are also shown, modified from Nicholas *et al.* (2006). After Cotton and Pearson, (2011).

Based on micropalaeontological event correlations the three sections were given “composite depths” to enable direct comparison of results between the three sites (Pearson *et al.*, 2008). These composite depths are specified in the distribution tables following the systematical part..

The meters added are, for TDP 17 cores 1-35: 0 m; for TDP17 cores 36-42: 12 m; for TDP11 cores 1-28: 17 m, and for TDP11 cores 29-36: 20 m.

The Eocene Tanzanian shelf is thought to have been narrow, much like today, and the three boreholes are estimated to be approximately 50 km from the palaeo-shoreline (Kent *et al.*,

1971; Nicholas *et al.*, 2006, 2007). Based on the sedimentary facies and smaller benthic foraminiferal assemblage the palaeo-depth is estimated to be approximately 300-500 m, although it is possible it is deeper than this (Nicholas *et al.*, 2006; 2007). The dominant sediment is dark green-grey clay with < 10% CaCO₃. The half-round core-samples of the hemi-pelagic clays, varying in length from 1.5 to 15 cm, were washed through a 63 µm sieve and the residues dried (Lear *et al.*, 2008; Pearson *et al.*, 2008; Wade and Pearson, 2008). Pteropods were then picked using a paintbrush under a binocular microscope. The Kilwa Group is characterised by excellent preservation of its carbonate microfossils, this extends to the pteropods, which is related to the high clay content and shallow maximum burial depth (Pearson *et al.*, 2001; Van Dongen *et al.*, 2006; Bown *et al.*, 2008). However, the condition of pteropod specimens varies from perfect shell preservation to shells filled with pyrite and specimens preserved as pyritic internal moulds with partly or completely dissolved aragonitic shell material. Local dissolution may be related to pyritisation. Some specimens are badly deteriorated and indeterminate. Specimens are housed in the fossil holoplanktic mollusc collection of Naturalis Biodiversity Center, Leiden (The Netherlands), registered with RGM-registration numbers.

Scanning electron microscopy

SEM micrographs were made by Renate Helwerda from uncoated specimens at 1.000 kv with a Jeol Field emission scanning electron microscope type: JSM-7600F at Naturalis Biodiversity Center, Leiden, The Netherlands.

213

214 **Micro - Computed Tomography**

215 Micro-computed tomography (Micro-CT or X-ray microtomography) scanning of the
216 three holotypes (RGM 777415a, RGM 777428b and RGM 1007748b) was carried out by
217 Dirk van der Marel, of Naturalis Biodiversity Center in Leiden, the Netherlands, using a
218 Bruker SkyScan 1172 micro-CT scanner. 1601 projections of 3 -15 sec exposure were
219 collected with a 2000x1336 detector, using a source voltage of between 80 and 140 kV.
220 No filtration was used, and the scan provided a reconstructed dataset with 0.8 – 1.4 μm
221 voxels.

222 The pteropod shells were segmented in Avizo 9.0.0. and surface files exported as .wrl.

223 The files were subsequently edited in Meshlab 1.3.3. Laplacian smoothing was applied
224 and Quadric Edge collapse Decimation used, reducing the polygon count by
225 approximately a factor of 10. Files were saved as .ply and u3d for viewing as three
226 dimensional PDFs.

227

228 **SYSTEMATICS (AWJ)**

229 Phylum **MOLLUSCA** Linnaeus, 1758

230 Class **GASTROPODA** Cuvier, 1795

231 Subclass **HETEROBRANCHIA** Burmeister, 1837

232 Superorder **PTEROPODA** Cuvier, 1804

233 Order **THECOSOMATA** de Blainville, 1824

234 Suborder **EUTHECOSOMATA** Meisenheimer, 1905

235 Superfamily **LIMACINOIDEA** Gray, 1847

Commented [LC2]: Do we need to remove this as suggested by the editors comments?
A WJ: No, that would be quite unusual.

236 Family LIMACINIDAE Gray, 1847

237 Genus ALTASPIRATELLA Korobkov, 1966 (= *Plotophysops* Curry, 1982)

238

239 *Type species.* – “*Limacina elongatoides*” [sic] (Aldrich), by original designation of

240 Korobkov (1966, p. 74) = *Physa elongatoidea* Aldrich, 1887 = *Altaspiratella*

241 *elongatoidea* (Aldrich, 1887) (Eocene, early Ypresian, Wilcox Group, Hatchetigbee

242 Formation, Bashi Member; zone NP 10).

243

244

245 *Altaspiratella bearnensis* (Curry, 1982)

246 Figure 2.1-2

247

248 *v 1982 *Plotophysops bearnensis* Curry, p. 40, pl. 1, figure 9a-c.

249 v. (1986) *Spiratella tutelina* Curr. – Merle, p. 43 (*non* Curry).

250 v. 1990b *Altaspiratella bearnensis* (Curry 1981) – Janssen, p. 68.

251 ? 1992 *Altaspiratella bearnensis* (Curry) – Hodgkinson *et al.*, p. 13, pl. 1, figures 1,

252 2.

253 . (1996) *Altaspiratella bearnensis* (Curry, 1981) – Kunz, p. 164, pl. 30, figures. 1-3.

254 v. 2010 *Altaspiratella bearnensis* (Curry, 1982) – Cahuzac and Janssen, p. 24, pl. 2,

255 figures 1-4, pl. 3, figure 1.

256 v. 2013 *Altaspiratella bearnensis* – King *et al.*, pp. 192, 193.

257

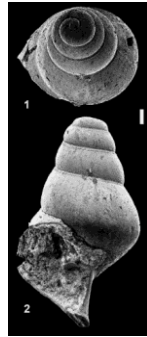


Figure 2. *Altaspiratella bearnensis* (Curry, 1982); RGM 777374. TDP 12.28.1, 66-76 cm; 1: apical view, 2: apertural view. Bar = 100 μ m.

Material examined. – TDP 12.28.1, 66-76 cm, depth 90.91-91.01 m below surface, composite depth 148.91-149.01 m (Table 2); RGM 777374 (1 specimen, Fig. 2.1-2, H = 1.36, W = 0.80 mm).

Description. – Only available specimen with high conical shell, 1.7 times higher than wide, and apical angle of *c.* 40°. Four and a half slightly convex and comparatively high whorls, regularly increasing in diameter and separated by incised, oblique suture. Whorls attach below periphery of preceding whorl. Specimen incomplete, last whorl missing, in shell preservation, but filled with pyrite.

Discussion. – Two closely resembling *Altaspiratella* species are currently known.

Altaspiratella elongatoidea (Aldrich, 1887) occurred during the earliest Eocene (Ypresian, nannoplankton zones NP 9 and 10) of the USA (Hodgkinson *et al.*, 1992; Janssen *et al.*, in review); *A. bearnensis* (Curry, 1982), introduced from the Ypresian

(zone NP 12/13) of SW France, is also known from the USA (with some doubt) in rocks of middle Eocene, Lutetian age (zone NP 15). These two species differ only very slightly in the proportions of their early whorls. Of *A. elongatoidea* no specimens preserving apertural structures are known, so there might be differences in that respect as well. In the single available specimen from Tanzania apertural structures are missing and its apical whorls take a position in between the two known species mentioned above. This could be an indication that these two taxa represent a single, long-ranging species. However, as long as apertural structures cannot be compared it seems better to keep them apart and the one available Tanzanian specimen is here considered, for the time being, to be the youngest representative of the *A. elongatoidea* - *A. bearnensis* lineage. The specimen extends the vertical range to the Priabonian (Zones P 16-17, NP 19-20; Nicholas *et al.*, 2006, figure 16).

Genus HELICONOIDES d'Orbigny, 1835

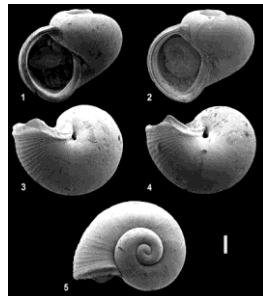
Type species. – *Atlanta inflata* d'Orbigny, 1834, by subsequent designation of Herrmannsen (1846, p. 514) = *Heliconoides inflatus* (d'Orbigny, 1834) (Recent).

Heliconoides nikkieae sp. nov.

Figures 3.1-5; 4

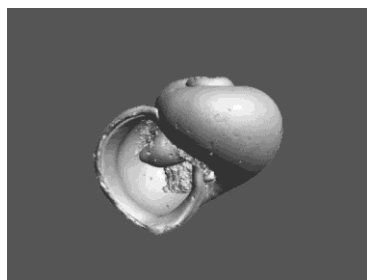
Type material. – Holotype RGM 777415a (Fig. 3.1); paratypes 1-2, RGM 777415b-c (Figures 3.3, 3-5) from the type locality. Kilwa Group, Pande Formation (upper Eocene,

299 Priabonian), biozones P 18 and NP 21; paratype 3, RGM 777 381 (Fig. 3-4), from TDP
300 17.21.1, 9-20 cm, Kilwa Group, Pande Formation (early Oligocene, Rupelian), biozones
301 P18 and NP 21.



