



# Permeability and Surface Hardness Surveying of Stone Damaged by Ballistic Impact

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**Abstract:** Recent instances of the destruction of cultural assets in conflict zones have demonstrated the need to develop methods which will allow for the assessment of damage to heritage stone in the field. In particular, non-destructive methods would be invaluable when working on sites damaged by contemporary ballistics. Permeability (TinyPerm 3) and surface hardness (Equotip) surveys of stone damaged by 7.62 × 39 mm (AK-47) projectiles were undertaken to determine the ability of these methods to identify the spatial distribution of damage patterns such as shear faces and surface fractures. Results demonstrate the ability of surface hardness surveys to distinguish between non-impacted surfaces of the target stone and surfaces which shattered/sheared upon impact. Whilst spatial distribution analysis ("heat mapping") of Equotip data did not correlate directly with surface fractures, permeability data heat maps were found to be indicative of surface fracture distribution. The data suggests that compaction of the stone matrix at the impact crater results in a lesser reduction of hardness in this area relative to the wider damaged surface. Surveys of impacted stone using the methods outlined here can identify damage patterns that are not visible to the naked eye, thus aiding in damage identification on fragile sites.

Keywords: heritage; conservation; non-destructive testing; ballistics

# 1. Introduction

Damage to heritage monuments, both deliberate and inadvertent, has been a frequent aspect of armed conflict throughout history [1]. However, recent examples of cultural destruction, such as the demolition of archaeological sites by the Islamic State armed group using explosives [2] and the breaching of the old city walls of Raqqa by coalition airstrikes [3], have brought the issue to prominence in public discourse. Constant improvement in firearms technology means that infantry small arms now also present a significant threat to built heritage. Despite the growing potential for the destruction of these small arms, the damage that ballistic impact and subsequent weathering can cause to buildings is often overlooked in conversations around cultural destruction [4].

The increased capacity for damage that has resulted from the evolution of firearms technology has been explored in a number of studies. Krenn, Kalaus and Hall demonstrated through test-firing early small arms that modern assault rifles have significantly increased accuracy, range and kinetic energy at impact when compared with muskets available in the 16th century [5].

These conclusions are supported by the findings of Mol and Gomez-Heras [6], who investigated the surface hardness of sandstone building blocks impacted by musket balls during the English Civil war, and limestone blocks impacted by modern spitzer shape (pointed) projectiles fired by the Mauser M1893 [7]. The stone impacted by the lower velocity, rounded musket ball was found to have been



compacted at the impact site, raising the surface hardness of the stone in these areas. Conversely, the stone impacted by the Mauser projectiles was found to have reduced surface hardness in the immediate impact sites and the areas surrounding them [6]. Thus, it has been shown that analysis with surface hardness probes can distinguish between impact craters in stone and surrounding non-impacted areas. Therefore, it is likely that surface hardness surveys of impacted stone could compliment other techniques such as permeability analysis in identifying areas of weakness in impacted stone which are most at risk of further degradation.

Miao et al. have shown that an increasing fracture density increases the permeability of a fracture network in stone [8]. Therefore, a permeability survey of the surface of stone impacted by a small arms projectile would likely aid in the identification of surface fracture networks. In addition to contributing to the degradation of the surface of the stone, the presence of surface fractures caused by projectile impact has been shown to be indicative of a wider fracture network in the interior of the impacted material [9]. Thus, a permeability survey could also identify areas of interest on the surface of impacted stone which could then be assessed for interior fracture networks using other techniques.

Some work has been done to explore the damage mechanisms by which high velocity (500–1300 ms<sup>-1</sup>) impacts weaken stone. Dynamic fracture propagation has been demonstrated to be a principle damage mechanism for impacts into brittle materials [10–12]. Fractures exhibit three principle morphologies: radial fractures which propagate outwards from the point of impact, both in a spoke like fashion, and conically into the rock; concentric fractures in a hemispherical orientation around the impact site; and spallation fractures parallel to free surfaces [13–15]. Radial fractures form first, directly behind the initial, quasi-spherical compressive phase of the shock wave. This results in a peak tension parallel to the wave, and thus a radial fracture orientation at 90° to this [14,16]. Concentric fractures form quasi-spherically around the impact as they are the result of the tensional phase of the shockwave, thus peak tension is at 90° to the compressional phase [14]. Finally, spallation fractures form as the compressional phase of the shock wave reflects off a free surface as a tensile wave of equal magnitude, generating a peak tension perpendicular to the free surface [13,16].

Additional research into the propagation of fractures in rock and other brittle materials has allowed the development of numerical modelling of fracture induced damage to rock, stated in the Johnson–Holmquist damage model (J–H) [10]. The model has been refined in successive iterations, and J–H2 numerical modelling has been applied to understand the damage to brittle materials caused by ballistic impact induced fracture networks [17]. However, the J–H2 model does not describe the spatial distribution of fractures within rock, and an understanding of this spatial distribution will be crucial in assessing damage to heritage buildings for the purposes of conservation work [18]. In addition to this, understanding fracture networks will aid in identifying likely points of ingress for weathering agents such as moisture and salt which may further degrade the structure [19,20].

