

## **Diversity and distribution of methane-oxidizing microbial communities associated with different faunal assemblages in a giant pockmark of the Gabon continental margin**

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### **Abstract:**

A giant 800-m-diameter pockmark named REGAB was discovered on the Gabon continental margin actively emitting methane at a water depth of 3200 m. The microbial diversity in sediments from four different assemblages of chemosynthetic organisms, Mytilidae, Vesicomidae, Siboglinidae and a bacterial mat, was investigated using comparative 16S rRNA gene sequence analysis. Aggregates of anaerobic methanotrophic archaea (ANME-2) and bacteria of the *Desulfosarcina/Desulfococcus* cluster were found in all four chemosynthetic habitats. Fluorescence in situ hybridization targeting the ANME-2/*Desulfosarcina/Desulfococcus* aggregates showed their presence few centimeters (3–5 cm) below the surface of sediment. 16S rRNA gene sequences from all known marine ANME groups were detected in the pockmark sediments, as well as from both known bacterial partners. The archaeal diversity was limited to the ANME cluster for all investigated samples. The bacterial diversity included members of the *Proteobacteria*, Bacilliales, Cytophaga/Flavobacteria, *Verrucomicrobia*, JS1 and *Actinobacteria* clusters. Bacterial 16S rRNA gene sequences related to those of known sulphide-oxidizing symbionts were recovered from tissues of several invertebrates including vesicomid clams and siboglinid tubeworms of REGAB.

**Keywords:** REGAB; AOM; Chemosynthetic ecosystems; Cold seep; Faunal assemblage; Microbial phylogeny

## 1. Introduction

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Deep-sea cold-seep ecosystems are characterized by high fluxes of hydrocarbons and sulphide originating from the subsurface seabed by thermogenic or microbial production. Chemotrophic microorganisms, either free-living or in symbiosis with invertebrates, are abundant at these sites where they derive energy from the oxidation of sulphide or methane (Paull et al., 1984; Sibuet et Olu, 1998; Pancost et al., 2000; Valentine et al., 2000; Orphan et al., 2001; Tunnicliffe et al., 2003). Deep-water cold seeps often show substantial accumulations of methane in the seafloor in gaseous, liquid and solid form, the latter also known as gas hydrates (Kvenvolden, 1998; Milkov and Sassen, 2002). A major fraction of the methane rising from deeper sediments to the seafloor is consumed by microorganisms within the seabed, mainly through the anaerobic oxidation of methane (Reeburgh, 1976). The anaerobic oxidation of methane (AOM) by sulphate reduction leads to the formation of sulphide and carbonate (Barnes and Goldberg, 1976; Reeburgh, 1976). This process is mediated by yet uncultivated methanotrophic archaea related to the methanogens, and sulphate reducers of the Deltaproteobacteria clade (Hinrichs and Boetius 2002). The sulphate-reducing bacteria (SRB) and anaerobic methanotrophic archaea (ANME) often form aggregates and occur at very high abundances in cold-seep sediments (Boetius et al., 2000; Orphan et al. 2001). Several distinct phylogenetic clusters of *Archaea* are involved in AOM as shown in various observations including proteomic and metagenomic data (Knittel et al. 2005, Krüger et al. 2003; Hallam et al. 2004; Meyerdierks et al., 2005). The ANME-1 group is related to the *Methanosarcinales* and *Methanomicrobiales* orders (Hinrichs et al. 1999), the ANME-2 and ANME-3 groups are related to the *Methanosarcinales* (Boetius et al, 2000; Niemann et al., 2006). The ANME occur in sediments in consortium with sulphate-reducing bacteria (Michaelis et al, 2005; Knittel et al, 2005) or as single cells or small assemblages (Orphan et al., 2001; Knittel et al., 2005). ANME-1 and ANME-2 are generally associated with sulphate-reducing bacteria of the *Desulfosarcina/Desulfococcus* cluster, and ANME-3 with the *Desulfobulbus* genus (Lösekann et al., 2007). The AOM process provides high fluxes of sulfide, which nourishes several types of microbe-invertebrate symbioses such as bivalves from the Lucinidae, Thyasiridae, Solemyidae and Vesicomidae families, as well as annelid polychaetes of the family Siboglinidae (Goffredi et al., 2004; Duperron et al., 2007a; Bright and Giere, 2005). Mussels from cold seeps (family Mytilidae) host methanotrophic bacteria (Cavanaugh et al., 1987; Barry et al., 2002) or live in dual symbiosis with both methanotrophic and thiotrophic bacteria (Won et al., 2003; Duperron et al., 2005; Duperron et al., 2007b). In most cold-seep environments, such chemosynthetic organisms dominate the community biomass and form specific habitats generally associated with high AOM rates. Pockmarks are a special type of cold-seep characterized by negative seafloor features caused by the seepage of fluids through the seabed (Hovland and Judd, 1988). They generally occur in unconsolidated fine-grained sediments as cone-shaped circular or elliptical depressions ranging from a few metres to 300 or more in diameter and from 1 m to 80 m in depth. The giant 800 m diameter pockmark REGAB was discovered along the Gabon continental margin at a water depth of 3200 m during the ZAIANGO cruise by geophysical surveys with the research vessel Suroit (April 2000, Ifremer). REGAB is characterized by high methane emissions and by the presence of methane hydrates in subsurface and surface sediments (Charlou et al., 2004). First dives on REGAB using the remotely operated vehicle (ROV) VICTOR 6000 (cruises ZAIANGO and BIOZAIRE I - 2001) showed the occurrence of dense communities of typical seep-associated invertebrates, some of which represent chemosynthetic symbioses (Ondréas et al., 2005; Olu-Le Roy et al., 2007a). Among these were Mytilidae (*Bathymodiolus* aff. *boomerang*, Olu-Le Roy et al. 2007b), Vesicomidae bivalves (see Cosel and Olu-Le Roy, this volume), and Siboglinidae polychaetes (Vestimentifera, *Escarpia southwardae* n. sp., Andersen et al., 2005). A few mats of giant sulfide oxidizing bacteria have also been observed in the pockmark. Three main faunal assemblages were observed (Figure 1): Mytilidae and Siboglinidae occur in the centre area of the pockmark associated with carbonate crusts, and Vesicomidae mostly in

the periphery on soft sediments. The bacterial mats patches were observed close to Vesicomidae. Olu et al. (2007a) found that methane concentrations in the bottom water were the main factor explaining the distribution and high spatial variability of the faunal assemblages. Methane measurements in these habitats suggested that the *Bathymodiolus* species are exposed to highest methane concentrations while Vesicomidae and Siboglinidae occur at the periphery of methane-emitting patches, experiencing lower methane concentrations. The occurrence of *Bathymodiolus* mussels in very large and dense aggregates in the most methane-rich areas is probably linked to their dual symbiosis with methane and sulphide-oxidizing Gammaproteobacteria (Duperron et al. 2005). Vesicomidae and Siboglinidae were assumed to harbour only sulphide-oxidizers (Olu-Le Roy et al., 2007a).

In this study, we present a molecular survey of microbial communities associated with different types of habitats in the REGAB giant pockmark. The two main goals were to assess microbial diversity associated with the different faunal assemblages including their bacterial symbionts, and to assess the distribution of microorganisms involved in AOM in the seafloor.