302
303 **Figure 3.** *Heliconoides nikkieae* sp. nov.; 1: **Holotype**, RGM 777415a, apertural view;
304 2: apertural view, specimen lost; 3: **paratype 1**, RGM 777415b, umbilical view; 4:
305 **paratype 3**, RGM 777 381, umbilical view; 5: **paratype 2**, RGM 777415c, apical view.
306 Figures 3-1, -2, -3 and -5 from the type locality, TDP 17.36.1, 10-25 cm; Figure 3-4 from
307 TDP 17.21.1, 9-20 cm. Bar = 100 μ m.

308



309
310 **Figure 4.** *Heliconoides nikkieae* sp. nov.; **holotype**, RGM 777415a, 3dPDF.

311

312 *Additional specimens.* – TDP 11 (2 specimens), 12 (1 specimen) and 17 (49 specimens,
313 some of which with a query because of poor preservation) (Tables 1-3).

314

315 *Type locality.* – Stakishari (Tanzania, Kilwa region), cored borehole TDP 17, sample
316 17.36.1, 10-25 cm, 104.00-104.10 m below surface, composite depth 116.00-116.10 m.

317

318 *Etymology.* – Named after Nikkie Elert, the author's second granddaughter. At age six she
319 is, in many respects, more up-grown than many grown-ups. *Heliconoides* gender
320 masculine (ICZN 1992, art. 30.1.4.4).

321

322 *Diagnosis.* – Very small limacinid of 2.5 whorls in low-conical spiral, about as high as
323 wide, with large aperture, apertural margin externally thickened and internally doubled in
324 some specimens (Figure 3-2), basal margin with denticle. Apertural margin preceded by
325 about ten fine, margin parallel riblets.

326

327 *Description.* – Strikingly small limacinid of 2.5 whorls, height/width-ratio variable
328 between *c.* 1.19 (holotype H = 0.62, W = 0.52 mm, Figure 3-1) and 0.89 (*e.g.*, H = 0.50,
329 W = 0.56 mm, Figure 3-2), with depressed, low conical spire and distinct, incised suture.
330 Shell surface smooth and shiny, growth lines invisible. Last whorl large, occupying
331 almost entire shell height. Aperture relatively large, slightly higher than wide to almost
332 circular, occupying 4/5th of entire shell height. Apertural margin externally thickened by
333 narrow ridge, running all around margin, flexuous at base of shell, reaching umbilicus,
334 internally (Figure 3-2) in some specimens. Protruding denticle on basal part of margin

(Figures 3-3 and -4). Marginal thickening preceded by about ten fine, margin parallel orthocline riblets. Shell base regularly rounded, umbilicus very narrow, *c.* 1/20th of shell diameter.

Discussion. – In general shape the new species resembles somewhat the holotype of *Limacina wechesensis* Hodgkinson (1992, p. 21, pl. 5, figures 4-6) from the Lutetian of Texas, USA, but that species has one whorl more and reaches double the size of *H. nikkieae*, does not have the apertural structures of that species and its umbilicus is considerably wider. Size and apertural structures of the new species form a unique combination and cannot be compared to any limacinid currently known. The holotype was chosen from the sample with most specimens (13) of Priabonian age, but the species continues, in low numbers, well into the Rupelian part of the TDP 17 section.

Genus LIMACINA Bosc, 1817

Type species. – *Clio helicina* Phipps, 1774 by monotypy = *Limacina helicina* (Phipps, 1774) (Recent).

Limacina robusta (Eames, 1952)

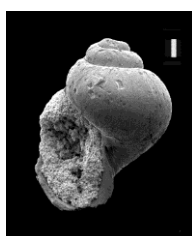
Figures 5, 6

*1952 *Aplexa robusta* Eames, p. 152, pl. 6, figure 149.

358

359 *Type material.* – Holotype (Figure 6) in the Natural History Museum, London, NHMUK
360 BM 68457, presented by the Burma Oil Co. Ltd, March 1950; Eames (1952) furthermore
361 recorded 19 paratypes from the type locality.

362



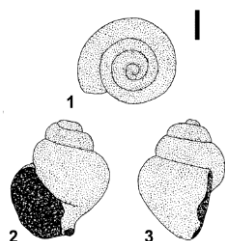
363

364 **Figure 5.** *Limacina robusta* (Eames, 195). TDP 17.37.1, 0-13 cm; RGM 777423b,
365 apertural view. Bar = 100 μ m.

366

367 *Type locality.* – Rahki Nala (Pakistan, western Punjab), 255' above base, local zone 9.
368 Lower Chocolate Clays; according to Afzal *et al.* (2009, p. 20) nowadays indicated as
369 Rahki Nala (Pakistan, western Punjab), 255' above base, local zone 9. Lower Chocolate
370 Clays; according to Afzal *et al.* (2009, p. 20) nowadays indicated as Kirthar Formation of
371 late Lutetian - Priabonian age of late Lutetian - Priabonian age.

372



373

Figure 6. *Limacina robusta* (Eames, 1952), **Holotype**, Natural History Museum, London
BM 68457. Rahki Nala (Pakistan, western Punjab); Kirthar Formation, Lower Chocolate
Clays (late Lutetian – Priabonian); 1: apical, 2: apertural, 3: lateral view. Bar = 100 μ m.

Material examined. – TDP 11 (1 specimen), TDP 12 (1 specimen), TDP 17 (16
specimens, all from the Priabonian), see Tables 1-3 for specification.

Description. – Most available specimens are juveniles and most in pyritic internal mould
preservation. Largest and best preserved specimen (Fig. 5) higher than wide ($H = 0.94$, W
 $= 0.72$ mm) with four convex whorls rapidly increasing in diameter. Aperture large, oval,
occupying more than half shell height. Base of shell regularly rounded with narrow
umbilicus.

Discussion. – Most of the available specimens are poorly preserved and juvenile, but the
illustrated Tanzanian specimen (Fig. 5)
has a shell height of 0.94 mm, whereas the holotype of *Limacina robusta* has a shell
height of 0.75 mm. Some of the smaller specimen have a somewhat wider apical angle
than the illustrated specimen. In spite of these small differences the Tanzanian specimens
are thought to represent the same species as the Pakistanian *L. robusta*, which has a
comparable age and was described from a pre-eastern-Paratethys locality under influence
of the Indian Ocean.

Initially the Tanzanian specimens were thought to represent *Limacina conica* (von
Koenen, 1892, p. 994, pl. 62, figures 5-6), a species described from the “early Oligocene”

of two localities (Atzendorf, Unseburg) in the eastern part of Germany. However, specimens from Atzendorf (NP 21 interval), made available by Arnold Müller (Leipzig, Germany) differ in shape and reach far larger dimensions. Their apical angle is smaller, the whorls are more convex and the aperture remains smaller than half shell height. Several species described by Hodgkinson *et al.* (1992), from the Paleogene of the United States also resemble the Tanzanian shells. Especially similar is *Limacina smithvillensis* Hodgkinson (*in* Hodgkinson *et al.*, 1992, p. 19, pl. 3, figure 16) from the Lutetian of Texas. However, that species reaches a shell height of 1.5 mm, has less convex whorls, a wider apical angle and its aperture occupies just half the shell height. Finally, that species is considerably older (Lutetian, NP 15) than the Priabonian (NP 21) specimens from Tanzania. At first glance also *L. stenzeli* Garvie (*in* Hodgkinson *et al.*, 1992, p. 19, pl. 4, figure 1) from the NP13-14 interval (Ypresian) of Texas is similar. However, that species should be included in the genus *Heliconoides*, because of its reinforced apertural margin.

Limacina tanzaniaensis sp. nov.

Figures 7.1-4; 8

Type material. – Holotype (Fig. 7.1) RGM 777428b; paratype 1 (Fig. 7.2), RGM 777438; paratypes 2 and 3 (Figures 7.3-4) RGM 777416b-c. Kilwa Group, Pande Formation (Eocene, Priabonian, biozones P 18 and NP 21).

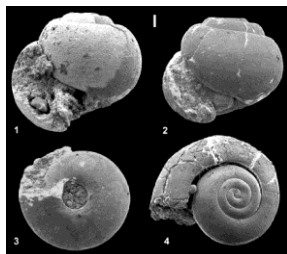


Figure 7. *Limacina tanzaniaensis* sp. nov. 1: **Holotype**, apertural view, RGM 777428b, TDP 17.38.2, 22-29 cm; 2: **paratype 1**, apertural view, RGM 777438, TDP 17.36.2, 80-95 cm; 3: **paratype 2**, umbilical view, RGM 777416b, TDP 17.36.1, 10-25cm; 4: **paratype 3**, apical view, RGM 777416c, same data as 3. Bar = 100 μ m.

Additional specimens. – Fourteen specimens from TDP 17 (Table 3), all from the Priabonian part of the section. Most specimens poorly preserved as pyritic internal moulds.