As well as macroscopic fracture damage, the fracturing of constituent grains is a common feature observed through microscale analysis with both electron and optical microscopy techniques. Polanskey and Ahrens described a hemispherical region directly below the impact site with a high density of small fractures and significant grain and pore size reduction through crushing and shear fractures [13]. This zone of compaction and shear fracture is observed in impact experiments into water ice [16], and by using micro-computed tomography ( $\mu$ CT) on stone impacts [6]. Further microstructural observations made by Mol highlight the association of quartz fracturing with proximity to the impact [21]. They present optical microscopy evidence of networks of parallel and high angle micro-fractures transecting grains directly adjacent to the impact site. This is supported by back scatter electron (BSE) imaging of quartz grains from the impact site, depicting a high density of micro fracturing and material loss [6,21].

Previous studies have focused on the effects of low velocity, deformable projectiles such as musket balls and 0.22 lead bullets [4,6]. Having highlighted the use of permeametry and surface hardness in identifying damaged areas of stone, this work will seek to investigate fracture networks arising in sandstone from modern 7.62  $\times$  39 mm ammunition (fired by AK-47 type weapons) using these methods. This research is necessary to understand the damage that modern military projectiles with higher kinetic energies can cause to stone used in built heritage and will inform future conservation strategies. Furthermore, given the difficulty in accessing heritage sites damaged during armed conflict, and the irreplaceable nature of heritage assets, the development of methods which are portable, rapid, and non-destructive for detecting and assessing ballistic damage to stone will be invaluable in ongoing attempts to develop conservation strategies in conflict areas. This work also has the potential to increase the number of techniques available to conservation professionals when assessing ballistic damage to stone, as none of the previous works discussed here have explored the potential of permeametry analysis for this purpose.

## 2. Materials and Methods

### 2.1. Sample Preparation

A sample of well-consolidated mesoporous sandstone (average pore size between 40 and 70  $\mu$ m; water absorption capacity of 1.8%) was sourced from the Huesca region of Spain. The stone was freshly quarried, so as to avoid any pre-existing defects caused by exposure to weathering. The block measured 14.7 × 14.7 × 14.7 cm. The sample was then shot with a 7.6 × 39-mm cartridge fired from an AK-103 assault rifle at a range of 200 m. Figure 1 shows the allocation of side labels on the block, whereas Figure 2a–f shows the individual sides of the block post-impact.



**Figure 1.** A schematic diagram showing the system devised for assigning each sample side a number, and how the sides relate to each other. NB: Side 1 contains the projectile impact crater and sides 2–4 are assigned according to successive 90° clockwise rotations in the same alignment as when the sample was shot. The top of the block is side 5, the bottom of the block is side 6.



Figure 2. Images showing the extent of damage to each side of the sample. Sides are shown 1–6 (**a**–**f**).

On each of the sides of the sample, a grid was created to facilitate accurate sampling using permeametry and Equotip surface hardness readings. The cell size of the sampling grids was defined by the diameter of the Equotip type D impact device  $(2 \times 2 \text{ cm})$ . This cell size allowed for 49 cells on complete faces, whilst those faces which had fractured at impact contained less cells. Grids were

measured using the upper and left-hand edges of each side as a baseline. Each cell was assigned an XY alpha-numeric code, such that the lower left cell was A1 (see Figure 3).



**Figure 3.** Images showing the alphanumeric grid system on an intact side of the sample (side 3, **a**) and a sample significantly fractured at impact (side 4, **b**).

As seen in Figure 3, the fractured nature of the sample meant that some of the surfaces could facilitate less cells/data points than others, as listed in Table 1.

Surface	Number of Cells		
Control	49		
Side 1	22		
Side 2	49		
Side 3	49		
Side 4	30		
Side 5	26		
Side 6	49		
Impact crater/shattered surface	10		

Table 1. A table showing the number of cells surveyed for each surface.

The impact crater of the sample and larger shattered surface could not be incorporated into this grid sampling regime as the surface was too uneven. This also presented difficulties when attempting to obtain Equotip readings, as the instrument requires relatively flat surfaces of at least  $2 \times 2$  cm to obtain a reading. Therefore, ten areas of this shattered surface which were flat enough to obtain a reading were sampled to assess both surface hardness and permeability, and the points which were sampled were numbered and marked on an image of the surface (Figure 4). Each point sampled was measured 3 times and the arithmetic mean was calculated.



**Figure 4.** An image showing the various sampling points on the shattered surface of the sample, and assigned numbers.

## 2.2. Surface Hardness Survey

The surface hardness survey of the sample was conducted using a Proceq Equotip portable Rockwell probe with a D type impact device. The Equotip measures surface hardness by firing a 3-mm tungsten ball at the target surface at a known velocity. The rebound velocity is measured, and the ratio between the initial and rebound velocities is multiplied by 1000 to give a hardness in Leebs (L) [22]. In this study, the Equotip was calibrated on an Equotip hardness test block D with a known hardness of 762 L. Each cell was measured 3 times and the arithmetic mean for that cell was calculated (all data available in supplementary material). All measurements were conducted with the test surface horizontal and the Equotip perpendicular to this. Thus, the influence of gravity on the tungsten ball is systematically applied to all data points, and so can be disregarded from any conclusions [23].

## 2.3. Permeametry Survey

The permeability of the sample's surface was investigated using a New England Research TinyPerm3 air Permeameter. This instrument assesses permeability to air of the stone by creating a vacuum through a piston stroke, drawing air from the sample. The instrument monitors the volume of air withdrawn, and the transient vacuum pulse created at the surface. This data is computed by the instrument and converted into a permeability value in Darcys (D) [24]. The instrument was calibrated using the manufacturer's standards. As with the surface hardness readings, the permeability of each cell was measured 3 times and the mean value was calculated (all data available in supplementary material). It should be noted that as air permeability results presented here serve only to highlight relative differences across the sample, and cannot be used to accurately describe the likely ingress of water as a weathering agent.