## 2. Material and methods

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### 2.1. Sample collections

During the BIOZAIRE II cruise in November 2001, aboard the research vessel L'Atalante, sediment samples were collected with push cores operated by the ROV VICTOR 6000 (Figure 1). The target sites in the REGAB pockmark at 3152 m depth included sites M2 and M3 close to Mytilidae (dives 146 and 147, cores CT15 and CT2 respectively), sites V1 and V3 close to Vesicomidae (dives 145 and 147, cores CT1 and CT8 respectively), site Vest close to Siboglinidae (dive 146, core CT4) and site "Mat" on bacterial mat (dive 146, core CT10) (Table 1). Lengths of sediment cores were 15 cm except for M2, V3 and Vest sites because of carbonate concretions below 8, 10 and 10 cm sediment depth, respectively. The site nomenclature follows the one that had been used before (Olu-Le Roy et al., 2007a; Duperron et al., 2007). Description of the faunal communities associated with the sites can be found in Olu-Le Roy et al., 2007a. Siboglinidae, Mytilidae and Vesicomidae were collected by the ROV grab and rapidly dissected on board. Bivalve gills and siboglinid trophosomes were frozen for phylogenetic analysis of bacterial symbionts.

### 2.2. Methane concentration measurement

Fluids were sampled using the water sampling pump and collected in 200 ml titanium bottles at the different sites (Table 1). Methane concentrations were measured as described previously (Olu et al., 2007a).

### 2.3. Fluorescence in situ hybridization (FISH)

Subsamples of sediment cores were sliced into 1 cm intervals and fixed for 2h with 2% formaldehyde solution. After washing with PBS (10 mM sodium phosphate, 130 mM NaCl), samples were stored in PBS/ethanol (1/1) at -20°C until used in the laboratory. Fixed sediment samples were diluted ten times in PBS/ethanol and sonicated 20s. Twenty µL were filtered on 0.2 µm GTTP polycarbonate filters. Separated hybridizations with two monolabelled probes were done on half cut filters. Microbial aggregates were counted after hybridization with Cy3 monolabelled probes, probe EelMS932 (5'-AGCTCCACCCGTTGTAGT-3') for *archaea* ANME-2 and probe DSS658 (5'-TCCACTTCCCTCTCCCAT-3') for sulfate-reducing bacteria (*Desulfosarcina* /*Desulfococcus*). After staining with 4,6-diamidino-2-phenylindole (DAPI), aggregates were counted with the appropriate fluorescent filters on a Olympus BH-2 microscope. For each

sample slide, 500 microscope fields were counted and reported as number of aggregates per ml.

#### **2.4. DNA extraction**

DNA from the different samples (0.8 g per sample) was extracted with FastDNA SPIN kit for soil (Bio 101 System, Qbiogen) with the modifications previously described by Webster et al. (2003). The DNA extraction and amplification procedures gave good results on all samples except the bacterial mat sample. Therefore, the DNA extraction of the bacterial mat sample (dive 146 core CT10) was carried out on 5 x 0.8 g replicates. The replicates were pooled, concentrated and purified with YM-100 centrifugal filters (Millipore Corporation) to give a final volume of 70  $\mu$ L.

#### **2.5. PCR and cloning**

For cores CT1 and CT2, amplifications were performed using the 16S rRNA gene universal primers for Bacteria or Archaea: U1492r (5'-GTT ACC TTG TTA CGA CTT-3') as reverse universal primer, E8f (5'-AGA GTT TGA TCA TGG CTC AG-3') for Bacteria and A8f (5'-CGG TGG ATC CTG CCG GA-3') for Archaea. The amplifications for cores CT4, CT8, CT15 and Bacteria in core CT10 were performed with: U1407r (5'-GAC GGG CGG TGW GTR CAA-3') as reverse universal primer, E338f (5'-ACT CCT ACG GGA GGC AGC-3') for Bacteria and E344f (5'-AYG GGG YGC ASC AGG SG-3') for Archaea.

For amplification of Archaea in core CT10, nested PCR was conducted with combination of 16S rRNA gene primers for Archaea with 8f (5'-CGG TTG ATC CTG CCG GA-3') and 1492r (5'-GGC TAC CTT GTT ACG ACT T-3') in the first round, with 344f (5'-AYG GGG YGC ASC AGG SG-3') and 915r (5'-GTG CTC CCC CGC CAA TTC CT-3') in the second round. Five independent PCR products from the first round were pooled and purified (QIAquick PCR purification Kit, Qiagen) and used as a template for the second PCR round.

The bulk DNA was amplified in a 50  $\mu$ l reaction mix containing (final concentration): 1X Taq DNA polymerase buffer (Q biogen, Strasbourg, France), 20  $\mu$ M of each dNTP, 20  $\mu$ M of each primer and 0.5  $\mu$ l of Taq DNA polymerase (Q Biogen, Strasbourg, France). Reaction mixtures were held at 94°C for 3 min followed by 30 cycles of 94 °C for 60 s, 49 °C (except for primer couples U1407r-E338f, A1492r-A344f and A915r-A344f: 54°C, 55°C and 57°C respectively) for 90 s and 72 °C for 120 s, with a final extension step of 6 min at 72 °C. PCR products were then visualized on a ethidium bromide containing agarose gel before cloning. Before cloning, two to four PCR products were pooled, purified and concentrated using the QIAquick PCR purification kit (Quiagen SA, Grenoble, France) following the manufacturer's instructions. Clone libraries were constructed by transforming *E. coli* TOP10F' using the TOPO TA Cloning kit (Invitrogen Corp., San Diego CA USA) according to the manufacturer's protocol. Clones harbouring the expected insert were then cultured and treated for sequencing on the "Ouest Genopole Plateform" (Roscoff, France, <http://www.sb-roscoff.fr/SG/>) on a Abi prism 3100 GA (Applied Biosystem), using the big-Dye Terminator V3.1 (Applied Biosystem) following the manufacturer's instructions.

#### **2.6. Phylogenetic analyses**

To determine preliminary phylogenetic affiliations, sequences were compared to those available in databases using the BLAST network service (Altschul et al., 1990). Alignments of 16S rRNA gene sequences were performed using CLUSTALW (Thompson et al., 1994), further refined manually using SEAVIEW (Galtier et al., 1996). The trees were constructed by PHYLO-WIN (Galtier et al., 1996). Only homologous positions were included in the phylogenetic comparisons. For the 16S rRNA phylogenetic reconstruction, the robustness of inferred topology was tested by bootstrap resampling (1000) (Felsenstein, 1985) of the tree calculated on the basis of evolutionary distance (neighbour-joining algorithm with Kimura 2 correction). For the Gamma and and Deltaproteobacteria phylogenetic reconstruction, the

Maximum Likelihood was used with 100 bootstraps resampling. Sequences displaying more than 97% similarity were considered to be related, and grouped in the same phylotype.

## 2.7. Nucleotide sequence accession numbers.

Sequences have been deposited at EMBL for partial 16S rRNA gene sequences. The clones were named according to the sample core number (CTi), and A for *Archaea* and B for *Bacteria*. For the symbionts, the entire name of the host was used. The accession numbers are AM888202 to AM888272.