Type locality. – Stakishari (Tanzania, Kilwa region), cored borehole TDP 17.38.2, 22-29 cm, 111.12-111.19 m below surface = 123.12-123.19 m composite depth.

Diagnosis. – Spherical limacinid of almost four whorls, apex flattened. Last whorl inflated, 95% of total shell height, aperture *c.* 75% of total shell height, margin simple, base of shell umbilicate.

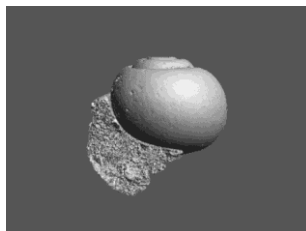


Figure 8. *Limacina tanzaniaensis* sp. nov.; **holotype**, apertural view, RGM 777428b, 3dPDF.

Description. – Limacinid of spherical shape, measurements of holotype H = 0.92, W = 1.00 mm, consisting of 3.75 moderately convex whorls with convex tangent. First two whorls flattened, last whorl very large, inflated, occupying 95% of total shell height. Aperture large, about 75% of total shell height, attaching on (holotype, Fig. 7.1) or slightly below (paratype 1, Fig. 7-2) periphery of penultimate whorl. Apertural margin simple, semicircular, inner margin and columella invisible as all specimens are in pyritic internal mould preservation. Base of shell regularly rounded, umbilicus present, 20-25% of shell diameter.

Discussion. – The available specimens demonstrate variability in height of the apex, in some the initial flattening continues to the third whorl, resulting in an only slightly raised apical shell part. There is some resemblance to the Ypresian species *Limacina heatherae* Hodgkinson in Hodgkinson *et al.* (1992, p. 17, pl. 2, figures 15-18). In that species, however, the whorls attach higher on the foregoing whorl, the aperture is narrower and its umbilicus smaller.

Limacina timi sp. nov.

Figures 9.1-4; 10

Type material. – Holotype (Fig. 9-1), RGM 1007748b; paratype 1 (Fig. 9-2), RGM777408a, from TDP 17.34.1, 0-7 cm, 98.90-98.97 m below surface and composite depth; paratype 2 (Fig. 9-3), RGM 777408b, TDP 17.34.1, 0-7 cm; paratype 3 (Fig. 9-4), RGM 777414m TDP 17.36.1, 0-5 cm. Kilwa Group, Pande Formation (Eocene, Priabonian); planktic foraminefera zone P 16-17, calcareous nannoplankton zone NP 19-20.

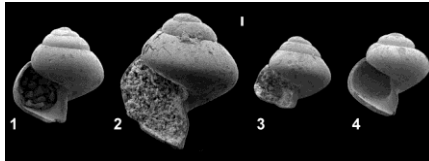
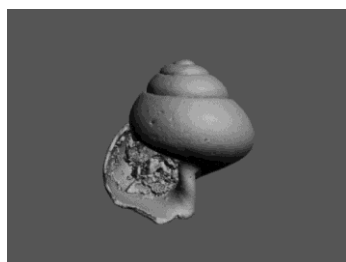


Figure 9. *Limacina timi* sp. nov.; 1: **holotype**, RGM 1007748b, TDP 12.27.1, 35-45 cm; 2: **paratype 1**, RGM 777408a TDP 17.34.1, 0-7 cm; 3: **paratype 2**, RGM 777408b, TDP 17.34.1, 0-7 cm; 4: **paratype 3**, RGM 777414m TDP 17.36.1, 0-5 cm. Apertural views, bar = 100 μ m.

Additional specimens. – Boreholes TDP 11, 12 and 17: 15 specimens (see Tables 1 and 3).

474 *Type locality.* – South of Stakishari, Tanzania, Kilwa Region, cored borehole TDP 12,
475 coordinates UTM 37L 560222-8981309, sample TDP 12.27.1, 35-45 cm = 89.60-89.70 m
476 below surface = 147.60-147.70 m composite depth.

477



478

479 **Figure 10.** *Limacina timi* sp. nov.; **holotype**, RGM 1007748b, 3dPDF.

480

481 *Etymology.* – Named after Tim Janssen, the author's six year old second grandson. For
482 him life is nonstop fun. Keep it that way, my friend!

483

484 *Diagnosis.* – Small limacinid of 3.75 whorls, slightly higher than wide. Whorls rounded
485 angular in juveniles, more regularly rounded in adults. Aperture lunate, somewhat higher
486 than half shell height. Base imperforate.

487

488 *Description.* – Holotype (H = 1.04, W = 1.00 mm) a small, regularly coiled, conical
489 limacinid of 3.75 slightly convex whorls, separated by incised suture. Last whorl
490 occupying 80% of shell height, slightly angular at periphery. Shell surface smooth,
491 growth lines invisible. Aperture large, lunate, 65% of shell height, no apertural
492 reinforcements present. Columellar side of aperture smooth, slightly concave internally,

straight externally. Base of shell regularly convex, umbilicus absent in holotype, extremely small or absent in juveniles. Angularity of whorls more clearly developed in juvenile specimens, resembling a rounded keel, but covered by following whorls attaching at the place of angularity, almost disappeared in adults. Juvenile specimens wider than high, in some of these the keel is rather strong and the apical angle wider (RGM 777425).

Discussion. – The angularity of especially juvenile specimens is not very clearly visible in the SEM images given here in Figures 9.3-4, but under light microscope, with illumination from left above it cannot be overlooked.

Limacina timi resembles somewhat the Ypresian species *L. gormani* (Curry, 1982) described from the Marnes de Gan Formation, of Gan, SW France (Curry, 1982; Cahuzac and Janssen, 2010). The same species or a closely related form was also described from the Stone City and Cook Mountain formations (Lutetian/Bartonian) of Texas, USA (Hodgkinson *et al.*, 1992, p. 19, pl. 3, figures 14-15) as *Limacina pygmaea* (*non* Lamarck, 1805), in which also the juvenile whorls are angular and the base imperforate. The adult shell, however, reaches one and a half times the size of *L. timi*, has a different, more spherical shape and a distinctly wider apical angle. Subsutural crests (Cahuzac and Janssen, 2010, pl. 11, figure 2) as seen in *L. gormani* are not present in *L. timi*.

Adult specimens of *Limacina timi* also resemble *Heliconoides nemoris* (Curry, 1965), described from the Bartonian of the UK and also recorded from the Priabonian of Biarritz, SW France by Curry (1982) and Cahuzac and Janssen (2010). That species belongs to the genus *Heliconoides* because of its apertural reinforcements that are apparently absent in

516 *L. timi*, but also its apex is flattened and juvenile specimens have no angular periphery.

517 *Limacina timi* straddles the Eocene/Oligocene boundary.

518

519

520 *Limacina* sp. 1

521 Figures 11.1-2

522

523 *Material examined.* – Five specimens, all from the same sample TDP 17.37.1, 0-13 cm;

524 106.90-107.03 m below surface, 118.90-119.03 composite depth; RGM 777440a-c (Table

525 3).



526

527 **Figure 11.** *Limacina* sp. 1. 1: RGM 777440c, TDP 17.37.1, 0-13 cm, apical view; 2:

528 RGM 777440b, same data, apertural view. Bar = 100 μ m.

529

530 *Description.* – Low conical limacinid of 3.75 moderately convex whorls attaching on

531 periphery of foregoing whorls. Aperture semicircular, slightly more than half shell height,

532 Base regularly rounded, umbilicus c. 1/6th of shell diameter.

533

534 *Discussion.* – Three of the available specimens are juveniles in poor preservation. In
535 specimen RGM 777440b (Fig. 11.2) the first whorl is missing. The two more adult
536 specimens show a striking resemblance with a limacinid illustrated by Lokno and Kumar
537 (2008, figure 3-2) from the Upper Disang Formation (Bartonian - Priabonian) of the Phek
538 District, south-central Nagaland (Assam - Arakan Basin), northeastern India, indicated by
539 these authors as “Limacinidae type A”. As these specimens were recorded from Indian
540 Ocean Basin rocks of more or less similar age they could very well represent the same
541 species as the Tanzanian ones. However, the material is insufficient for a reliable
542 identification. Lokno and Kumar compared their specimens with “*Limacina pygmaea*”, as
543 illustrated by Hodgkinson *et al.* (1992, pl. 3, figures 14-15), correctly stating that that
544 name “most certainly is incorrect”, as Hodgkinson *et al.*’s species seems to be closely
545 related to *L. gormani* (Curry, 1982) (see above) and not to *Limacina pygmaea*.

546
547 *Limacina* sp. 2

548 Figures 12.1-4

549
550 *Material examined.* – One specimen from TDP 12 (Table 2); 4 poorly preserved and
551 presumably juvenile specimens in pyritic internal mould preservation, from TDP 17
552 (Table 3).



Figure 12. *Limacina* sp. 2, RGM 777373, TDP 12.27.1, 35-45 cm; 1: apertural, 2: apical, 3: oblique apical, and 4: umbilical views. Bar = 100 μ m.