#### 2.4. Control and Calibration Measurements

As shown in Figure 3, the optimal distribution of sample cells across the block surfaces left a margin measuring 5–7 mm on the right and bottom edges which was not sampled. This margin, coupled with the calibration for the edge effect detailed below, which was applied to the upper and left edges of each face (cells A1–A7 and A7–G7) ensured that any observed changes to permeability or hardness could be attributed to ballistic damage rather than damage sustained during the cutting process prior to ballistic impacts.

This was necessary as it was hypothesised that the cutting process could potentially exacerbate the edge effect suggested by Viles et al. [26], which may lead to an artificial lowering of surface hardness values. It was also hypothesised in this study that a similar edge effect might be witnessed for air permeametry readings, whereby permeability is increased towards the edge of the sample surface, as this has been discussed in previous works [27].

To ascertain the surface hardness and permeability of non-impacted stone and investigate the presence or absence of an edge effect, a control sample of non-impacted Huesca sandstone was used. Following the alpha-numerical sampling grid seen in Figure 3, 49 cells were sampled across the face of a non-impacted sample measuring  $14 \times 14 \times 4.5$  cm. This would allow for the average permeability and surface hardness of the 24 edge cells to be ascertained and compared with the 25 inner cells (see Figure 5) as well as the averages for the non-impacted stone overall. All statistical significance comparisons between the control and shot sample were undertaken using SPSS Statistics [25] and are contained in the results section.



Figure 5. An image showing the control sample and the distinction between inner and edge cells.

Figure 6 demonstrates that there is a difference between edge cells and inner cells for mean surface hardness. Further statistical analysis using SPSS Statistics 25 and the significance values (p values) for the Shapiro–Wilk tests suggest that both data sets are normally distributed (inner cells: n = 25,  $\bar{x} = 468$  L, p = 0.374, edge cells: n = 24,  $\bar{x} = 405$ , L p = 0.73), whilst the significance value for the 2-tail t-test suggests that there is a significant difference between the mean hardness of the inner cells and edge cells (p = 0.000).

Having established that there is a statistically significant edge effect which lowers surface hardness, the difference in mean surface hardness between the inner cells and the edge cells (63 L) was used as a calibration constant when conducting Inverse Distance Weighted (IDW) heat mapping of surface hardness using ArcGIS Pro. This meant that a 63 L calibration constant was added to the values for the cells abutting the edge of the impacted sample (cells A1–A7 and A7–G7). Thus, any lowering of the surface hardness at the edge of the sample could be attributed to ballistic damage, as was observed by Mol et al. [6], rather than the measured edge effect.

It should be noted that if the standard deviation of the mean of cell readings is compared to the standard deviation of the data set as a whole, there is a difference. The maximum difference between these two standard deviations in surface hardness was on the shattered surface: 14.6 L. For permeability, the largest difference was on side 1: 0.27 D. Using cell average values was necessary to represent spatial data, which is the purpose of this research. Therefore, throughout this research, the standard deviation of the population of cell means for a given surface was used in analysis rather than the standard deviation of the overall data set because this approach maintains a statistical methodology consistent with that used for creating the heat maps. This means that the variability is consistent with the spatial variability of generated heat maps.



**Figure 6.** Mean surface hardness across the control sample, as well as for inner and edge cells of the sample. Error bars are the standard deviation of the sample population.

As discussed earlier, previous studies have stated a discernible edge effect in relation to permeability. This is likely due to the fact that in a free-standing block, the edges of a face act as an interface between two free surfaces, meaning that permeability readings taken at an edge are higher due to air or water flow through the adjacent free face [27]. Supporting this notion, the edge effect was found to be responsible for an increase in permeability towards the control sample's edge (Figure 7), which was also tested for statistical significance. The permeability readings for the inner cells were found to be non-normally distributed using a Shapiro–Wilk test (n = 25, p = 0.000). A Mann–Whitney U test was conducted in SPSS to compare the means of the two sample populations. A Mann–Whitney U test is a non-parametric test which assesses the likelihood that a given data population is significantly different from another. Unlike a t-test for statistical significance, it does not require that the data populations be normally distributed and is therefore appropriate for this data set. The results of this

test suggested a statistically significant increase in average permeability at the edge of the sample (p = 0.000). The difference in mean permeability between the inner cells and the edge cells (inner cells:  $\bar{x} = 0.0072$  D, edge cells:  $\bar{x} = 0.0026$  D, difference: 0.0046 D) was used as a calibration constant when conducting IDW heat mapping of permeability. This means that a 0.0046 D calibration constant was subtracted from the values for the cells abutting the edge of the impacted sample (A1-A7 and A7-G7). Therefore, any increase in permeability at the edge of the sample can be attributed to ballistic damage rather than a potential edge effect.



**Figure 7.** A chart showing mean permeability across the control sample, as well as for inner and edge cells of the sample. Error bars are the standard deviation of the sample population.