## 3. Results and discussion

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### 3.1. Distribution of AOM consortia

Fluorescence in situ hybridization of pockmark sediments showed abundant microbial aggregates targeted with ANME-2 and DSS probes. All the counted aggregates were also visualized with DAPI staining. The aggregates consisted of ANME-2 archaea surrounded by sulphate-reducing bacteria of the *Desulfosarcina/Desulfococcus* cluster (Figure 2). No monospecies aggregates of ANME-2 were found as previously detected in cold seep sediments from Eel River Basin (Orphan et al., 2002). Depth profiles of the number of AOM consortia differed between the habitats (Figure 3). Sediment from the Mytilidae site M3 contained the highest number of aggregates ( $1.8 \times 10^7 \text{ cm}^{-2}$ ) integrated over 15 cm depth. This was twice as much as at the Vesicomidae site V1 ( $0.99 \times 10^7 \text{ cm}^{-2}$ ) and ten times more than at the microbial mat site ( $0.16 \times 10^7 \text{ cm}^{-2}$ ). At all sites except the Vest site the surface sediments did not contain aggregates. The reason could be the depletion of methane in the surface layers and/or the oxygenation of the surface sediments. The maximum of aggregate counts in sediments from the microbial mat site occurred between 5 and 6 cm ( $7 \times 10^5 \text{ cm}^{-3}$ ) and between 10 and 11 cm for the Vesicomidae site ( $30 \times 10^5 \text{ cm}^{-3}$ ). The density of aggregates for the Mytilidae sites (M2 and M3) showed a maximum between 13 and 14 cm ( $28 \times 10^5 \text{ cm}^{-3}$ ). At the microbial mat site, aggregates were restricted to the uppermost centimetres (between 3 and 8 cm) whereas at all other sites the aggregate numbers did not decrease with depth. We assume that the bioturbation by the bivalves and hence oxygenation in surface and greater availability of sulfate in subsurface sediments, can explain the deeper distribution of aggregates in M3 and V1 sites compared to the bacterial mat site. Distribution of aggregates in sediments below siboglinid tubeworms (Vest site) showed no trend with sediment depth. Aggregates were detected in the first ten centimetres of sediment. Tubeworms use their roots to take up hydrogen sulphide, which fuels the thiotrophic symbiotic bacteria in their trophosome (Julian et al., 1999; Freytag et al., 2001). Then the tubeworms release sulfate across their roots into the surrounding sediments (Dattagupta et al., 2008), probably creating variable microhabitats for the AOM consortia. Moreover, Vestimentiferan tubeworm assemblages are usually associated with reduced sediments where oxygen is rapidly consumed in the first centimetres (Ritt, Sarrazin et al. in prep).

The densities of microbial aggregates in the pockmark sediments follow the same trend as the methane concentrations in the bottom waters, which were much higher at M3 (23.5  $\mu\text{M}$ ) than at V1 (0.9  $\mu\text{M}$ ) (Table 1). Probably bottom water concentrations are linked to the gradient in total methane fluxes from the seafloor to the hydrosphere, decreasing from the mussel sites to the bacterial mats.

### 3.2. Symbiont sequences

Invertebrate communities at REGAB are dominated by taxa previously reported to harbour chemosynthetic bacterial symbionts, including Mytilidae, Vesicomidae and Siboglinidae (Fisher 1990). Specimen of the mytilid *Bathymodiolus* sp. from the Gabon margin were investigated previously and shown to harbour two phylotypes related to sulphide- and methane-oxidizing symbionts (Duperron et al., 2005). No evidence for methane-oxidizing symbionts was found in the bacterial clone libraries from the two invertebrate species investigated here (Figure 6). A single phylotype, related to the sulphide-oxidizing endosymbiont of *Calyptogena pacifica* was recovered from the vesicomid clam. Vesicomid clams are known to harbour chemoautotrophic symbionts and to colonize sulphide-rich habitats at vents and seeps (Peek et al., 1998). Symbiosis in the tubeworm was also investigated, and preliminary results based on electron micrographs suggested the presence of sulphide-oxidizing symbionts (Andersen et al., 2005) confirmed by the sequence obtained here. Although FISH was not performed on animal samples, the close relationship between sequences recovered and known symbionts suggests that these sequences represent the dominant symbionts, although the presence of additional symbionts remains to be addressed in more specific studies. Symbiosis is known to be of prime importance in the productivity of cold seep communities (Sibuet and Olu, 1998), and recent results confirm that availability of energy in the form of methane and sulfide to the symbiosis is a key parameter for distribution and structure of communities at REGAB (Duperron et al., 2005, Olu et al., 2007a).

### 3.3. Evidence of AOM in REGAB pockmark through molecular survey

DNA was extracted, according to FISH results, from each sample where the AOM consortia were most abundant (Table 1). As the present study was mainly focused on the identification of the microorganisms involved in AOM in the seafloor, only a limited number of clones were sequenced. For the CT10 sample from the thiotrophic mat, the use of nested PCR was necessary to obtain a correct amplification of the archaeal 16S rRNA genes, confirming the low abundance of *Archaea* in this sample. From the Mat sample, we sequenced 78 clones from the bacterial library and 39 from the archaeal one. For each other sediment sample, 25 to 30 clones were sequenced both for *Archaea* and *Bacteria*. 18 to 20 sequences were suitable to be used for phylogenetic analyses, the others 10 being too short for proper phylogenetic analysis.

#### 3.3.1. Archaeal diversity

In all samples, the archaeal diversity comprise only members of the ANME-1b, ANME-2 (a, b and c) and ANME-3 clusters (Table 2 and Figure 4). For the CT10 sample from the thiotrophic mat, nearly half of the sequences obtained were related to ANME-2a and formed two clusters closely associated to uncultured *Methanosarcinales* (> 98% similarity) and to clones retrieved at Kazan mud volcano (ANME-2ab). The other sequences are related to the ANME-3 group, forming a cluster closely related to environmental sequences from methane-rich sites such as Eel River Basin. The seawater overlying the mytilid sites M2 and M3 displayed the highest methane concentration (Table 1), corresponding to higher densities of aggregates in the seafloor. As encountered in other studies, few archaeal sequences were retrieved (Inagaki et al., 2004) despite the high abundance of aggregates according to FISH results (this study, Lösekann et al., 2007). All archaeal sequences were affiliated to the ANME-2c and ANME-3 groups. For the Vesicomidae and Siboglinidae fields, few archaeal sequences were retrieved and these were also all affiliated to ANME clusters.

This relative low archaeal diversity is to be compared to other methane rich environments such as Kazan mud volcano (Heijs et al., 2007), Haakon Mosby mud volcano (Niemann et al., 2006), Black-Sea (Treude et al., 2005, Knittel et al., 2005), Hydrate Ridge (Knittel et al., 2005), Eel River basin (Orphan et al., 2001), Ryukyu Arc seeps (Inagaki et al., 2004) and Gulf of Mexico (Lloyd et al., 2006). In all these methane-rich environments, archaeal diversity is clearly dominated by ANME members but in most of these sites, sequences related to

other Euryarcheota (e.g. *Thermoplasmatales*) or *Crenarcheota* (marine benthic group B, C) were also found, which is not the case here (Knittel et al. 2005). However, this apparently low diversity may also be related to the low number of sequences analysed for this preliminary study.