Description. – Limacinid of very low conical shape with almost flat, slightly raised or slightly concave apical plane. Width of illustrated specimen 1.20 mm, height 0.86 mm. Whorls 3.75, regularly increasing in diameter. Aperture semicircular, occupying c. 80% or more of total shell height, reaching to far beyond base of preceding whorl. Apertural structures absent or not preserved. The apparently present groove along the apertural margin, as visible in Fig. 13.2-3, is considered to be caused by damage of the mould. Base perforated by umbilicus of 1/5th to 1/7th of shell diameter.

Discussion. – Several limacinids with an almost planorboid shell shape are known from the Eocene-Oligocene interval in Europe, Asia and the USA. Some of these are characterised by having a slightly concave apical plane, or, in other cases, by an irregular

development of early whorls. In the present specimens, however, the whorls are in a regular spiral and the apical plane is a bit raised, with the first 1.5 whorls flattened (Fig. 12.3).

Very similar is a species from the early Oligocene of Japan, described as *Limacina karasawai* Ando (2011, p. 248, figures 3.1-2. This species was said to have three quarters of a whorl more than the Tanzanian specimen illustrated here, but we fail to see that from Ando's photographs. Also closely similar is *Limacina canadaensis* Hodgkinson (in Hodgkinson *et al.*, 1992, p. 16, pl. 2, figures 4-6), but its last whorl seems to be relatively lower. This species was collected from downhole contaminated cutting samples and could be anything between early Eocene and early Oligocene. An occurrence of similar age (earliest Oligocene) was described from the North Sea Basin and is also known from contemporaneous rocks (base of Viborg Formation) in Jylland, Denmark, as *Limacina mariae* Janssen (1989, p. 111, pl. 4, figures 2-5), but that species always has a concave apical spiral, has a somewhat wider umbilicus and reaches to over 2 mm shell width.

Finally, as yet unpublished similar material is available from the Eocene - Oligocene interval in the NE United States and from the eastern part of Germany.

A reliable interpretation of all these forms depends on a larger material for comparisons and therefore the present specimens, apart from the illustrated specimen in poor condition, are left in open nomenclature.

Limacinidae indet.

591 *Material examined.* – One specimen from TDP 12.24.3, 0-10 cm; 10 specimens from
592 TDP 17 (see Table 3).

593

594 *Description.* –These specimens are in internal pyritic mould preservation, more or less
595 strongly deteriorated and cannot be identified any further.

596

597

598 Superfamily CAVOLINIOIDEA Gray, 1850 (1815) [= Hyaline Rafinesque, 1815]

599 Family CRESEIDAE Rampal, 1973

600 Genus BOVICORNU Meyer, 1886

601

602 *Type species.* – *Bovicornu eocenense* Meyer, 1886, by monotypy (early Oligocene, USA).

603

604 *Discussion.*– Shortly after its introduction by Meyer (1886) the validity of the genus

605 *Bovicornu* was denied by Dall (1892, p. 302), who considered both species introduced in

606 that genus by Meyer (1886, 1887) to belong to the caecid (benthic) genus *Meioceras*

607 Carpenter, 1858. Dall's point of view was followed, with doubts, by Cossmann (1912, p.

608 154-155), but rejected by later authors (Collins, 1934, p. 212; van Winkle Palmer, 1947,

609 p. 464; Hodgkinson *et al.*, 1992, p. 24). Indeed, in *Meioceras* the shell wall is

610 considerably thicker and its protoconch is utterly different, whereas in *Bovicornu* the shell

611 is thin-walled and the larval parts agree with species of the pteropod genus *Creseis*. Zilch

612 (1959, p. 49) followed Collins and included *Bovicornu* with a query in the cavoliniid

613 pteropods, considering it a possible synonym of *Euchilotheca* Fischer, 1882, which is also

614 a creseid genus, but its type species, *E. succincta* (Defrance, 1828), shows only very faint
615 traces of spiralisation. We agree with Hodgkinson *et al.* that *Bovicornu* should be
616 considered an independent genus in Creseidae.
617 *Bovicornu* species demonstrate a certain resemblance with *Hameconia edmundi* Janssen,
618 2008, described from the late Oligocene (Chattian) of SW France. In that species the
619 bilaterally symmetrical shell has a curvature of *c.* 180°, but it is curved in one plain, not in
620 a spatial spiral. The curvature is dorso-ventral, as is clear by the presence of lateral
621 carinae. Its larval stage differs from *Bovicornu* in having separate protoconchs 1 and 2.
622 Janssen and Maxwell (*in* Janssen, 1995, p. 164), Janssen (2008, p. 160) and Cahuzac and
623 Janssen (2010, p. 111) included the genus *Hameconia* in the Sphaerocinidae family.

624

625

626 *Bovicornu* aff. *eocenense* Meyer, 1886

627 Figures 13.1-2

628

629 cf 1886 *Bovicornu eocenense* Meyer, p. 79, pl. 3, figure 12 (not figure 2).

630 cf 1892 *Meioceras eocenense* (Meyer) – Dall, p. 302.

631 cf 1912 *Bovicornu eocænense* [*sic*] Meyer – Cossmann, p. 155.

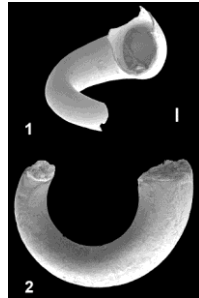
632 cf 1934 *Bovicornu eocenense* Meyer – Collins, p. 212, pl.9, figure 3, pl. 13, figure 5.

633 cf 1959 *Bovicornu eocenense* O. Meyer – Zilch, p. 49, figure 164.

634 cf 1992 *Bovicornu eocenense* Meyer – Hodgkinson *et al.*, p. 24, pl. 7, figures 9-10 (with
635 additional synonymy).

636

637 *Type material.* – Holotype (H = 2,8, W = 0,7 mm) United States Natural Museum
638 (Smithsonian Institution) nr. 644596.
639



640
641 **Figure 13.** *Bovicornu* aff. *eocenense* Meyger, 1886; 1. TDP 12.23.3, 89-96 cm; RGM
642 777370, apertural view; 2. TDP 17.36.1, 20-35 cm, RGM 1007784, basal view. Bar = 100
643 μm .

644
645 *Type locality.* – Red Bluff, USGS locality 5264, Mississippi, USA (Red Bluff Clay;
646 Oligocene, Rupelian, NP 21).

647
648 *Material examined.* – Only fragments were found, 2 from TDP 11, 9 from TDP 12 and 11
649 from TDP 17 (see Tables 1-3 for details).

650
651 *Description.* – Spatially spiralised tube with free volutions. Transverse section of tube
652 circular, no surface ornamentation or growth lines visible. Only smaller fragments are
653 available among which no complete aperture or protoconch. Diameter of tube doubles in
654 about half a volution.

655

656 *Discussion.* – Two species of the genus *Bovicornu* are currently known and both are
657 exclusively recorded from the USA. The older one of these, *B. gracile* Meyer, 1887 (p. 9,
658 pl. 2, figure 17), of the Moodys Branch Formation, Texas, has an age of Bartonian (NP
659 17). The other species, *B. eocenense* Meyer, 1886 (p. 79, pl. 3, figure 12), from the Red
660 Bluff Formation of Mississippi, USA occurred during the Priabonian and early Rupelian
661 (NP19-21) (Hodgkinson *et al.*, 1992, figure 3). In both species the shell is creseid, but
662 instead of being straight or slightly curved, as in *Creseis* species, the tube demonstrates
663 clear twisting in a wide spatial spiral that was said to be stronger in *B. gracile*.
664 Hodgkinson *et al.* (1992), however, collected numerous specimens at both type localities
665 and noted that in many specimens of both species the twisting is stronger and that the
666 species cannot be distinguished on the degree of twisting. There is, however, a clear
667 difference in protoconch morphology: an inflated bulb in *B. eocenense* and a more
668 cylindrical shape in *B. gracile*. The holotypes of both species were re-illustrated in
669 Hodgkinson *et al.* (1992, pl. 7, figures 9-10 and 11-12), reproduced herein as Figs 14.1-4.
670 Although the few larger fragments from Tanzania seem to indicate a considerably
671 stronger spirally twisted shell than in either of the holotypes it is preferred to indicate the
672 Tanzanian species as related to the younger of the two American species and is indicated
673 here as *Bovicornu* aff. *eocenense*. Once specimens preserving their protoconch become
674 available this position might be revised.



Figure 14. Holotypes of *Bovicornu eocenense* Meyer, 1886 (1, 2) and *B. gracile* Meyer, 1887 (3, 4). Shell height of 1 = 2.8 mm, of 2 = 2.7 mm; 2 and 4 are magnifications of 1 and 3, respectively. Photographs after Hodgkinson *et al.* (1992, pl. 7, figures 9-12).

RESULTS

The number of specimens per species/per sample are specified in Tables 1-3. Vertical ranges of the species in the three Tanzanian sections are shown together in Figure 15. Eight pteropod species are recognised, seven of them belonging to the Limacinidae and one (*Bovicornu*) to the Creseidae. Three of the limacinids are represented by a single or very few specimens only. Three species could only be identified in open nomenclature. Three of the limacinid species are described as new.