## 2.5. Spatial Distribution Analysis (Heat Mapping) Methods

Images of the sample were processed in the ArcGIS Pro Geographic Information Systems software package. The permeametry data was assigned to a point cloud, with each cell shown in Figure 3 given a point with a permeability value in D. Using this point cloud, and the Inverse Distance Weighted (IDW) interpolation spatial analyst tool in ArcGIS Pro, a series of heat maps of likely permeability across the sampling grid was created with an IDW analysis output cell size of  $0.1 \times 0.1$  cm. This was done under The IDW stretch function, with a custom minimum-maximum data set so that the data range for the whole sample (0.0001-5.34 D) could be assigned to a single temperature colour ramp. This colour scheme allowed the visualisation of differences in permeability across individual sample sides, as well as across the sample overall.

Heat mapping of surface hardness data was conducted using the same parameters as those for permeability but using a bathymetric colour ramp which better highlighted local differences in hardness and the range of 275–596 L across the whole sample.

## 3. Results

## 3.1. Surface Hardness Survey Results

As shown in the table associated with Figure 8, the average surface hardness of the six numbered faces (1–6) range from 435 to 475 L. The control sample fits inside this range and, in terms of surface hardness, is consistent with the six numbered sides of the impacted sample. The fact that these sides

have an average hardness comparable to the control sample suggests that surface hardness is not a useful metric when seeking to assess which of the numbered sides of the sample are most damaged by the impact.



Side	Number of Cells	Average Equotip rock surface hardness (L)	Maximum Value (L)	Minimum Value (L)	Range (L)	Standard Deviation (L)
Control	49	438	528	303	225	49.1
All Shot Sample Cells (Excluding Shattered Surface)	225	452	547	275	272	35.9
1	22	461	533	275	258	51.2
2	49	475	547	379	168	29.7
3	49	442	501	374	127	30.7
4	30	447	489	372	116	23.8
5	26	453	533	389	143	29.1
6	49	439	494	318	176	37.1
Shattered Surface Average	10	295	353	225	128	35.6

**Figure 8.** A chart and table showing surface hardness across the faces of the sample. Error bars are the standard deviation of the data for the given surface. NB: Figures in the chart are rounded to three significant figures.

To demonstrate this, the overall hardness data of all 225 cells of the shot sample (excluding the shattered surface) was compared to the control sample data using a Mann–Whitney U test, and found to have no statistically significant difference from the control sample (p = 0.116).

However, the mean hardness for the shattered surface (295 L) was found to be significantly lower than both the mean hardness of the control, and the non-shattered sides of the sample (both tests: Mann–Whitney U test, p = 0.000). Therefore, the Equotip surface survey demonstrates a statistically significant reduction in hardness across the shattered surface.

Furthermore, in Figure 8, side 1 exhibits a larger standard deviation (error bar) than the surfaces which were not directly impacted by the bullet: 51 L for side 1 vs. an average of 30 L for the other numbered sides of the impacted sample. A higher standard deviation of Equotip readings can be indicative of advanced weathering of a rock surface [26,28]. Therefore, the fact that side 1 has the highest standard deviation could suggest that the degradation effects of projectile impact are comparable to those of weathering systems but act instantaneously. However, given that this data set only includes one surface which experienced direct impact (side 1), more investigation is required to ascertain whether the heightened standard deviation of the impacted side of a sample is statistically significant.

Figures 8 and 9 demonstrate that the impact crater has a higher average hardness than the rest of the shattered surface (353 L mean for the impact point vs. 288 L mean for the rest of the shattered surface). This is likely indicative of compaction and realignment of clay minerals within the matrix, as observed using optical microscopy by Mol et al. [4] on samples of the same Huesca sandstone. X-ray diffraction analysis (XRD) conducted by Mol et al. on the Huesca sandstone identifies the relatively high proportion of clay minerals (13% muscovite, 6% kaolinite), which facilitate compaction as a response to impact shock at the impact site [4]. The comparatively lower reduction in surface hardness at the impact site, compared to the wider shattered surface, can be viewed as comparable to the apparent surface hardness increases seen on heritage stone damaged by ballistic impacts observed in the field [6].



**Figure 9.** A chart comparing the hardness of the impact point with the rest of the shattered surface average.

#### 3.2. Permeametry Survey Results

In contrast to the data obtained from the Equotip survey, the permeametry data for all 225 non-shattered cells was found to be significantly different from the control sample (Mann–Whitney U test, p = 0.000). This is reflected in the fact that all surfaces of the impacted sample show a great deal of variation, both in relation to the control and in relation to each other (Figure 10).

Side 1 and the shattered surface show the highest permeability, reflecting the large number of visible fractures across these areas of the sample. The increased permeability seen across side 1 and the shattered surface are due to the fact that the projectile impacted side 1 directly, leading to greater damage in this area and across the adjacent shattered surface.

The sides with the lowest permeability readings were sides 3 and 6. This is likely because side 3 is the antipode of side 1, and thus furthest away from the impact. Therefore, the reduced damage exhibited by side 3 corresponds well with known models of impact damage in rock, in which shock wave energy reduces with distance from the point of impact [29]. Side 6 probably experienced less damage because, at the time of shooting, this side was facing downwards on the target area. This meant that it was the only side with compression/containment, i.e., the mass of the block above it pushing it downward onto the target structure on the firing range.

Permeability readings across a given side of the sample were often found to have ranges spanning orders of magnitude, which would give standard deviation error bars that were so large as to obscure the data itself. For this reason, the standard deviation of the permeability data has been plotted separately in Figure 10 rather than as an error bar, and other statistical information on the data set is listed in the associated table.