### 3.3.2. Bacterial diversity

The bacterial diversity is apparently higher than the archaeal one (Table 2 and Figure 5 and 6), as described for the other environments cited above. Nevertheless, the sequences obtained for each sample are all affiliated to a small number of phylogenetic groups.

Twenty six sequences were affiliated to the Epsilonproteobacteria cluster (Figure 5), most of them being closely related to sequences retrieved from deep-sea sediments (NKB sequences) and to methane-rich environments. Some sequences are related to environmental clones associated to chemosynthetic deep-sea marine invertebrates and could thus represent sulfur oxidizers.

Few sequences (CT10 and CT2) were related to the Alphaproteobacteria, *Agrobacterium-Halophaga* division (>98% similarity), showing similar results to those reported from Kazan Mud volcano (Heijs et al., 2007) and Eel River basin (Orphan et al., 2001). This group, closely related to *Hyphomicrobium* sequences, could represent aerobic methylotroph species.

Fourteen sequences were related to the Gammaproteobacteria and three of them were related to mytilid-associated methanotrophic endosymbiont sequences (>98%). This result could suggest the occurrence of free-living, close relatives of mussel symbionts in the environment. A more general survey of methanotrophs is needed, including a complete phylogenetic analysis of the first centimetres where lipid signatures indicate the presence of aerobic methanotrophs (Bouloubassi pers. com.).

Relatives of the Deltaproteobacteria dominated the bacterial clone libraries with 42 sequences (Figure 6). Most of the clones were related to the DSS subgroup (*Desulfosarcina/Desulfococcus* relatives) to which the bacterial partners in AOM belong. Some sequences were closely related to the DBB subdivision (*Desulfobulbus* relatives), which also includes bacteria involved in AOM associated with members of the ANME3 archaeal group (Niemann et al., 2006).

Other clones were related to clusters usually encountered in deep-sea sediments such as Bacilliales (8 clones), *Verrucomicrobiales* (7 clones), Green non sulphur (16 clones), *Actinobacteria* (21 clones), and the *Cytophaga – Flexibacter – Bacteroidetes* (CFB) (19 clones). These clusters are generally supposed to be heterotrophic and present quite versatile groups found in many environments (Alain et al., 2002). In the present study, our clone sequences are mostly related to environmental sequences originated from methane and hydrocarbon-rich areas (Kazan mud volcano, Gulf of Mexico or Guaymas Basin).

Members of the JS1 group (8 clones) were also retrieved under the mat sample and from the mytilid field. Representative sequences of the JS1 group have previously been found in methane-rich subsurface sediments where gas hydrates are present (Webster et al., 2004). This suggests that some members of the JS1 division may be associated with methanogenic consortia. JS1 inhabit strictly anoxic organic-rich environments (Webster et al., 2006).

Results from this preliminary study clearly indicate that key microorganisms mediating the anaerobic oxidation of methane occur below each type of faunal assemblage (thiotrophic mat, Mytilidae, Vesicomidae) on the REGAB pockmark. Too few sequences were recovered from the Siboglinidae fields to draw any conclusion, but these preliminary results seem to confirm also the presence of AOM consortia under the Siboglinidae fields. Interestingly, all known marine AOM-mediating groups were present in clone libraries of the pockmark sediments. Below the *Bathymodiolus* field, aggregates were most abundant but the archaeal PCR amplification yielded poor results. The low number of ANME-related sequences in sediments characterized by high methane concentration such as observed here was previously reported (Inagaki et al., 2004, Lösekann et al., 2007) and could be due to PCR inhibition or biases (Lösekann et al., 2007). In the absence of quantitative data regarding

abundances, the relative importance of each group cannot be assessed. Nevertheless, the occurrence of strongly  $^{13}\text{C}$ -depleted archaeal lipids (such as archaeol, sn-2 hydroxyarchaeol, crocetane and PMI), clearly points out to substantial biomass of anaerobic methanotrophic archaea (Bouloubassi, this volume). Their composition and relative abundance indicate a methanotrophic archaeal community dominated by ANME-2 members. In addition, the co-occurrence of  $^{13}\text{C}$ -depleted monoalkyl glycerol ethers (MAGE) indicates the presence of sulphate reducing bacteria among the *Desulfosarcinales* group, also implicated in AOM with ANME-2. The downcore distribution of archaeal and bacterial (SRB) lipids provides evidence for dominant anaerobic oxidation of methane (AOM) and concomitant sulphate reduction below the upper 2 cm layer of the sediments (Bouloubassi pers. comm.). Within the surface sediment layer, AOM-related biomarkers show low concentrations, whereas, in contrast, amounts of  $^{13}\text{C}$  depleted diploptene are highest, most likely reflecting aerobic methanotrophy in the zone where oxygen penetrates from the overlying bottom water (Bouloubassi this volume).

#### 4. Conclusion

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The sediments collected from four main types of chemosynthetic habitats of the REGAB pockmark, characterized by aggregations of Mytilidae, Vesicomidae, Siboglinidae and thiotrophic bacterial mats, have been found to contain abundant AOM consortia. The ANME-2/DSS shell-type aggregates were concentrated a few centimetres below the surface (3-5 cm). Highest densities of AOM consortia were discovered at the Mytilidae sites, which also showed the highest methane concentrations in the bottom waters.

Due to the necessary use of 30 PCR cycles, and the investigation of limited numbers of clones, we were not able to give an exhaustive description of the diversity of microbial communities present on the REGAB site. Nevertheless, representatives of all ANME groups were found at REGAB, comprising all recovered archaeal sequences. Bacterial diversity in clone libraries from the REGAB sediment was relatively high for the number of clones examined and included sequences related to aerobic methanotrophs, sulphide oxidizing and sulphate reducing bacteria, as well as diverse heterotrophs. The bacterial 16S rRNA genes associated with the 3 main types of chemosynthetic invertebrate symbioses of mytilids, vesicomids and siboglinids were closely related to known thiotrophic symbionts of the respective groups. Only the mytilids contained a second symbiont related to aerobic methanotrophs. This indicates that the rich chemosynthetic communities of the REGAB pockmark are mostly dependent on AOM as the sulphide generating process in this ecosystem.