TDP Core section	Depth interval (cm)	Subsurface depth (m)	Composite depth (m)	<i>Heliconoides nikkieae</i>	<i>Limacina robusta</i>	<i>Limacina timi</i>	<i>Bovicornu</i> aff. <i>eocenense</i>
11.26.2	64-74	82.14	98.14	1	-	-	-
11.32.1	33-40	89.63	109.63	-	1	-	-
11.32.3	64-72	91.94	111.94	1	-	1	-
11.33.2	74-84	94.04	114.04	-	-	-	2

Table 1. Distribution of pteropod species in core TDP 11.

TDP Core section	Depth interval (cm)	Subsurface depth (m)	Composite depth (m)	<i>Altaspiratella bearnensis</i>	<i>Heliconoides nikkieae</i>	<i>Limacina robusta</i>	<i>Limacina timi</i>	<i>Limacina</i> sp. 2	<i>Limacina</i> indet.	<i>Bovicornu</i> aff. <i>eocenense</i>
12.11.4	20-26	38.20	96.20	-	-	-	1	-	-	-
12.12.1	23-31	38.23	96.23	-	1	-	-	-	-	-
12.14.1	47-48,5	44.47	102.47	-	-	-	-	-	-	2
12.14.1	51-53	44.51	102.51	-	-	-	-	-	-	1
12.14.1	56-58	44.56	102.56	-	-	-	-	-	-	1
12.14.3	23-31	46.23	104.23	-	-	-	-	-	-	1
12.18.3	65-76	58.65	116.65	-	-	-	1	-	-	-
12.19.1	22-34	59.22	117.22	-	-	-	1	-	-	-
12.21.1	37-48	65.37	123.37	-	-	-	1	-	-	-
12.23.3	89-96	73.89	131.89	-	-	-	-	-	-	1
12.24.3	0-10	76.30	134.30	-	-	-	-	-	1?	-
12.26.2	54-62	81.79	139.79	-	-	-	1	-	-	-
12.27.1	35-45	89.60	147.60	-	-	1	1	1	-	-
12.28.1	66-76	90.91	148.91	1	-	-	-	-	-	2
12.29.1	25-35	91.90	149.90	-	-	-	1	-	-	-
12.29.2	20-30	92.85	150.75	-	-	-	-	-	-	-

Table 2. Distribution of pteropod species in core TDP 12.

731
732

733
734
735

736

737
738
739
740
741
742
743
744
745
746
747
748
749
750
751
752
753
754
755
756
757
758
759
760
761
762
763
764
765
766
767
768
769
770
771
772

TDP Core section	Depth interval (cm)	Subsurface depth (m)	Composite depth (m)	<i>Heliconoides nikkieae</i>	<i>Limacina robusta</i>	<i>Limacina tanzaniaensis</i>	<i>Limacina timi</i>	<i>Limacina</i> sp. 1	<i>Limacina</i> sp. 2	Limacinae indet.	<i>Bovicornu</i> aff. <i>eocenense</i>
17.15.1	50-63	47.95	47.95	2	-	-	-	-	-	-	-
17.17.2	0-14	54.45	54.45	2?	-	-	-	-	-	-	-
17.18.3	10-25	58.55	58.55	2	-	-	1	-	-	-	-
17.21.2	9-20	62.87	62.87	1	-	-	-	-	-	-	-
17.23.2	0-10	66.57	66.57	-	-	-	1?	-	-	-	-
17.23.3	0-13	67.54	67.5	2?	-	-	-	-	-	-	-
17.23.3	99-105	68.56	68.56	2	-	-	-	-	-	-	-
17.24.1	20-35	68.30	68.30	5	-	-	-	-	-	-	-
17.24.2	0-15	69.10	69.10	-	-	-	1	-	-	-	-
17.25.1	0-15	71.40	71.40	2	-	-	-	-	1	-	-
17.25.1	50-60	71.90	71.90	1	-	-	-	-	-	-	-
17.25.2	48-58	72.88	72.88	1	-	-	1	-	-	-	-
17.26.3	10-25	77.90	77.90	2	-	-	-	-	-	-	-
17.31.1	20-35	89.10	89.10	2?	-	-	-	-	-	-	-
17.31.2	0-15	89.90	89.90	1	-	-	-	-	-	-	-
17.31.3	0-15	90.90	90.90	3?	-	-	-	-	-	-	-
17.31.4	0-12	91.90	91.90	1?	-	-	-	-	-	-	-
17.32.1	10-25	92.00	92.00	2?	-	-	-	-	-	-	-
17.32.2	0-15	92.90	92.90	3	-	-	-	-	1	1?	1
17.32.4	14-20	95.04	95.04	1	-	-	-	-	-	-	1
17.33.1	3-18	95.93	95.93	1	-	-	-	-	-	-	2
17.33.3	0-15	97.90	97.90	3	-	-	1	-	-	-	1
17.34.1	0-7	98.90	98.90	4?	-	-	2	-	-	3?	1
17.34.2	91-99	100.81	100.81	-	-	-	-	-	-	-	2
17.36.1	0-5	103.90	115.90	1	-	-	1	-	2?	1?	2
17.36.1	5-13	103.95	115.95	-	-	1	-	-	-	-	-
17.36.1	10-25	104.00	116.00	11	-	7	-	-	-	-	3
17.36.2	52-59	105.42	117.42	-	1	-	-	-	-	3?	-
17.36.2	80-95	105.70	117.70	3+5?	2	1	2?	-	-	2?	-
17.37.1	0-13	106.90	118.90	-	6	-	-	5	-	-	-
17.37.1	32-47	107.22	119.22	-	1	-	-	-	-	-	-
17.38.1	25-35	110.15	122.15	-	-	5	1	-	-	-	-
17.38.2	22-29	111.12	123.12	-	1	2	-	-	-	-	-
17.38.2	62-70	113.52	123.52	-	-	-	1	-	-	-	-
17.41.1	8-18	119.48	131.48	-	-	1	2?	-	-	-	-
17.41.3	0-15	121.40	133.40	-	1	1	-	-	-	-	-

Table 3. Distribution of pteropod species in core TDP 17 (question mark denotes poorly preserved specimens).

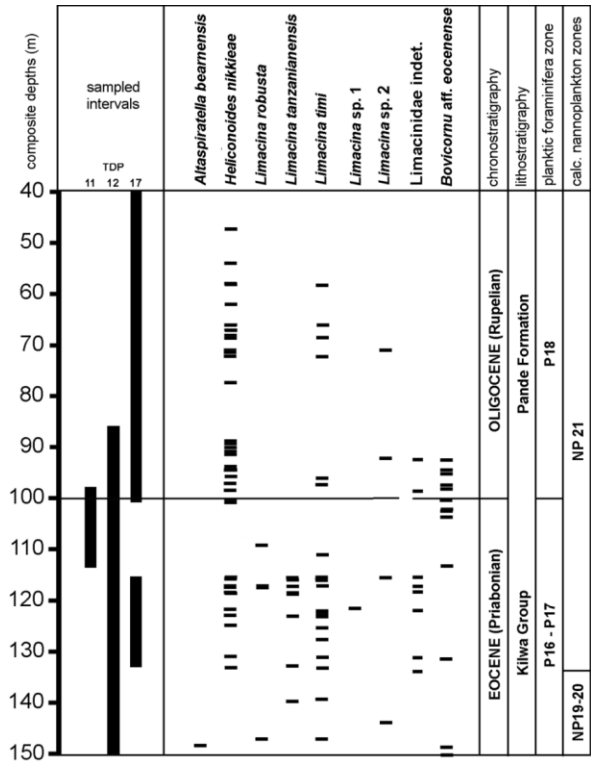


Figure 15. Range chart of pteropods, combined data of TDP 11, 12 and 17, calibrated to composite depths. Basic stratigraphical data mainly from Pearson *et al.* (2008, figure 16).

The stratigraphic ranges are remarkable. The two most commonly occurring species, *Heliconoides nikkieae* and *Limacina timi* (both introduced herein), and the less frequently represented *Limacina* sp. 2 occur in comparable numbers both below and above the EOB

and do not seem to be influenced by changing environmental conditions. Two other species, *Limacina robusta* and *L. tanzaniaensis*, on the contrary, disappear some 10 m below the EOB, at a level close to the extinction of the *Turborotalia cerroazulensis*-group of planktic foraminifera and the first oxygen isotope step (Step 1 of Pearson *et al.*, 2008, Lear *et al.*, 2008) although improved sampling could of course extend the ranges up to the EOB or beyond. One species, *Altaspiratella bearnensis*, was found as a single specimen only in one of the lowermost samples. Finally, a single species (*Bovicornu* aff. *eocenense*), is irregularly distributed in the Eocene part of the sections, but disappears in the basal 10 m of the Oligocene. No pteropods are found exclusively in the Oligocene part of the record.