It is worth noting in Figure 10 that the sides previously hypothesised to be least damaged by the projectile impact, sides 3 and 6, have the lowest standard deviations. This suggests that, as with the Equotip measurements, the standard deviation of a permeability data set for a given surface of an impacted sample may be related to the damage sustained by it during impact. This is supported by statistical reasoning, as a lower standard deviation would imply a largely homogenous surface without surface fractures which increase the range and standard deviation of the data set for that side. This is indeed the case with side 3. This notion is also supported by the fact that the non-impacted control sample had both the lowest average permeability and the lowest standard deviation.

The hypothesised compaction at the impact site (point 1 of the shattered surface), which is thought to result in a lower reduction of surface hardness, may also be responsible for a reduction in permeability. The shock generated upon impact can realign clay minerals and reduce moisture flow, as discussed by Mol et al. [4]. This phenomenon might be expected to influence our permeability measurements at the impact site. As seen in Figure 11, point 1 does have a lower average permeability than those areas of the shattered surface which show visible fractures or have become friable after shearing (points 3,4,8,9). However, the permeability of the impact point is higher than that for other distal points of the shattered surface (2,5,6,7,10). This is possibly due to the generation of micro-fracture networks through quartz grains, which has previously been observed using optical microscopy [4] and scanning electron microscopy (SEM) [21], creating connected networks that may act as potential pathways for moisture.

The greatly increased permeability of points 8 and 9 are due to sizeable surface fractures close to these points, whilst the intermediate permeability of points 3 and 4 are attributable to flaking of the stone matrix in these areas. Some of these features are observable in Figure 4.



Surface	Number of	Average	Maximum	Minimum	Range (D)	Standard
	Cells	Permeability	Value (D)	Value (D)		Deviation (D)
		(D)				
Control	49	0.00496	0.01643	0.00023	0.0162	0.0047
All Shot Cells	225	0.134	5.33	0.00030	5.33	0.604
(Excluding						
Shattered						
Surface)						
1	22	0.454	2.93	0.000564	2.932	0.892
2	49	0.171	5.33	0.00030	5.33	0.829
3	49	0.0164	0.0555	0.00143	0.054	0.0153
4	30	0.0929	1.6	0.00065	1.599	0.314
5	26	0.203	4.86	0.00041	4.86	0.951
6	49	0.0600	1.60000	0.00078	1.599	0.261
Shattered	10	0.532	2.033	0.01487	2.018	0.699
Surface Average						

**Figure 10.** A chart and table showing the variation in average permeability and standard deviation of permeability readings across the sample surfaces.



**Figure 11.** A chart showing the variation in permeability across the impact crater and wider shattered surface. Error bars are the standard deviation of the data set for each sample area.

## 3.3. Spatial Distribution Analysis ("Heat Mapping") of Sample Permeability Data

In addition to distinguishing between the levels of damage across the sample surfaces, permeametry is effective at identifying the areas of a given surface that have been significantly fractured as a result of projectile impact. This is best demonstrated by presenting the permeametry data in relation to its spatial distribution across the sample surface, as seen in Figure 12. Fractures have been highlighted in black. With the exception of side 1, which was impacted face on (Indicated by a red ' $\otimes$ '), the direction of projectile impact is indicated with a red arrow. As discussed previously, a calibration constant of -0.0046 D was applied to the cells contacting the sample edge, so as to remove any edge effect when visualising ballistic damage at the sample's edge.

A number of important trends become evident from the visualisation of the spatial distribution of the permeability data. Most obvious of these is that for every surface displaying a fracture network, the areas of highest permeability correspond directly to the position of the fractures. This is perhaps an obvious observation, as a fracture will drastically increase the permeability of the local surface it runs through. However, this clearly demonstrates the efficacy of permeametry analysis in identifying the position of surface fractures in an impacted sample. Some areas with surface fractures appear to have higher permeability that other areas exhibiting surface fracture. This is possibly due to differences in the sub-surface extent of the fractures, i.e., that those fractures which extend further into the sample or connect with a fracture network of greater density will give higher permeability readings, but further work will be necessary to confirm this [8].

The area of increased permeability to the upper left corner of side 3 of the sample is notable and although it is markedly lower than the permeability averages for surface fractures on other sides (the highest reading for a side 3 cell is 0.056 D) it could be indicative of sub-surface micro-fractures or defects. This could be the result of natural variation in the stone or could be related to residual shock wave damage towards the rear of the sample. Further work would be necessary to determine the cause of this area of increased permeability.



0.0001 D

Figure 12. A series of Inverse Distance Weighted (IDW) permeability heat maps generated for the impacted sample (sides 1–6, **a**–**f**).

# 3.4. Spatial Distribution Analysis ("Heat Mapping") of Sample Equotip Surface Hardness Data

As can be seen in Figure 13, spatial analysis of the distribution of Equotip surface hardness readings corresponded less directly with identifiable surface fractures than similar analysis for permeability data. This can be seen in Figure 13 and although there is some agreement between those areas exhibiting surface fractures and areas of lower hardness, such as those seen in the heat map for side 6 in Figure 13, other areas of lower hardness appear to be unrelated to the presence of surface fractures, such as sides 2 and 4.



**Figure 13.** A series of IDW Equotip surface hardness heat maps generated for the impacted sample (sides 1–6, **a**–**f**).