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## Figures

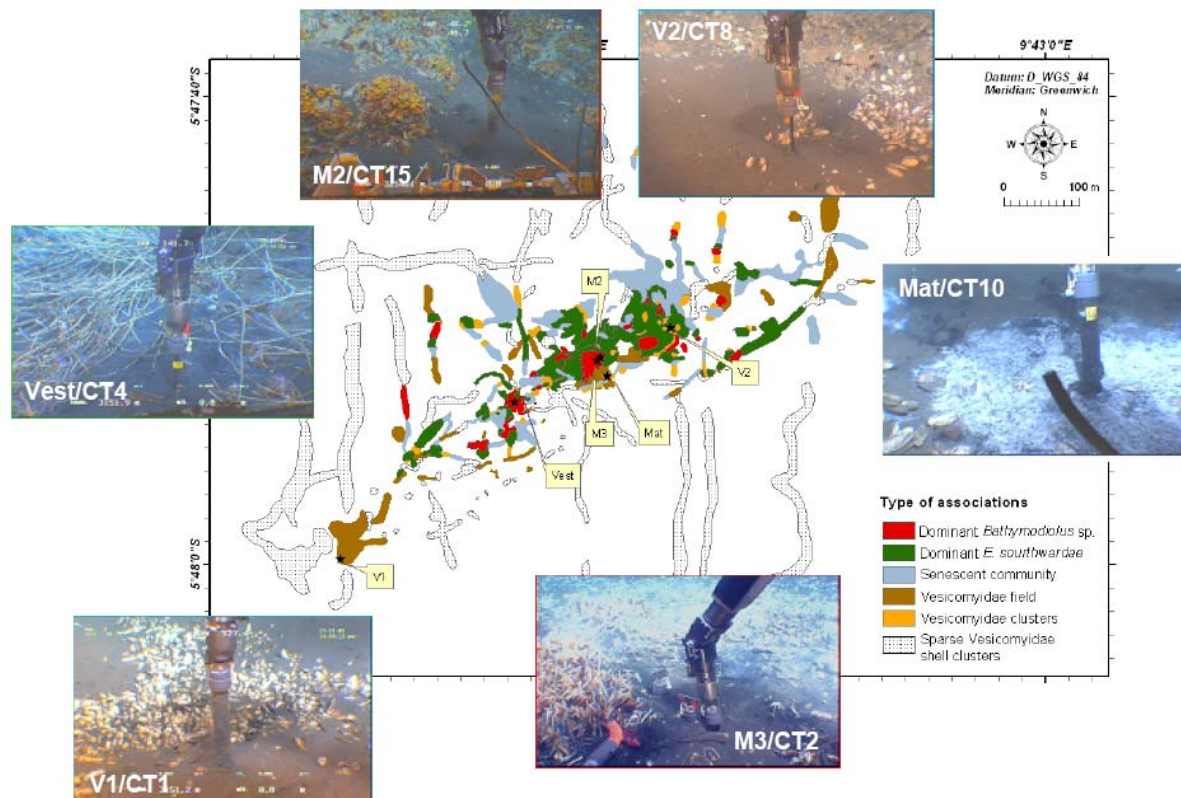


Figure 1 : Map of the REGAB site (adapted from the Olu-Le Roy et al. 2007a) and sediment sampling sites.

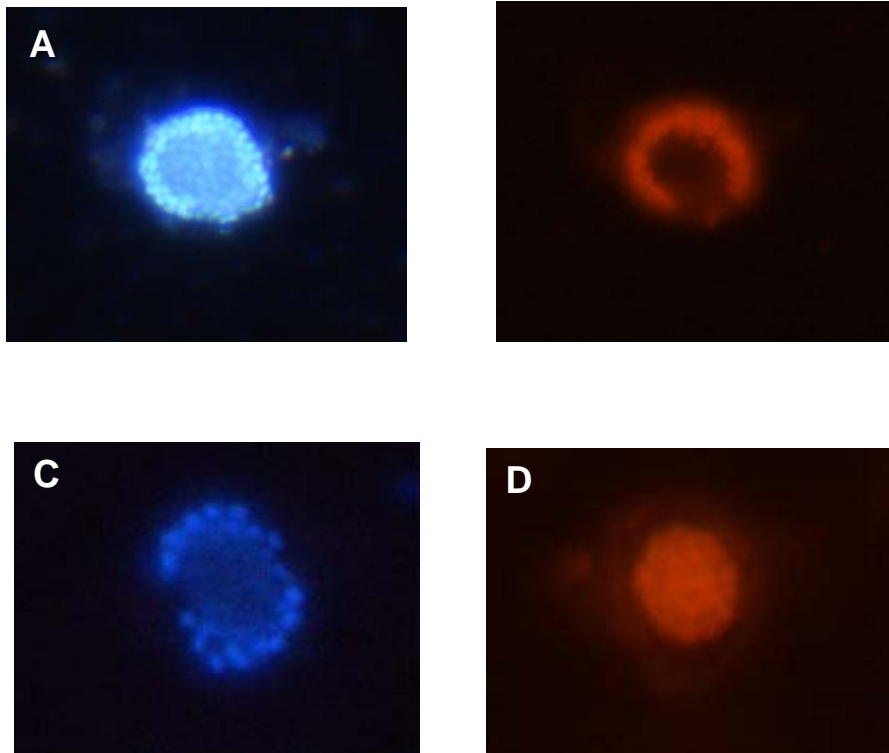


Figure 2 : Consortia of ANME-2 archaea and sulfate-reducing bacteria DSS in sediments of REGAB visualized by fluorescence microscopy and FISH. (A) DAPI staining of aggregate. (B) Corresponding FISH image of same aggregate with probe DSS658 specific for sulfate-reducing bacteria (*Desulfosarcina* /*Desulfococcus*). (C) DAPI staining of aggregate. (D) Corresponding FISH image of same aggregate with probe EeIMS932 specific for ANME-2

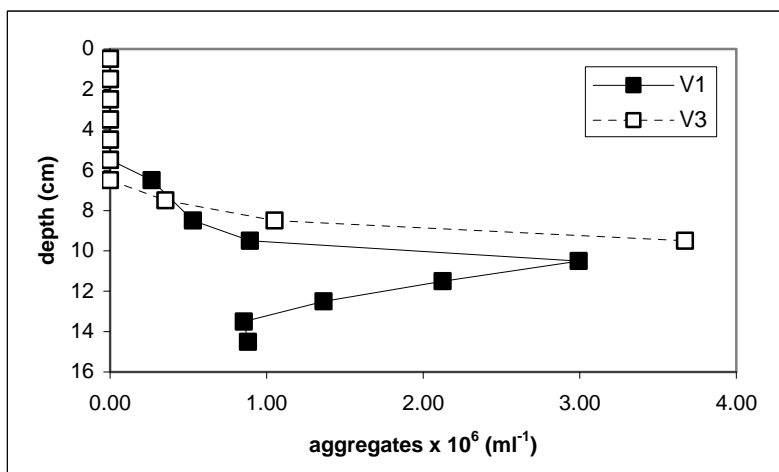
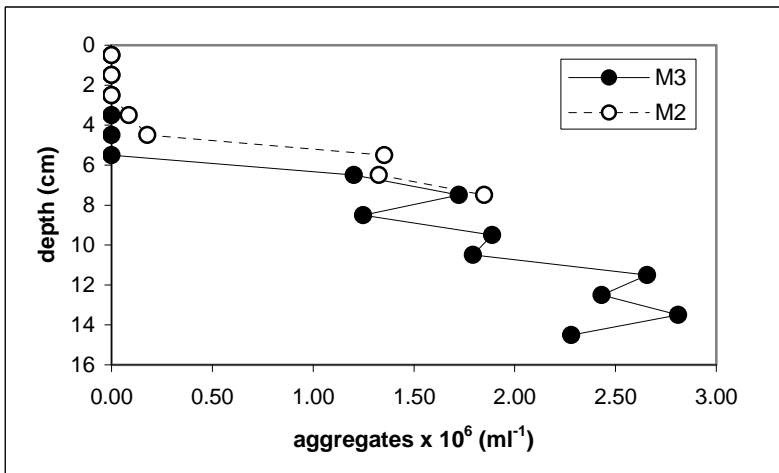
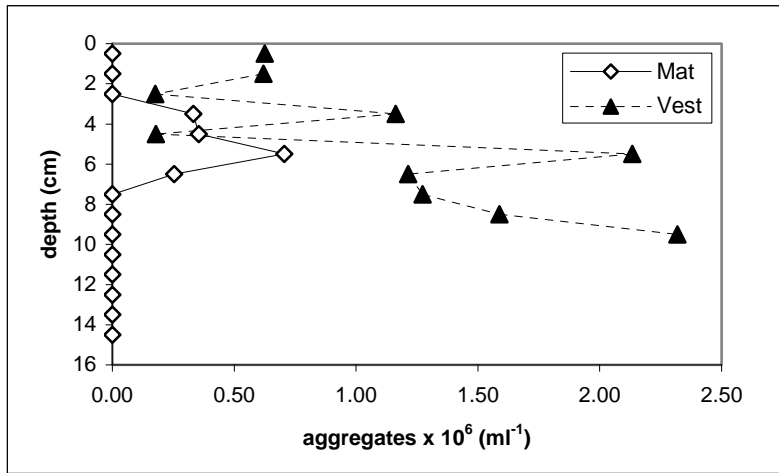