DISCUSSION

The newly described taxa, as well as species recorded in open nomenclature, do not present clues that can be applied in long distance correlations. The present material from Tanzania furthermore originates from a very large area from which Paleogene pteropods have never been recorded previously. From the African continent only a late Eocene occurrence is known from Nigeria (Bende Ameki; Curry, 1965 and Naturalis collection) and some Eocene (Ypresian, Bartonian) material was collected from Egypt (Valley of the Kings) by the late Chris King (Naturalis Biodiversity Center collections, unpublished). The presence of *Altaspiratella bearnensis* is quite unexpected. That species is only known from the Aquitaine Basin in SW France (type locality, Ypresian) and from the southern USA (Ypresian and Lutetian). Its presence in the Tanzanian Priabonian is surprising. The

807 record, however, is based on a single incomplete specimen and better preserved material
808 might prove that another species is involved.

809 The occurrence of *Limacina robusta* is interesting. That species was originally described
810 from Pakistan, a locality under the influence of the Indian Ocean, what is equally the case
811 for the Tanzanian material. If *Limacina* sp. 1 is indeed closely related to or even identical
812 with ‘Limacinidae type A’ as described by Lokho and Kumar (2008) from northern India
813 it similarly represents an interesting palaeogeographical occurrence on the Indian Ocean.

814 The creseid species *Bovicornu eocenense* is to date exclusively known from the United
815 States and its vertical distribution includes the Priabonian and the Rupelian transition,
816 which is consistent with its range in the Tanzanian material. However, the Tanzanian
817 material is poorly preserved and better specimens might lead to another specific
818 interpretation.

819 When compared with previous isotope and microfossils from the Tanzanian material it
820 may be concluded that three pteropod species (*Heliconoides nikkieae*, *Limacina timi* and
821 *Limacina* sp. 2) do not show clear response following the drastic climatic cooling at the
822 EOB, but two species (*Limacina robusta* and *L. tanzaniaensis*) seem to disappear close to
823 the cooling related step of the EOT, perhaps indicating a temperature sensitivity if a true
824 disappearance. The last occurrence of these two taxa precedes that of the planktic
825 foraminiferal family Hantkeninidae and the extinction level of the larger benthic
826 foraminifera (Pearson *et al.*, 2008; Wade and Pearson, 2008; Cotton and Pearson, 2011).
827 However, it is similar to the last occurrence of the planktic foraminifera *Turborotalia*
828 (Pearson *et al.*, 2008; Wade and Pearson, 2008). In addition, the nannofossil assemblage
829 has shown an increase in nutrient loving taxa close to the onset of the EOT (Dunkley

Commented [U3]: AWJ: previous isotope and microfossil data
??

830 Jones *et al.*, 2008), suggesting nutrient increase in the water column as a potential
831 contributing factor.

832 There are currently plans to re-core the EOB of Tanzania with wide diameter boreholes
833 (Pearson and Hudson, 2014). Much larger samples and denser sampling may shed further
834 light on the pteropod record across the EOB in the region.

Commented [U4]: AWJ: Do we keep this statement or is it merely wishful thinking ?

836 CONCLUSION

837
838 Here we have shown a small but important insight into pteropod fauna from both an
839 under-represented geological time and geographic region. Eight species were identified,
840 three of which were new. Two of the taxa show an apparent, at least, local extinction
841 close to the first major cooling step of the EOT whilst the others seem unaffected, or have
842 too few occurrences to tell. The occurrences of several taxa, though not enough for long
843 distance correlation, are surprising with *Altaspiratella bearnensis*, and *Limacina robusta*
844 only previously known from Europe and the U.S.A., and from the U.S.A., respectively.
845 The occurrence of possible *Bovicornu eocenense* which is only previously known from
846 Pakistan very tentatively suggests an Indo-Pacific connection. This therefore underlines
847 the need for increased studies of older pteropod occurrences, particularly in that may be
848 under-sampled but have good preservation potential. Furthermore it shows the importance
849 of carrying out studies of less conventional micro-fossils on cores generally used for
850 foraminiferal or nannoplankton studies. (or something like that?)

Commented [U5]: AWJ: better 'recognised', as some remain in open nomenclature

Commented [U6]: AWJ: better leave this out, as we do some long-distance correlation with *Limacina robusta*

Commented [U7]: No, *Bovicornu* aff. *eocenense*

Commented [U8]: AWJ: No, *Limacina robusta* should be here

Commented [U9]: AWJ: not so 'tentatively'... I think that quite clear. Better leave it out.

Commented [U10]: AWJ: I'd rather suggest to say Paleogene instead of older

Commented [U11]: AWJ: Here something seems to be missing. What did you want to say here ?

ACKNOWLEDGEMENTS

The authors thank Yusuke Ando (Graduate School of Environmental Studies, Nagoya University, Furo-cho, Chikusa-ku, Nagoya, Japan, for discussion on Paleogene pteropod occurrences, and Arnold Müller (Institut für Geophysik und Geologie, Geologisch-Palaeontologische Sammlung, Leipzig, Germany) for providing pteropod material for comparisons. Renate Helwerda of Naturalis Biodiversity Center, Leiden, NL, was kind enough to make the SEM images. Dirk van der Maarel of the same institute made the CT-scannings. The Tanzanian Petroleum Development Corporation and COSTECH are acknowledged for their support of the TDP. Two anonymous reviewers are thanked for critically reading the manuscript and for their helpful comments.

REFERENCES

- Afzal, J., Williams, M. and Aldridge, R.J. 2009. Revised stratigraphy of the lower Cenozoic succession of the Greater Indus Basin in Pakistan. *Journal of Micropalaeontology* 28: 7-23.
- Aldrich, T.H. 1887. Notes on Tertiary fossils, with descriptions of new species. *The Journal of the Cincinnati Society of Natural History*, 10(2):78-83.
- Ando, Y. 2011. Oligocene pteropods (Gastropoda: Thecosomata) from the Kishima Formation, Saga Prefecture, southwest Japan. *Revista Mexicana de Ciencias Geológicas* 28(2): 245-253.

874 Bednaršek, N., Ohman, M. D., 2015. Changes in pteropod distributions and shell
 875 dissolution across a frontal system in the California Current System. *Marine*
 876 *Ecology Progress Series*, 523: 93–103. doi: 10.3354/meps11199
 877 Blainville, [H.M.D.] de. 1824. Mollusques, Mollusca (Malacoz.). *Dictionnaire des*
 878 *Sciences naturelles*, 32:1-392.
 879 Bodiselitsch, B., Montanari, A., Koeberl, C. and Coccioni, R. 2004, Delayed climate
 880 change in the late Eocene caused by multiple impacts: High-resolution
 881 geochemical studies at Massignano, Italy. *Earth and Planetary Science Letters*
 882 223: 283–302. doi: 10.1016/j.epsl.2004.04.028.
 883 Bosc, [L.A.G.]. 1817. Limacine. *Nouveau Dictionnaire d'Histoire naturelle*, 18: 42.
 884 Bown, P.R., Dunkley Jones, T., Lees, J.A., Randell, R.D., Mizzi, J.A., Pearson, P.N.,
 885 Coxall, H.K., Young, J.R., Nicholas, C.J., Karega, A., Singano, J. and Wade, B.S.
 886 2008. A Paleogene calcareous microfossil Konservat-Lagerstätte from the Kilwa
 887 Group of coastal Tanzania. *Geological Society of America Bulletin* 120(1/2): 3–
 888 12. doi: 10.1130/B26261.1
 889 Burmeister, H. 1837. *Handbuch der Naturgeschichte zum Gebrauch bei Vorlesungen;*
 890 *zweite Abtheilung., Zoologie*. T.C.F. Enslin, Berlin.
 891 Cahuzac, B. and Janssen, A.W. 2010. Eocene to Miocene holoplanktonic Mollusca
 892 (Gastropoda) of the Aquitaine Basin, southwest France. *Scripta Geologica* 141:1-
 893 193.
 894 Cande, S.C. and Kent, D.V. 1995. Revised calibration of the geomagnetic polarity
 895 timescale for the late Cretaceous and Cenozoic: *Journal of Geophysical Research*
 896 100(B4): 6093-6095.

897 Carpenter, P.P. 1858. First steps towards a monograph of the Cæcidæ, a family of
898 rostriferous Gasteropoda. *Proceedings of the Zoological Society of London* 26:
899 413-443.

900 Collins, R.L. 1934. A monograph of the American Tertiary pteropod mollusks. *Johns*
901 *Hopkins University Studies in Geology* 11: 137-234.

902 Cossmann, M. 1912. *Essays de paléoconchologie comparée*, 9. Paris (Cossmann and
903 Lamarre): 1-215.

904 Cotton, L.J. and Pearson, P.N. 2011. Extinction of larger benthic foraminifera at the
905 Eocene/Oligocene boundary. *Palaeogeography, Palaeoclimatology,*
906 *Palaeoecology* 311: 281–296.

907 Cotton, L.J., Arciszewski, J., Reich, S., Wesselingh, F. and Pearson, P.N. (in prep.).
908 Molluscan response across the Eocene – Oligocene Transition: Tanzania Drilling
909 Project sites 11, 12 and 17.