Furthermore, some large fractures visible on the sample surface do not affect the hardness of the surrounding cell, as is the case for the heat map of side 1, and are in direct contrast to the permeability heat map for that side. The addition of the 63 L calibration constant appears to have been successful in creating a broadly homogenous surface hardness independent of the edge effect for many of the sample sides, as seen on sides 3,4 and 6. However, the left and upper edges for side 2 do appear to be significantly harder than the rest of that surface, indicating that further work may be required to thoroughly investigate the edge effect and to calibrate for its effects when investigating ballistic damage in stone.

# 4. Discussion

As shown in Figure 4, there is a discoloured area of lighter stone in the centre of the shattered surface. This is the impact point of the projectile (point 1 on the damaged surface) and as shown in Figure 9, it has a noticeably higher average hardness than the rest of the shattered surface (353 L for the impact point vs. 288 L average for the rest of the shattered surface). The fact that the impact crater has experienced less weakening than the rest of the shattered surface could indicate compaction of the stone matrix at the impact point, as found in previous studies [4,6]. Although the impact crater has a reduced hardness when compared with the cells sampled on non-shattered sides (353 L compared with an average of 452 L for non-shattered cells), the fact that it is harder than the surrounding shattered stone is an important finding when seeking to model impact damage to stone. It is notable because it suggests that the most damaged areas may not be the point of direct impact, but rather the wider shattered surface which has sheared away during the ballistic event.

The variability in hardness readings from side 1 is higher than those from other sides of the block, with a range of 258 L compared to an average range of 146 L for the other sides, whilst side 1 also has an increased standard deviation (Figure 8). This is demonstrated by the charts presented here, which show that Equotip surface hardness is a highly indicative method for quantifying the difference between the point of impact of a ballistic projectile and non-impacted areas. Equotip is also effective at identifying areas of wider damage across any shattered/shear faces arising from such an impact. This trend is indicated both by the Equotip absolute values and the standard deviation of the data sets for a given side, which may indicate that ballistic impact has the capacity to degrade stone to a comparable extent to known weathering processes, but much more rapidly [26,30]. More studies will be necessary to confirm that the link between point of impact and high standard deviation is statistically significant.

However, the Equotip surface hardness survey is likely not useful in detecting surface fractures on an otherwise undamaged surface. This can be seen in the heat maps presented in Figure 13 and although there is some agreement between surface fractures and areas of reduced surface hardness (side 1), for many surface fractures, there is no corresponding reduction in surface hardness (sides 2 and 4). It is possible that fractures resulting in a discernible reduction of hardness in the stone have wider micro-fracture networks associated with them. Investigations using Schmidt hammer by Young and Frowell [31] evidenced a decrease in surface hardness with smaller discontinuity spacing. Decreased surface hardness proximal to a visible fracture may therefore indicate a network of micro-fractures associated with it that are unobservable on the surface.

An alternative explanation may not stem from the spacing and number of fractures, but from the orientation of those fractures relative to the measured face. Rebound values are strongly reduced when discontinuities, such as internal layering, fracturing or mineral alignments are oriented perpendicular to the direction of Equotip impact [23]. It could therefore be that the noticeable decrease in hardness may result from a fracture intersecting the measured surface at a low angle. Fractures that display no noticeable change in hardness may be the result of fractures intersecting the surface at a high angle, or a discrete fracture with no associated micro-fractures. Further investigation through destructive methods such as thin sections may be necessary to quantify this.

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Mirroring the Equotip findings, permeability at the impact crater was found to be lower than the average for the wider shattered surface, which further supports the notion that there is compaction at this point which causes a smaller reduction in hardness and a smaller increase in permeability. However, whilst the impact crater has the highest hardness across the shattered surface when measured with the Equotip (Figure 9), it does not have the lowest permeability. This can be seen in Figure 11 and shows that many of the sampling points further from the impact crater have a lower permeability than the centre (points 2,5,6,7,10).

The fact that the impact crater has a slightly increased permeability relative to these points is possibly due to the fracturing of quartz grains at the impact site. The fracturing of quartz grains at the site of ballistic damage has been observed under scanning electron microscope (SEM) by Mol [21]. The increase in permeability at the point of impact is evidence of the same phenomenon being present with larger calibre projectile impacts and highlights that ballistic damage to stone causes both macroscopic (observable surface fractures) and microscopic changes (the fracturing of quartz grains). It is also possible that the discoloration in this area (Figure 4) is related to the change in refractive index observed in microscopic analyses of shocked quartz grains arising from shock waves associated with high-energy impacts [32]. Identifying areas of such discoloration may prove a useful diagnostic tool in identifying areas of impact in the field.

Another clearly identifiable spatial trend is that permeability, which can be taken to infer damage to the sample, correlates strongly with the direction of impact. In the heat map for side 1, the fractures and areas of increased permeability can be seen to radiate outwards from the point of impact. Radial fracturing is a classic sign of ballistic impact in brittle materials [33] and shows that permeametry analysis is accurately reflecting known projectile damage patterns. In sides 2, 4, 5 and 6, it is clear that the side of the surface closest the impact (indicated by the red arrow) experiences a far greater increase in permeability than the side furthest from the impact. This trend is especially clear in the heat map for side 2, where the direction of impact and the surface fractures are strongly reflected in the areas of highest permeability across the surface.

In the sides that did not experience direct impact (2,4,5,6), the fractures often run parallel to the direction of impact, and this is especially true on sides 2 and 6. However, some surface fractures clearly run perpendicular to the direction of projectile impact, as can be seen on sides 4 and 5. This is perhaps evidence that the shockwaves generated by the projectile impact are exploiting pre-existing defects in the stone which result in these fracture patterns. This would support previous findings which illustrated through computed tomography analysis that internal damage to the stone arising from bullet impact follows bedding planes within the sample [4]. However, fractures perpendicular to the direction of impact could also be the expression of the hemispherical concentric fractures described earlier [13,14]. Further work utilising thin section analysis or computed tomography imaging may clarify this.