Figure 3 : Detection and quantification of AOM consortia in REGAB sediments by FISH (archaeal probe EelMS932, and bacterial probe DSS658)

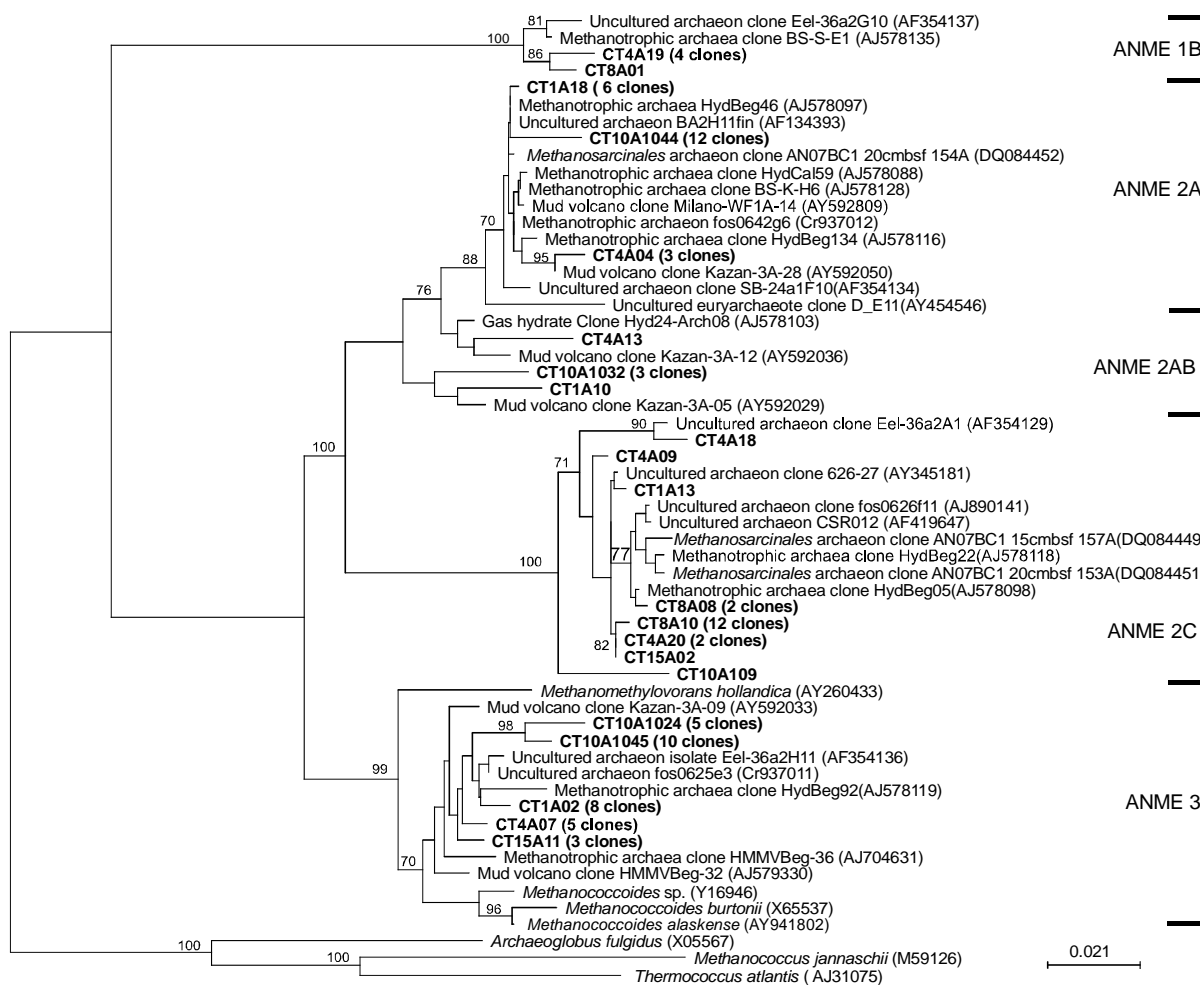


Figure 4 : Phylogenetic trees of Archaea obtained using Neighbour-joining analysis with bootstrap resampling (1000 replicates), based on 558 base pairs. Topologies were confirmed with Maximum Parsimony and Maximum Likelihood. Bootstraps values are indicated on nodes above 70%. The outgroup sequence used was *Thermococcus atlantis*.