910 Coxall, H.K., Pearson, P.N., 2007. The Eocene-Oligocene transition. In Williams, M.,
911 Haywood, A.M., Gregory, F.J., Schmidt, D.N. (eds). Deep-time perspectives on
912 climate change: marrying the signal from computer models and biological proxies:
913 *The Micropalaeontological Society, London*, Special publication: 351–387.

914 Curry, D., 1965. The English Palaeogene pteropods. *Proceedings of the Malacological*
915 *Society of London* 36: 357-371.

916 Curry, D. 1982. Ptéropodes éocènes de la tuilerie de Gan (Pyrénées-Atlantiques) et de
917 quelques autres localités du SW de la France. *Cahiers de Micropaléontologie* 4
918 (1981): 35-44.

919 Cuvier, G. 1795. Second mémoire sur l'organisation et les rapports des animaux à sang
 920 blanc, dans lequel on traite de la structure des mollusques et de leur division en
 921 ordre, lu à la Société d'Histoire Naturelle de Paris, le 11 prairial an troisième [30
 922 May 1795]. *Magazin Encyclopédique, ou Journal des Sciences, des Lettres et des*
 923 *Arts*, 1:433-449.

924 Cuvier, G. 1804. Mémoire concernant l'animal de l'Hyale, un nouveau genre de
 925 mollusques nus, intermédiaire entre l'Hyale et le *Clio*, et l'établissement d'un
 926 nouvel ordre dans la classe des mollusques. *Annales du Muséum national*
 927 *d'Histoire Naturelle* 4: 223-234.

928 Dall, W.H., 1892. Contributions to the Tertiary fauna of Florida, with especial reference
 929 to the Miocene silex-beds of Tampa and the Pliocene beds of the Caloosahatchie
 930 River, 2. *Transactions of the Wagner Free Institute of Science of Philadelphia*
 931 3(2): 201-473.

932 Defrance, [J.L.M.] 1828. *Vaginella*. (Foss.). *Dictionnaire des Sciences naturelles* 56: 427.

933 Diester-Haass, L, Zahn, R., 2001. Paleoproductivity increase at the Eocene-Oligocene
 934 climatic transition: ODP/DSDP Sites 763 and 592. *Palaeogeography,*
 935 *Palaeoclimatology, Palaeoecology*. 172(1): 153-170.

936 Di Martino, E., Taylor, P., Cotton, L.J., and Pearson, P.N. 2017. First bryozoan fauna
 937 from the Eocene/Oligocene transition in Tanzania. *Journal of Systematic*
 938 *Palaeontology*. <http://dx.doi.org/10.1080/14772019.2017.1284163>

939 Dunkley Jones, T., Bown, P.R., Pearson, P.N., Wade, B.S., Coxall, H.K. and Lear, C.H.,
 940 2008. Major shifts in calcareous phytoplankton assemblages through the

941 Eocene-Oligocene transition of Tanzania and their implications for low-latitude
 942 primary production. *Paleoceanography*, 23(4).

943 Eames, F.E. 1952. A contribution to the study of the rocks in western Pakistan and
 944 western India, C. The description of the Scaphopoda and Gastropoda from
 945 standard sections in the Rakhi Nala and Zinda Pir areas of the western Punjab and
 946 in the Kohat District. *Philosophical Transactions of the Royal Society of London*
 947 (B)631(236) : 1-168.

948 Fischer, P., 1880-1887. *Manuel de conchyliologie et de paléontologie conchyliologique*
 949 *ou histoire naturelle des mollusques vivants et fossiles suivi d'un appendice sur les*
 950 *brachiopodes par D.P. Oehlert*. Paris (Savy): 1-12, i-xxiv, 1-1369.

951 Gray, J.E. 1847. A list of the genera of recent Mollusca, their synonyms and types.
 952 *Proceedings of the Zoological Society of London*, 15 (78):129-219.

953 Herrmannsen, A.N. 1846-1852. *Indicis generum malacozoorum primordia. Nomina*
 954 *subgenerum, generum, familiarum, tribuum, ordinum, classium: adjectis*
 955 *auctoribus, temporibus, locis systematicis atque literariis, etymus, synonymis.*
 956 *Praetermittuntur Cirripedia, Tunicatae et Rhizopoda*, 1 (1846-1847), 2 (1847-
 957 1849), supplement (1852). T. Fischer, Cassel.

958 Hodgkinson, K.A., Garvie, C.L. and Bé, A.W.H. 1992. Eocene euthecosomatous
 959 Pteropoda (Gastropoda) of the Gulf and eastern Coasts of North America.
 960 *Bulletins of American Paleontology*, 103(341):5-62.

961 Janssen, A.W. 1989. Some new pteropod species from the North Sea Basin Cainozoic
 962 (Mollusca: Gastropoda, Euthecosomata). *Mededelingrn vsn de Werkgroep voor Tertiaire*
 963 *en Kwartaire Geologie* 26: 91-133.

964 Janssen, A.W. 1990. Long distance correlation of Cainozoic deposits by means of planktonic
 965 gastropods (“pteropods”); some examples of future possibilities. *Tertiary Research*,
 966 11,65-72.

967 Janssen, A.W. 1995. Systematic revision of holoplanktonic Mollusca in the collections of the
 968 “Dipartimento di Scienze della Terra” at Torino, Italy. *Monografie Museo Regionale di*
 969 *Scienze Naturali Torino* 17: 1-233.

970 Janssen, A.W. 2008. Notes on the systematics, morphology and biostratigraphy of fossil
 971 holoplanktonic Mollusca, 20. A new pteropod genus and species, *Hameconia*
 972 *edmundi* gen. nov. spec. nov. (Mollusca, Gastropoda, Sphaerocinidae), from the Late
 973 Oligocene of SW France. *Basteria* 72: 159-163.

974 Janssen, A.W., Sessa, J.A. and Thomas, E. Pteropoda (Mollusca, Gastropoda, Thecosomata) from
 975 the Paleocene-Eocene Thermal Maximum (United States Atlantic Coastal Plain).
 976 *Palaeontologia Electronica*, 19.3.47A: 1-26. [palaeo-electronica.org/content/2016/1662-](http://palaeo-electronica.org/content/2016/1662-pteropoda-from-the-usa-petm)
 977 [pteropoda-from-the-usa-petm](http://palaeo-electronica.org/content/2016/1662-pteropoda-from-the-usa-petm)

978 Jiménez Berrocoso, Á., MacLeod, K.G., Huber, B.T., Lees, J.A., Wendler, I., Bown, P.R.,
 979 Mweneinda, A.K., Londoño, C.I. and Singano, J.M. 2010. Lithostratigraphy,
 980 biostratigraphy and chemostratigraphy of Upper Cretaceous sediments from
 981 southern Tanzania: Tanzania drilling project sites 21–26. *Journal of African Earth*
 982 *Sciences* 57 (2010) 47–69.

983 Jiménez Berrocoso, Á., Huber, B.T., MacLeod, K.G., Petrizzo, M.R., Lees, J.A.,
 984 Wendler, I., Coxall, H., Mweneinda, A.K., Falzoni, F., Birch, H., Singano, J.M.,
 985 Haynes, S., Cotton, L., Wendler, J., Bown, P.R., Robinson, S.A. and Gould, J.
 986 2012. Lithostratigraphy, biostratigraphy and chemostratigraphy of Upper

987 Cretaceous and Paleogene sediments from southern Tanzania: Tanzania Drilling
 988 Project Sites 27–35. *Journal of African Earth Sciences* 70: 36–57.

989 Jiménez Berrocoso, Á., Huber, B.T., MacLeod, K.G., Petrizzo, M.R., Lees, J.A.,
 990 Wendler, I., Coxall, H., Mweneinda, A.K., Falzoni, A., Birch, H., Haynes, S.J.,
 991 Bown, P.R., Robinson, S.A. and Singano, J.M. 2015. The Lindi Formation (upper
 992 Albian–Coniacian) and Tanzania Drilling Project Sites 36–40 (Lower Cretaceous
 993 to Paleogene): Lithostratigraphy, biostratigraphy and chemostratigraphy. *Journal*
 994 *of African Earth Sciences* 101: 282–308.

995 King, C., Iakovleva, A., Steurbaut, E., Heilmann-Clausen, C. and Ward, D. 2013. The
 996 Aktulagay section, west Kazakhstan: a key site for northern mid-latitude early
 997 Eocene stratigraphy. *Stratigraphy*, 10 :171–209.

998 Koenen, A. von, 1892. Das norddeutsche Unter-Oligocän und seine Mollusken-Fauna,
 999 4(1). Rissoidae-Littorinidae-Turbinidae-Haliotidae-Fissurellidae-Calyptraeidae-
 1000 Patellidae; (2) Gastropoda Opisthobranchiata; (3)(1) Gastropoda Polyplacophora;
 1001 (2) Scaphopoda; (3) Pteropoda; (4) Cephalopoda. *Abhandlungen zur eologischen*
 1002 *Specialkarte von Preussen und den Thüringischen Staaten* 10(4): 819-1004, pls
 1003 53-62.