In areas distant from the surface fractures, the permeability heat maps are broadly homogenous, indicating that there has been no notable artificial reduction in permeability due to the calibration constant. Therefore, we can consider this a viable method for removing the edge effect when seeking to visualise increases in permeability due to ballistic damage.

IDW heat maps show that permeability surveys of impacted stone can help in identifying and visualising the spatial distribution of surface fractures, which in turn could inform conservation strategies by identifying likely points of ingress for weathering agents. Furthermore, permeability readings and their standard deviations indicate that the technique is effective at highlighting the sides of a sample block which are most and least damaged.

Given that both the instruments used in this work (TinyPerm 3 and Equotip) are portable, these findings demonstrate a potential new method for conservation professionals to assess ballistic damage to heritage stone in the field. They could therefore provide a new capability to conduct survey work rapidly on site, which is a key requirement given the time constraints often inherent to accessing ballistic damage to heritage sites during unpredictable conflicts.

#### 5. Conclusions

The data presented here supports previous work on ballistic damage to stone, confirming previous observations that ballistic impact can lead to higher Equotip surface hardness for the impact crater relative to some areas of the surrounding stone [4], whilst also demonstrating that these conclusions hold true for higher calibre, more damaging ammunition than previously studied. Furthermore, the significantly lower hardness values of the shattered surface compared with the non-shattered cells illustrate the efficacy of using Equotip surface hardness to identify those surfaces which have been sheared or shattered as a result of ballistic impact. Therefore, the Equotip surface hardness survey has been shown here to be a useful method in distinguishing between less-damaged sides and the impact crater and wider shattered surfaces arising from ballistic impact.

On sides distant from the impacted side of the sample, permeametry surveys have been shown in this work to be an effective method for identifying surface fracture networks and identifying which specific surfaces/areas of an impacted stone sample are most damaged. This will aid in conservation work by allowing the triaging of most damaged building materials and identifying points of ingress for weathering agents.

Statistical analysis undertaken here supports the likely presence of an edge effect for both Equotip surface hardness and permeability for these samples of cut stone. Whilst Viles et al. did not record this edge effect on natural stone blocks [26], this does appear to be an issue for blocks subjected to the strain of cutting equipment. This is worth considering when informing conservation strategies for ballistic damage to stone, as impacts which cause damage and fracturing to the edge of a block may be exacerbated by the edge effect. However, further work is necessary to properly understand the edge effect, both to confirm its presence for cut stone, and to understand how to effectively calibrate for it during ballistic investigations.

The principle findings of this paper are that an integrated dual survey using surface hardness and permeability greatly increases the potential for informative data. Critically, these methods can be utilised rapidly in the field and are non-destructive, pointing to a viable method for assessing damage to built heritage in conflict zones.

This work has highlighted many areas for further investigation, particularly using those methods which will allow insight into the interior of the impacted sample. Thin section microscopic analysis and X-ray computed tomography analysis might allow for the relationship between fracture length and density and corresponding permeability and hardness values to be explored. Scanning electron microscopy could also be used to explore the theorised effects of matrix compaction and grain fracturing at the impact site.

**Supplementary Materials:** The following are available online at http://www.mdpi.com/2571-9408/2/2/87/s1, Spread sheet 1: "Permeability Sample Data and Control Data", Spread sheet 2: "Surface Hardness Sample Data and Control Data".

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#### References

- 1. Mollick, J. *The Fate of Cultural Property in Wartime: Why it Matters and What Should Be Done;* Carnegie Council for Ethics in International Affairs: New York, NY, USA, 2013.
- Harmansah, O. ISIS, heritage, and the spectacles of destruction in the global media. *Near East. Archaeol.* 2015, 78, 170–177. [CrossRef]