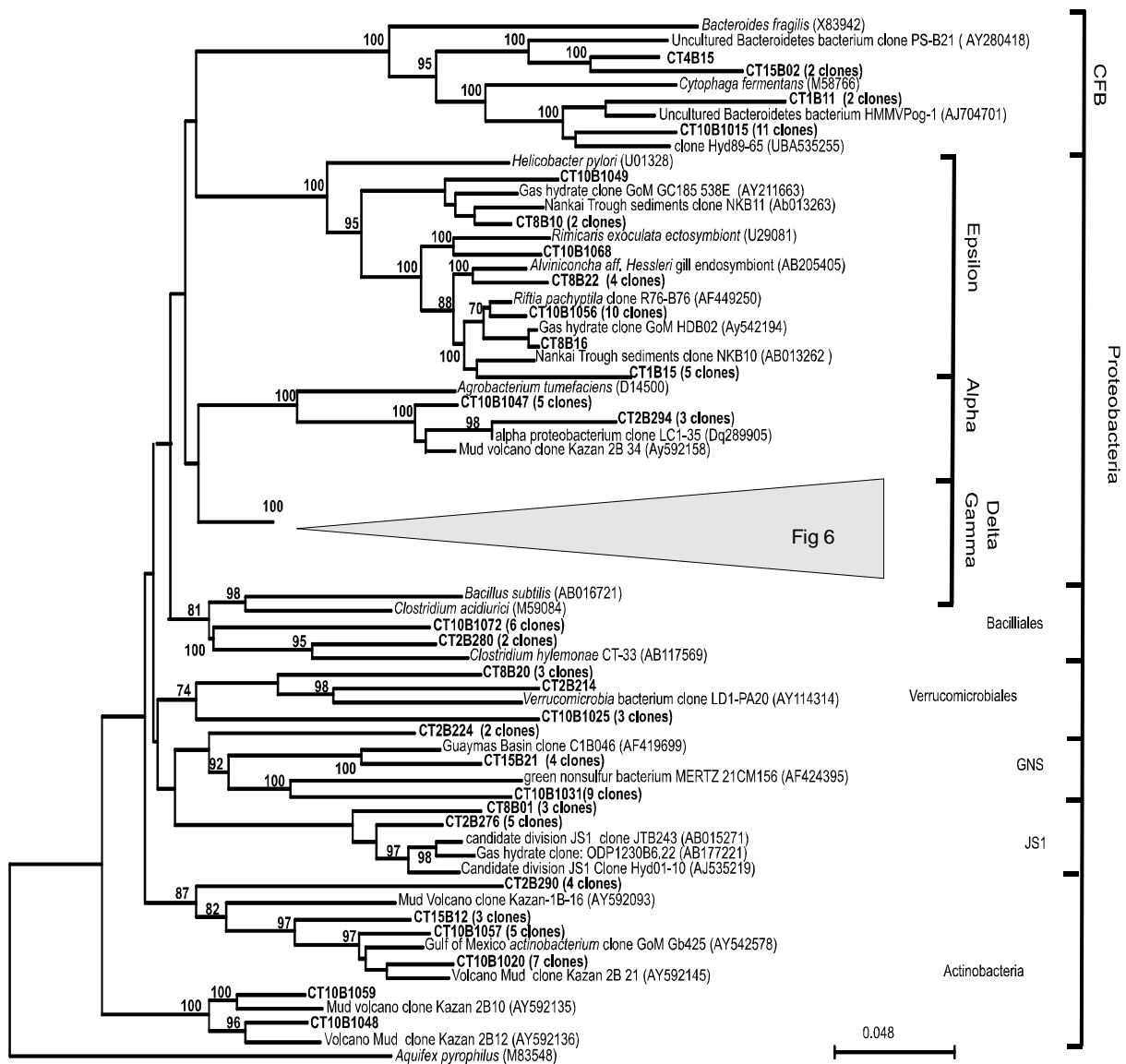
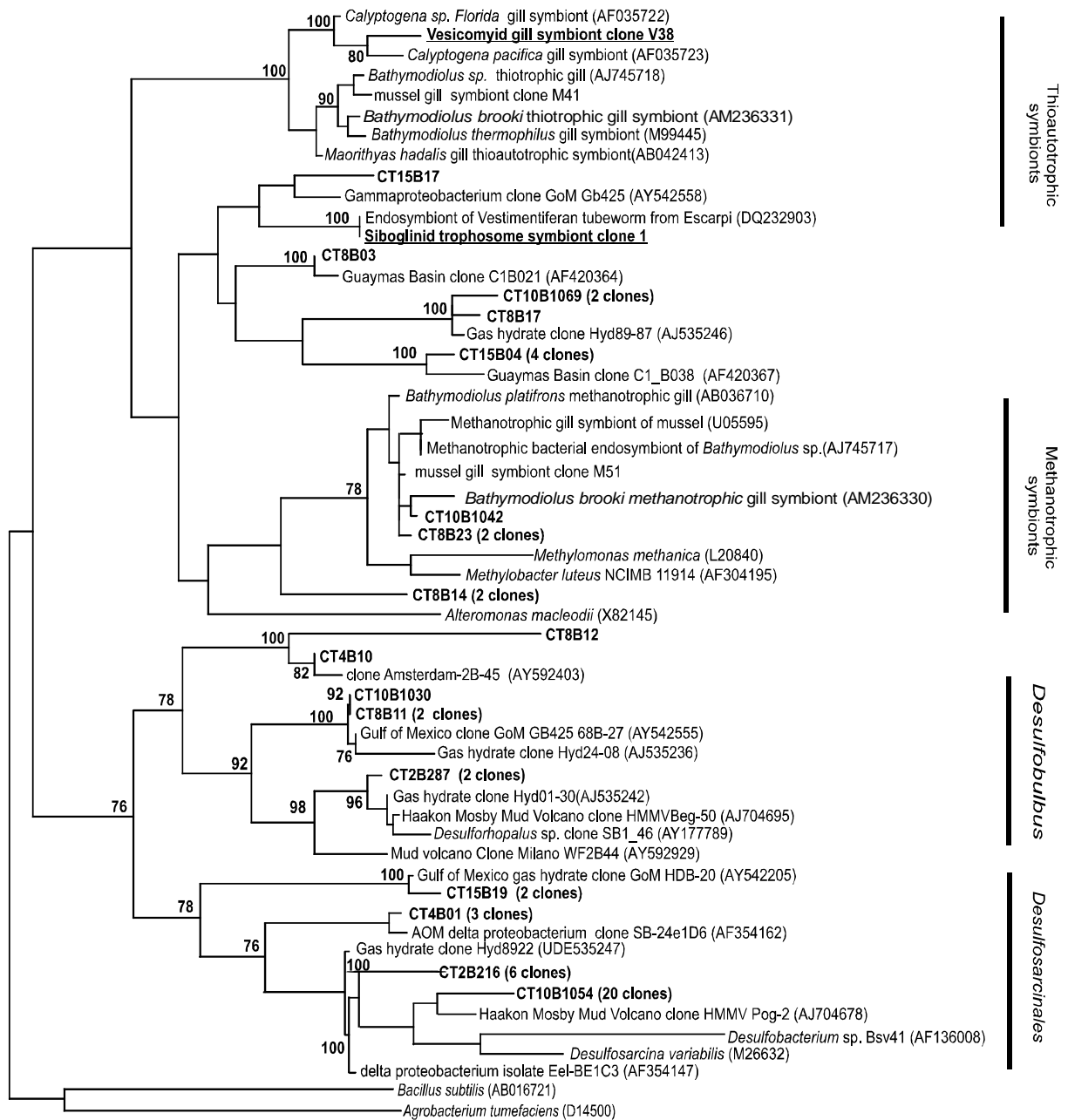


Figure 5 : Phylogenetic trees obtained using Neighbor-joining analysis with bootstrap resampling (1000 replicates), based on 812 base pairs. Topologies were confirmed with Maximum Parsimony and Maximum Likelihood. Bootstraps values are indicated on nodes above 70%. The outgroup sequence used was *Aquifex pyrophilus*.





0.035

Figure 6 : Phylogenetic trees obtained using Maximum Likelihood analysis with bootstrap resampling (100 replicates), based on 561 base pairs. Bootstraps values are indicated on nodes above 70%. The outgroup sequences used were *Agrobacterium tumefaciens* and *Bacillus subtilis*.