1004 Korobkov, I.A. 1966. Krylonogie (Mollusca Pteropoda) paleogenovykh otlozhenij juga
 1005 S.S.S.R. *Voprosy Paleontologii* 5:71-92. (In Russian)

1006 Kunz, A. 1996. *Schalenmorphologische Merkmale der Ontogenesestadien ausgewählter*
 1007 *Pteropoda (Gastropoda: Heterostropha)*. Hamburg, “Diplom-Arbeit im
 1008 Studienfach Geologie-Paläontologie an der Universität Hamburg”, Germany
 1009 (unpublished).

1010 Lamarck, [J.B.P.A. de] 1805-1806. Mémoire sur les fossiles des environs de Paris,
 1011 comprenant la détermination des espèces qui appartiennent aux animaux marins
 1012 sans vertèbres, et dont la plupart sont figurées dans la collection des vélins du
 1013 Muséum, 5. *Annales du Muséum d'Histoire Naturelle de Paris* 5: 28-36, 91-98,
 1014 179-188, 237-245, 349-357.

1015 Lear, C.H., Bailey, T.R., Pearson, P.N., Coxall, H.K. and Rosenthal, Y., 2008. Cooling and
 1016 ice growth across the Eocene-Oligocene transition. *Geology* 36(3): 251–254. doi:
 1017 10.1130/G24584A.1.

1018 Linnaeus, C. 1758. *Systema naturae per regna tria naturae, secundum classes, ordines,*
 1019 *genera, species, cum characteribus, differentiis, synonymis, locis*, 1 (editio
 1020 decima, reformata). Salvii, Holmiae.

1021 Liu, Z., Tou, S., Zhao, Q., Huang, W. 2004. Deep-water Earliest Oligocene Glacial
 1022 Maximum (EOGM) in the South Atlantic. *Chinese Science Bulletin* 49: 2190–
 1023 2197.

1024 Lokho, K. and Kumar, K. 2008. Fossil pteropods (Thecosomata, holoplanktonic
 1025 Mollusca) from the Eocene of Assam–Arakan Basin, northeastern India. *Current*
 1026 *Science* 94(5): 647-652.

1027 Meisenheimer, J. 1905. Pteropoda. *Wissenschaftliche Ergebnisse der Deutschen Tiefsee-*
 1028 *Expedition auf dem Dampfer "Valdivia" 1898-1899*, 9(1):1-314.

1029 Merle, D. 1986. *Contribution à l'étude paléontologique du gisement cuisien de Gan*
 1030 *(Pyrénées-Atlantiques): systématique, évolution et paléoécologie*, 1-2. PhD thesis
 1031 École Pratique des Hautes Études, Dijon, France (unpublished).

1032 Meyer, O. 1886. Contribution to the Eocene paleontology of Alabama and Mississippi, 2.
 1033 *In: Smith, E.A. (ed.). Geology of Alabama. Geological Survey of Alabama,*
 1034 *Bulletin 1: 61-85.*
 1035 Meyer, O. 1887. Beitrag zur Kenntnis der Fauna des Alttertiärs von Mississippi und
 1036 Alabama. *Bericht über die Senckenbergische Naturforschende Gesellschaft* (1886):
 1037 3-22.
 1038 Nicholas, C.J., Pearson, P.N., Bown, P.N., Dunkley Jones, T., Huber, B.T., Karega, A.,
 1039 Lees, J.A., McMillan, I.K., O'Halloran, A., Singano, J.M., Wade, B.S. 2006.
 1040 Stratigraphy and sedimentology of the Upper Cretaceous to Paleogene Kilwa
 1041 Group, southern coastal Tanzania. *Journal of African Earth Sciences* 45: 431–466.
 1042 Orbigny A. d' 1834-1847. *Voyage dans l'Amérique méridionale (le Brésil, la république*
 1043 *orientale de l'Uruguay, la république Argentine, la Patagonie, la république du*
 1044 *Chili, la république de Bolivie, la république du Pérou), exécuté pendant les*
 1045 *années 1826, 1827, 1828, 1829, 1830, 1831, 1832 et 1833*, 5. Bertrand, Paris;
 1046 Levraut, Strasbourg: xliii + 758 pp. (publication dates of separate issues see
 1047 Sherborn and Griffin, 1934). (In French)
 1048 Pearson, P.N., Ditchfield, P.W., Singano, J., Harcourt-Brown, K.G., Nicholas, C.J.,
 1049 Olsson, R.K., Shackleton, N.J. and Hall, M.A. 2001. Warm tropical sea surface
 1050 temperatures in the Late Cretaceous and Eocene epochs. *Nature* 413: 481-487.
 1051 Pearson, P.N., McMillan, I.K., Wade, B.S., Dunkley Jones, T., Coxall, H.K., Bown,
 1052 P.R. & Lear, C.H. 2008. Extinction and environmental change across the Eocene-
 1053 Oligocene. *Geology* 36(2): 179–182. doi: 10.1130/G24308A.1., 279–317

1054 Pearson, P.N., Nicholas, C.J., Singano, J.M., Bown, P.R., Coxall, H.K., van Dongen,
 1055 B.E., Huber, B.T., Karega, A., Lees, J.A., Msaky, E., Pancost, R.D., Pearson, M.,
 1056 Roberts, A.P., 2004. Paleogene and Cretaceous sediment cores from the Kilwa and
 1057 Lindi areas of coastal Tanzania: Tanzania Drilling Project Sites 1–5. *Journal of*
 1058 *African Earth Sciences* 39: 25–62.

1059 Pearson, P.N., Nicholas, C.J., Singano, J.M., Bown, P.R., Coxall, H.K., van Dongen,
 1060 B.E., Huber, B.T., Karega, A., Lees, J.A., MacLeod, K., McMillan, I.K., Pancost,
 1061 R.D., Pearson, M., Msaky, E., 2006. Further Paleogene and Cretaceous sediment
 1062 cores from the Kilwa area of coastal Tanzania: Tanzania Drilling Project Sites 6–
 1063 10. *Journal of African Earth Sciences* 45: 279–317.

1064 Pearson, P.N. and Wade, B. 2015. Systematic taxonomy of exceptionally well-preserved
 1065 planktonic Foraminifera from the Eocene/Oligocene boundary of Tanzania.
 1066 *Cushman Foundation Special Publication* 45: 1–85.

1067 Phipps, C.J. 1774. *A voyage towards the North Pole undertaken by his Majesty's*
 1068 *Command 1773*. Bowyer, W. and Nichols, J., London.

1069 Premoli Silva, I. and Jenkins, D.G. 1993. Decision on the Eocene-Oligocene boundary
 1070 stratotype. *Episodes* 16(3), 379–382.

1071 Rampal, J. 1973. Phylogénie des ptéropodes thécosomes d'après la structure de la
 1072 coquille et la morphologie du manteau. *Comptes Rendus de l'Académie des*
 1073 *Sciences de Paris* 277(D): 1345–1348.

1074 Rafinesque, C.S. 1815. *Analyse de la Nature ou tableau de l'univers et des corps*
 1075 *organisés*. Rafinesque, Palerme: 1–224.

1076 Sherborn, C.D. and Griffin, F.J. 1934. On the dates of publication of the natural history
 1077 portions of Alcide d'Orbigny's "Voyage Amérique Méridionale". *Annals and*
 1078 *Magazine of Natural History*, (10)13(73):130-134.
 1079 Van Dongen, B.E., Talbot, H.M., Schouten, S., Pearson, P.N. and Pancost, R.D. 2006.
 1080 Well preserved Palaeogene and Cretaceous biomarkers from the Kilwa area,
 1081 Tanzania. *Organic Geochemistry* 37: 539–557.
 1082 doi:10.1016/j.orggeochem.2006.01.003
 1083 Wade, B.S., Pearson, P.N. 2008. Planktonic foraminiferal turnover, diversity fluctuations
 1084 and geochemical signals across the Eocene/Oligocene boundary in Tanzania.
 1085 *Marine Micropaleontology* 68: 244–255.
 1086 Wall-Palmer, D., Hart, M.B., Smart, C.W., Sparks, R.S., Friant, A.L., Boudon, G.,
 1087 Deplus, C., Komorowski, J.C., 2012. Pteropods from the Caribbean Sea:
 1088 variations in calcification as an indicator of past ocean carbonate saturation.
 1089 *Biogeosciences*, 9(1): 309-15.
 1090 Winkle Palmer, K. van 1947. The Mollusca of the Jackson Eocene of the Mississippi
 1091 Embayment (Sabine River to the Alabama River), 2. Univalves. *Bulletins of*
 1092 *American Paleontology* 30: 207-563.
 1093 Zilch, A., 1959. Gastropoda, 2. Euthyneura, 1. In: Wenz, W. (ed.). *Handbuch der*
 1094 *Paläozoologie* 6: xii + 200 pp.