- 3. Asor Cultural Heritage Initiatives. Damage to al-Rafiqah Wall in Raqqa's Old City by US-Led Coalition Forces. Available online: https://www.asor-syrianheritage.org/update-damage-to-al-rafiqah-wall-in-raqqas-old-city-by-us-led-coalition-forces/ (accessed on 1 February 2019).
- 4. Mol, L.; Gomez-Heras, M.; Brassey, C.; Green, O.; Blenkinsop, T. The benefit of a tough skin: Bullet holes, weathering and the preservation of heritage. *R. Soc. Open Sci.* **2017**, *4*, 160335. [CrossRef]
- 5. Krenn, P.; Kalaus, P.; Hall, B. Material culture and military history: Test-firing early modern small arms. *Mater. Cult. Rev.* **1995**, *42*, 101–109.
- 6. Mol, L.; Gomez-Heras, L. Bullet impacts and built heritage damage 1640–1939. *Herit. Sci.* **2018**, *6*, 35. [CrossRef]
- 7. Grant, N. Mauser Military Rifles; Bloomsbury Publishing: London, UK, 2015; pp. 27–28.
- 8. Miao, T.; Yu, B.; Duan, Y.; Fang, Q. A fractal analysis of permeability for fractured rocks. *Int. J. Heat Mass Transf.* **2015**, *81*, 75–80. [CrossRef]
- 9. Ai, H.; Ahrens, T. Dynamic tensile strength of terrestrial rocks and application to impact cratering. *Meteorit. Planet. Sci.* **2014**, *39*, 233–246. [CrossRef]
- 10. Ma, G.; An, X. Numerical Simulation of blasting-induced rock fractures. *Int. J. Rock Mech. Min. Sci.* 2008, 45, 966–975. [CrossRef]
- 11. Krishnan, K.; Sockalingam, S.; Bansal, S.; Rajan, S. Numerical simulation of ceramic composite armor subjected to ballistic impact. *Compos. Part B Eng.* **2010**, *41*, 583–593. [CrossRef]
- 12. Neckel, L.; Hotza, D.; Stainer, D.; Janssen, R.; Al-Qureshi, H. Modelling of ballistic impact over a ceramic-metal protection system. *Adv. Mater. Sci. Eng.* **2013**, *2013*, 1–8. [CrossRef]
- Polanskey, C.A.; Ahrens, T.J. Impact spallation experiments: Fracture patterns and spall velocities. *Icarus* 1990, 87, 140–155. [CrossRef]
- 14. Ahrens, T.J.; Rubin, A.M. Impact-induced tensional failure in rock. *J. Geophys. Res.* **1993**, *98*, 1185–1203. [CrossRef]
- 15. Kumar, P.S.; Kring, D.A. Impact fracturing and structural modification of sedimentary rocks at Meteor Crater, Arizona. *J. Geophys. Res.* **2008**, *113*, E9. [CrossRef]
- 16. Arakawa, M.; Shirai, K.; Kato, M. Shock wave and fracture propagation in water ice by high velocity impact. *Geophys. Res. Lett.* **2000**, *27*, 305–308. [CrossRef]
- 17. Oucif, C.; Mauludin, L. Numerical modelling of high velocity impact applied to reinforced concrete panel. *Undergr. Space* **2018**, *4*, 1–9.
- 18. Warke, P.; Curran, J.; Turkington, A.; Smith, B. Condition assessment for building stone conservation: A staging system approach. *Build. Environ.* **2003**, *38*, 1113–1123. [CrossRef]
- Sousa, L.M.; del Rio, L.M.S.; Celleja, L.; de Argandona, V.G.R.; Rey, A.R. Influence of microfractures and porosity on the physico-mechanical properties and weathering of ornamental granites. *Eng. Geol.* 2005, 77, 153–168. [CrossRef]
- 20. Benavente, D. Why pore size is important in the deterioration of porous stones used in the built heritage. *Macla Mag. Span. Soc. Miner.* **2011**, *15*, 41–42.
- 21. Mol, L. Armed conflict impacts on the microscale. J. Phys. Conf. Ser. 2017, 902, 012032. [CrossRef]
- 22. Kompatscher, M. Equotip-rebound hardness testing after D. Leeb. In Proceedings of the Conference on Hardness Measurements Theory and Application in Laboratories and Industries, Washington, DC, USA, 11–12 November 2004.
- 23. Aydin, A. ISRM Suggested method for determination of the Schmidt hammer rebound hardness: Revised version. *Int. J. Rock Mech. Min. Sci.* **2009**, *46*, 627–634. [CrossRef]
- 24. Filomena, C.; Hornung, J.; Stollhofen, H. Assessing accuracy of gas-driven permeability measurements: A comparative study of diverse Hassler-cell and probe permeameter devices. *Solid Earth* **2014**, *5*, 1–11. [CrossRef]
- 25. McPhee, C.; Arthur, K. Klinkenberg permeability measurements: Problems and practical solutions. In Proceedings of the Second Society of Core Analysts European Core Analysis Symposium, London, UK, 20–22 May 1991.
- Viles, H.; Goudie, A.; Grab, S.; Lalley, J. The use of the schmidt hammer and equotip for rock hardness assessment in geomorphology and heritage science: A comparative analysis. *Earth Surf. Process. Landf.* 2011, 36, 320–333. [CrossRef]

- 27. Turkington, A.V.; Smith, B.J. Observations of three-dimensional salt distribution in building sandstone. *Earth Surf. Process. Landf.* 2000, 25, 1317–1332. [CrossRef]
- 28. Mol, L.; Viles, H. The role of rock surface hardness and internal moisture in tafoni development in sandstone. *Earth Surf. Process. Landf.* **2012**, *37*, 301–314. [CrossRef]
- 29. Thoma, K.; Hornemann, U.; Sauer, M.; Schneider, E. Shock waves—Phenomenology, experimental, and numerical simulation. *Meteorit. Planet. Sci.* 2005, 40, 1283–1298. [CrossRef]
- 30. Hall, K.; Arocena, J.; Boelhouwers, J.; Liping, Z. The influence of aspect on the biological weathering of granites: Observations from the Kunlun Mountains, China. *Geomorphology* **2005**, *67*, 171–188. [CrossRef]
- 31. Young, R.P.; Frowell, R.J. Assessing rock discontinuities. Tunn. Tunn. Int. 1974, 10, 45–48.
- 32. Grieve, R.; Langenhorst, F.; Stoffler, D. Shock metamorphism of quartz in nature and experiment: II. Significance in geoscience. *Meteorit. Planet. Sci.* **1996**, *31*, 6–45. [CrossRef]
- 33. Johnson, G.; Beissel, S.; Holmquist, T.; Frew, D. Computed radial stresses in a concrete target penetrated by a steel projectile. *Trans. Built Environ.* **1998**, *32*, 14.



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