## Tables

**Table 1:** Sampling sites and methane concentrations measured above sediment-water interface. DNA was extracted on sub-sample of each core and the extracted sample is indicated in centimeter below the surface.

Sampling site	Sample name	Fauna	Methane concentration ( $\mu\text{M}$ )	DNA extraction cm
M2	CT15	<i>Bathymodiolus</i> sp.	33.7	5-8
M3	CT2	<i>Bathymodiolus</i> sp.	11.8	10-15
V1	CT1	Vesicomidae	0.9	5-10
V3	CT8	Vesicomidae	4.4	5-10
Vest	CT4	<i>Escarpia</i> sp.	0.7	5-10
Mat	CT10	Bacterial mat	n. d.	5-10

n.d.: not determined

**Table 2 :** Closest 16S rRNA gene sequence matches to our clone type sequences using the NCBI BLASTN search tool .

Sample origin	Sequence type	Phylogenetic affiliation	Closest match by BLASTN search (accession number)	% of similarity	Number of related clone (similarity above 97%)
Microbial Mat (CT10) Partial sequences 1000 bp for bacteria 500 bp for archaea	CT10B1049	Epsilonproteobacteria	Uncultured clone NKB11 (AB011363)	98	1
	CT10B1068		Uncultured clone C1_B005 (AF420358)	98	1
	CT10B1056		Uncultured clone R76-B76 (AF449250)	98	10
	CT10B1046	Alphaproteobacteria	Clone LC1-35 (DQ289905)	99	5
	CT10B1047		Clone Hyd89-87 (AJ535246)	98	2
	CT10B1069	Gammaproteobacteria	Clone Milano Mud Volcano (AY592847)	98	1
	CT10B1042		Clone Hyd89-22 (AJ535247)	98	20
	CT10B1054	Deltaproteobacteria	Clone GoMGC2324463 (AM745212)	98	1
	CT10B1030		Clone HMMVPog-1 (AJ704701)	98	11
	CT10B1040	CFB	Clone HMMVPog-1 (AJ704701)	98	6
CT10B1030	Bacilliales	<i>Clostridium</i> sp. Kas303 (AB114244)	92	3	
CT10B1030	Verrucomicrobiales	Chlamydial symbiont (EF177461)	99	9	

	CT10B101 5	GNS	Clone GoM161-bac49 (AM745150)	98	5
		Actinobacteria	Clone Kazan 2B 21 (AY592145)	99	7
	CT10B107 2		Clone GoM Gb425 (AY542578)	99	1
			Clone Kazan 2B 10 (AY592135)		1
			Clone Kazan 2B 12 (AY592136)	99	12
	CT10B102 5	ANME-2	Clone HydBeg46 (AJ578097)	97	3
			Clone Kazan 3A 05 (Ay592029)	98	1
	CT10B103 1		Clone HydBeg22 (AJ578118)	98	5
		ANME-3	Clone fos0625e3 (CR937011)	99	10
	CT10B105 7		Clone fos0625e3 (CR937011)		
	CT10B102 0				
	CT10B105 9				
	CT10B104 8				
	CT10A104 4				
	CT10A103 2				
	CT10A109				
	CT10A102 4				
	CT10A104 5				
Mytilid fields (CT15 partial sequences 1000 bp, CT2 full length sequences)	CT2B294	Alphaproteobacteria	Clone Kazan 2B 34 (AY592158)	96	3
	CT15B04	Gammaproteobacteria	Guaymas clone C1-B038 (AF420367)	94	4
	CT15B17		Clone GoM Gb425 (AY542558)	95	1
	CT15B19	Deltaproteobacteria	Clone GoM HDB-20 (AY542210)	95	2
	CT2B216		Clone Eel-36e1G12(AF354163)	99	6
	CT2B287		Clone Hyd01-30 (AJ535242)	98	2
	CT15B02	CFB	Clone GoM161-Bac19 (AM745134)	97	2
	CT2B280	Bacilliales	Clone RL199-aaj41a08 (DQ793227)	95	1
	CT2B214	Verrucomicrobiales	Clone LD1-PA20 (AY114314)	89	1
	CT2B224	GNS	Clone C5LK53 (AM086136)	97	2
	CT15B21		Clone C1B046 (AF419699)	96	1
	CT2B276	JS1	Clone Hyd01-10 (AJ535219)	92	5
	CT2B290	Actinobacteria	Clone B22 (EF89984)	91	4
CT15B12	Clone GoM GB425 (AY542578)		89	3	
CT15A02	ANME-2	Clone fos0644c1 (CR937009)	95	1	
CT15A11	ANME-3	Clone Kazan 3A 09 (AY592033)	98	3	
Vesicomid fields (CT1 full length sequences, CT8 partial sequences 1000 bp)	CT8B10	Epsilonproteobacteria	clone GoM GC185 538E (AY211663)	96	2
	CT8B22		Uncultured clone NKB10 (AB013262)	95	4
	CT8B16		clone GoM HDB-02 (AY542194)	98	1
	CT1B15		clone GoM HDB-02 (AY542194)	96	5
	CT8B17		Clone hyd89-87 (AJ535246)	97	1
	CT8B14	Gammaproteobacteria	Clone <i>Bathymodiolus</i> sp (AB036710)	91	2
	CT8B03		Clone C1B021 (AF420364)	98	1

	CT8B23		Clone <i>Bathymodiolus</i> sp (AJ745717)	98	2
	CT8B11		Clone GoM GB425 (AY542555)	96	8
	CT8B12		Clone 2B-45 (AY592403)	89	1
	CT1B11	Deltaproteobacteria	Clone HMMVPog-1 (AJ704701)	97	2
	CT8B20	CFB	Clone Kazan 3B 36 (AY592202)	86	3
	CT8B01	Verrucomicrobiales	Clone JTB243 (AB015271)	98	3
	CT8A01	JS1	Clone Eel-36a2G10 (AF354137)	97	1
	CT1A10	ANME-1	Clone Kazan 3A-05 (AY592029)	97	1
	CT1A13		Clone HydBeg 05 (AJ578098)	97	1
	CT1A18	ANME-2	Clone fos0642g6 (CR9370012)	98	8
	CT8A10		Clone AN07BC1 (DQ084451)	95	12
	CT8A08		Clone AN07BC1 (DQ084451)	96	2
	CT1A02		Clone fos0625e3 (CR9370011)	98	8
		ANME-3			
Siboglinid field (CT4, partial sequences 1000 bp)	CT4B01	Deltaproteobacteria	Clone SB-24e1D6 (AF354162)	95	3
	CT4B10		Clone 2B45 (AY592403)	98	1
	CT4B15	CFB	Clone GoM161 B19 (AM745134)	98	1
	CT4A19	ANME-1	Clone GoM140A80	98	4
	CT4A04	ANME-2	Clone fos0642g6 (CR937012)	98	3
	CT4A13		Clone Kazan 3A-12 (AY592036)	94	1
	CT4A18		Clone Eel-36a2A1 (AF354129)	97	1
	CT4A09		Clone fos0644c1 (CR937009)	92	1
	CT4A20		Clone AN07BC1 (DQ084451)	98	2
	CT4A07	ANME-3	Clone fos0625e3 (CR9370011)	97	5